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**FACULTY OF SCIENCE AND TECHNOLOGY**

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## Abstract

Southeast directed fluvio-deltaic systems transported across Svalbard during Early Cretaceous are age correlated to southeast progradation deposits in the Fingerdjupet Subbasin. Both local (few km) and distal (>300 km) source regions has been suggested for this system. The recent exploration well 7321/4-1 drilled on the flank of the basin encountered sandstones in the Lower Cretaceous interval. This well data has not been implemented in the previous studies.

This study uses detrital zircon U/Pb geochronology from well 7321 / 4-1 and 7322 / 7-1 in the Fingerdjupet Subbasin to interpret the source of coarse grain sediments. Palynological analyses, well and seismic data are integrated to interpret the distribution of the sandstones encountered in recent well 7321 / 4-1. The aim of this study is to evaluate the provenance of Lower Cretaceous sandstones in encountered by well 7321 / 4-1 and the implications for reservoir development in the Fingerdjupet Subbasin.

Two seismic units (SU1 and SU2) is defined based on seismic downlap terminations and seismically guided well correlations. The SU1 shows wedge-shaped packages thickening towards fault in the central part of the basin. The SU2 consist of the encountered by well 7321/4-1 and is characterized by shelf-edge clinoforms that prograded southeast into the basin. The dominant detrital zircon age populations in SU2 unit from well 7321 / 4-1 is: (1) 2.6-2.75, (2) 1.7-1.5 Ga and (3) 1.2-1 Ga. Only a few Detrital Zircon ages were measured in well 7322 / 7-1 and it was thus not possible to compare the result with this well. By comparing Detrital Zircon Ages from the SU2 unit with Lower Cretaceous formations on Svalbard and in the Barents shelf, it became clear that both regions contain similar dominant age distribution.

The source of sediment is interpreted to originate from north Greenland and/or Arctic Canada based on the similar detrital age distribution and the southwestern progradation direction of the clinoforms. Seismic interpretation suggests that the well 7321 / 4-1 penetrates the topsets segment of the shelf-edge clinoforms, which interpreted as a potential sandy system in the topsets.

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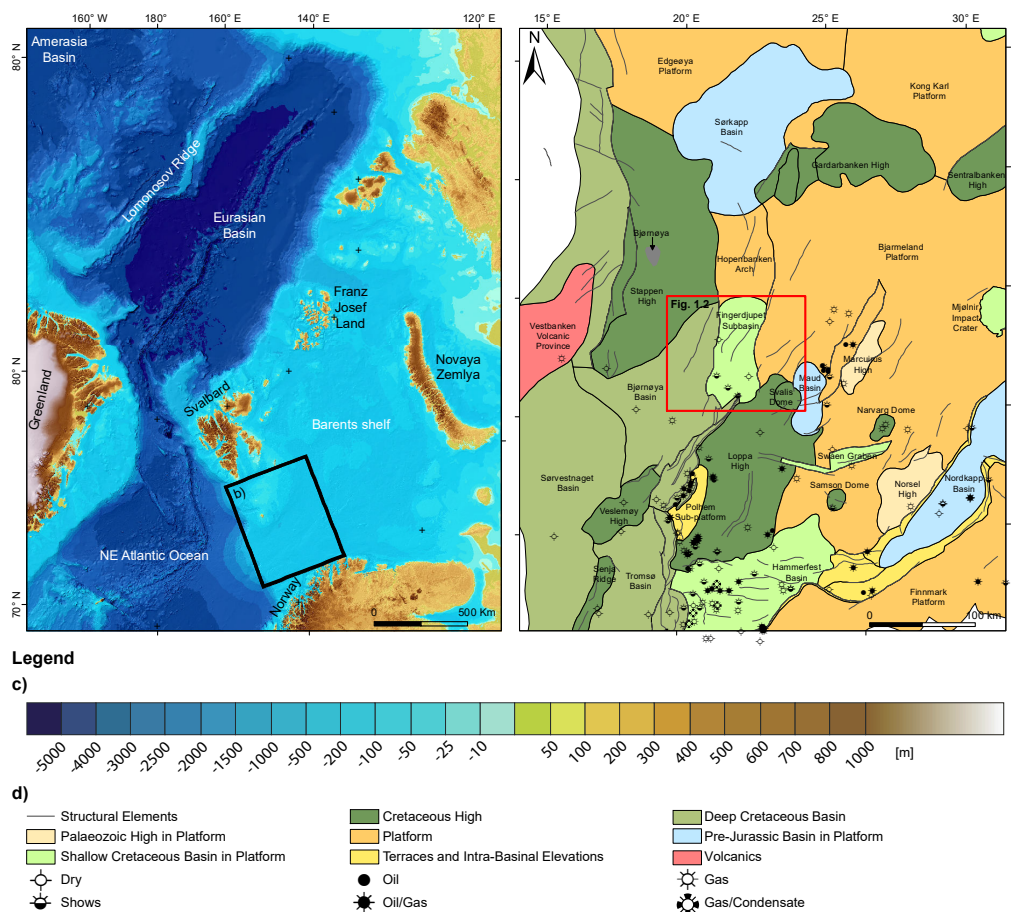
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# 1 INTRODUCTION

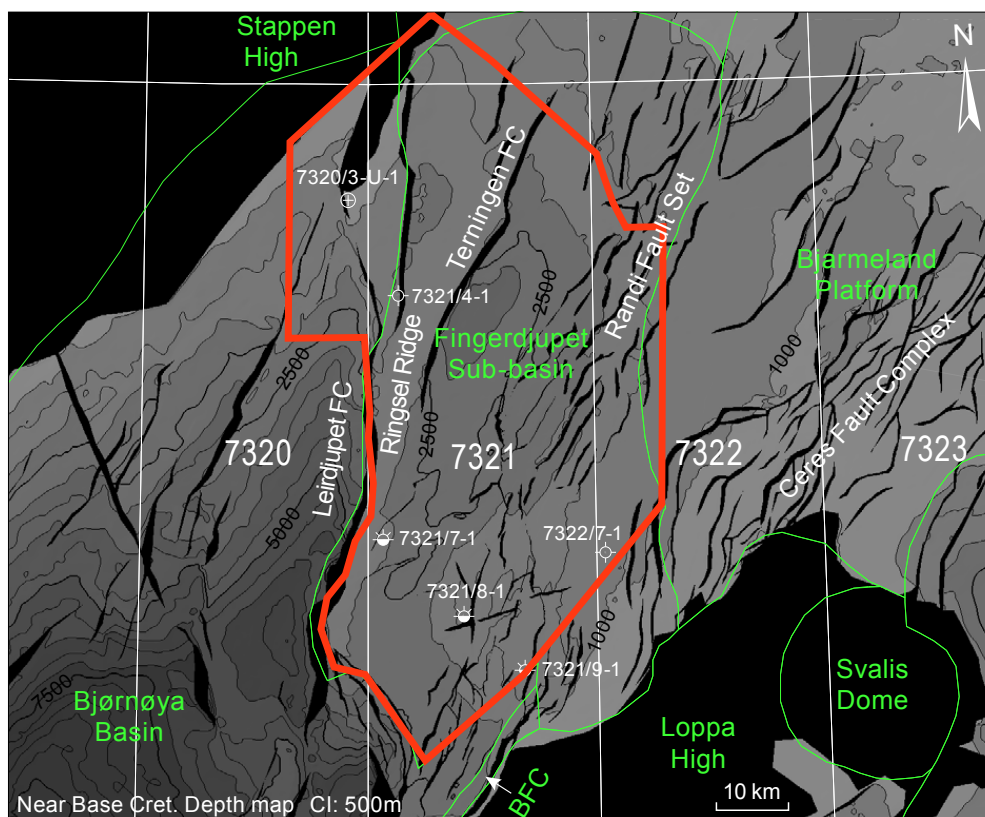
The western Barents shelf borders the mainland of northern Norway in the south, Northeast Atlantic Ocean in the west and is exposed in the Svalbard archipelago in the north-western corner (Fig. 1.1a) (Smelror et al., 2009). The northern margin of the Barents shelf is interpreted to have been uplifted during Cretaceous time (e.g. Maher, 2001; Polteau, 2015). Outcrop studies on Svalbard suggests that a northern tilt of the shelf triggered a southeast evolving paleoshoreline during Barremian (Early Cretaceous, Steel & Worsley, 1984; Gjeldberg & Steel, 1995; Worsley, 2008; Midtkandal and Nystuen, 2009). This regressive system consists of sandy fluvio-deltaic deposits visible in outcrops across Svalbard and are interpreted to extend into the south-western Barents shelf (Fig. 1.1b, e.g. Grundvåg et al., 2017; Midtkandal et al., 2019).



**Fig. 1.1 Structural elements in the Arctic and the southwestern Barents shelf** a) Map view of topographic and bathymetric features in the Arctic region after Jakobsson et al., (2012). The coordinate system used is the WGS 1984 with polar stereographic projection. The square marks the map area in Fig. 1b. b) Regional map of the southwestern Barents shelf showing the structural elements based on NPD (2020). The square marks the map area in Fig. 1.2. c) Color coding used in Fig. 1a and the corresponding meters above and below mean sea level. d) Color coding and well symbols used in Fig. 1b.



The terminal deposits of the regressive system observed on Svalbard is age correlated to a southeast directed clinoform system in the Fingerdjupet Subbasin and on the western Bjarmeland Platform (Grundvåg et al., 2017; Midtkandal et al., 2019). Recent exploration well 7322/7-1 targeted topsets of southeast prograding, delta-scale (25-80 m), clinoforms in the Fingerdjupet Subbasin (Fig. 1.2 , Bryn et al. 2019). The well penetrated the topsets less than one kilometer away from the delta-scale foresets which were interpreted to be of coarse-grained lithology based on the high angle foresets (10–12°) and modern quantitative clinoform analysis (Bryn et al., 2019, Patruno & Helland-Hansen, 2018). The reservoir proved to be gas-charged siltstones of Barremian age with a non-commercial volume. Bryn et al. (2019) studied the reservoir potential of clinoforms in the Fingerdjupet Subbasin and argued that the high angle foresets near the well may still be sand-prone. Thus, the reservoir potential of the Barremian clinoform system in the Fingerdjupet Subbasin remains unproven. Furthermore, Bryn et al. (2019) concluded that linking a coarse-grained source area to clinoforms of coarse grain character can be an important first step towards predicting the reservoir potential of the Barremian clinoforms.



**Legend**  
— Structural Elements  
— Study Area

**Fig. 1.2 Study Area.** Depth structure map of a reflector near the Base Cretaceous Unconformity (edited from Bryn et al. (2019)). The study area is illustrated in red color. For location of map see Fig. 1.1b.

Three source regions have been suggested to supply coarse grain sediments to the southeast directed clinoform system observed in the Barents shelf. (1) A western region corresponding to north-eastern Greenland (Grundvåg et al., 2017), (2) a north-western region corresponding to Crockerland (Paleohigh in northern Canada, Midtkandal et al., 2019) and (3) the Stappen High, located at the western margin of the south-western Barents shelf (Fig. 1.1b, Bryn et al., 2019). Hence, the source of sediments is poorly constrained.

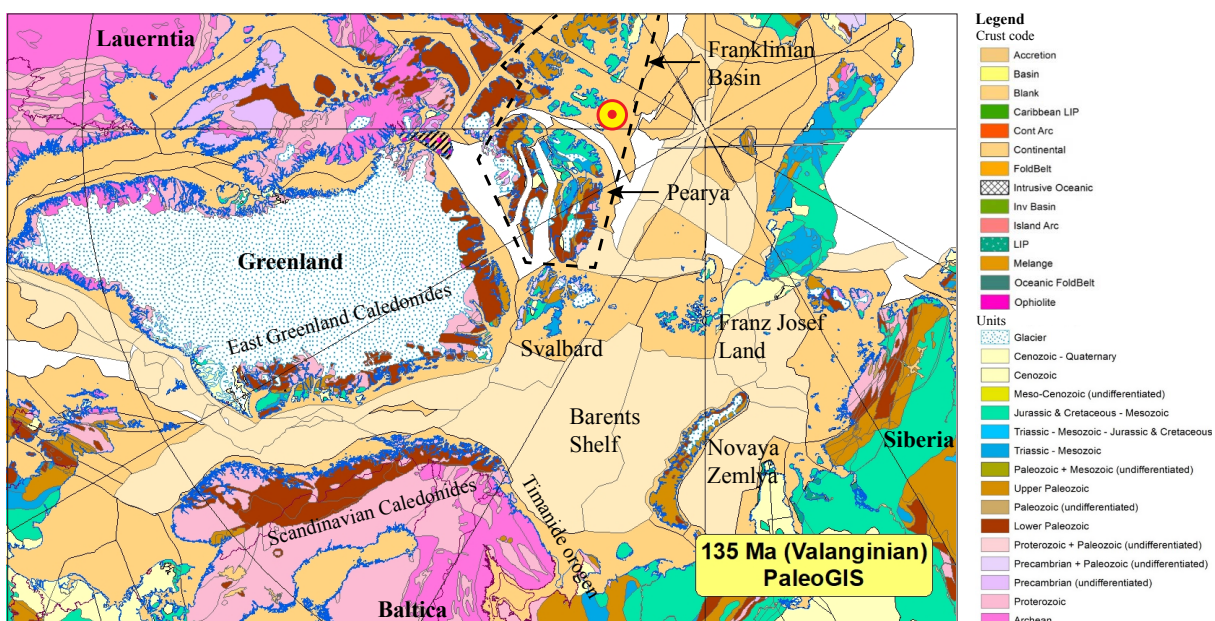
Recent exploration Well 7321/4-1 (2018) was drilled on the Ringsel Ridge, a horst structure defined by the Leirdjupet Fault Complex (LFC) and the Terningen Fault Complex, in the Fingerdjupet Subbasin (Fig. 1.2). This well targeted Jurassic and Triassic sandstones but also encountered sandstones in the Lower Cretaceous interval (B.K. Bryn. 2019. Pers. Comm.). This well data has not been implemented by previous studies.

This study uses biostratigraphic data and detrital zircon U/Pb geochronology (provided by GEUS) to interpret the age and origin of the Lower Cretaceous sandstones in well 7321/4-1 and the siltstones in well 7322/7-1. Biostratigraphy and petrophysical logs from the southern wells (7321/7-1, 7321/8-1 and 7321/9-1, Fig. 1.2) and seismic data is used to interpret the distribution of the sandstones encountered in well 7321/4-1. The aim of this study is to evaluate the provenance of Lower Cretaceous sandstones on the Ringsel Ridge and the implications for reservoir development in the Fingerdjupet Subbasin.

## 2 GEOLOGICAL SETTING

### 2.1 Potential Source Areas

During the Early Cretaceous a widespread epicontinental sea (Boreal Sea) covered the present day Barents shelf, parts of northeast Greenland and Sverdrup Basin. The northwest hinterlands facing the Barents shelf were Svalbard, Pearya and Laurentia (Fig. 2.1). The Baltic plate was situated along the south-western margin and Siberia in the south-eastern margin of the Barents shelf ( e.g. Torsvik et al., 2002; Torsvik et al., 2012, Fig. 2.1.).



**Fig. 2.1 Plate Reconstruction (Valanginian, 135Ma)** Plate tectonic reconstruction of the Barents shelf and basement geology (figure modified from LOCRA Final Report. (2017).

The crystalline basement of the northeast Laurentia interior consists of Archean cratons, reworked Archean rocks and intervening Palaeoproterozoic orogenic belts (Fig. 2.1, 2006). Towards the boarder of Franklinian Basin (Fig. 2.1), the basement is of Archean to early Proterozoic age with intrusive igneous rocks of mafic to felsic composition, in addition to Proterozoic sedimentary and volcanic rocks (Trettin, 1991). Large parts of the Laurentia basement in the northern part are covered by Neoproterozoic to Devonian siliclastic and carbonate rocks in the Franklinian Basin (Fig. 2.1), interpreted to represent a passive margin succession (Dewing et al., 2008).

The Pearya Terrane (Fig. 2.1) is located in the northernmost region of Ellesmere Island (Canada) and has been characterized as an accreted or exotic terrane(relative to the adjacent terranes) along the northern margin of Laurentia (Trettin, 1991). Accretion against the Franklinian Basin has been inferred

to occur during the middle Palaeozoic Ellesmerian Orogeny (Malone et al., 2017; Piepjohn & Von Gosen, 2017), interpreted as being the equivalent to the Caledonian Orogeny along the North Canada and Greenland margins (Gasser, 2014; Gee, 2015). The basement in the Pearya Terrane is dominated by metasedimentary rocks and granitoid gneiss (Trettin et al., 1991) of Tonian age (972–962 Ma) (Malone et al., 2017). The basement is overlaid by Neoproterozoic to Middle Ordovician metasedimentary and metavolcanic rocks (Fig. 2.1). These deposits also includes igneous intrusions of ultramafic-mafic composition aged 481 Ma and younger felsic intrusions of 462 Ma age. The intrusions were followed by nearly unmetamorphosed volcanic, siliciclastic and carbonate sedimentary rocks of middle Ordovician to late Silurian age (Trettin, 1991).

The northeastern margin of Laurentia is dominated by the 1300 km long East Greenland Caledonides (Fig. 2.1, Gee et al., 2008; Higgins et al., 2008). As a result of the collision of Laurentia and Baltica which culminated during the latest Silurian–early Devonian (Gee et al., 2008). The East Greenland Caledonides can be divided into three segments: northern (80–82°N), central (76–80°N) and southern (70–76°N) (Gasser, 2014). The crystalline basement is dominated by gneisses and metagranitic rocks (2 - 1.85 Ga) which are exposed in the central segment (Kalsbeek et al., 1999; 2008). Caledonian thrust sheets are composed of Paleoproterozoic to Neoproterozoic rocks of mainly siliciclastic and sedimentary carbonate, volcanic and metasedimentary rocks (Fig. 2.1, Higgins et al., 2008). While Lower Paleozoic sediments occupies large parts of western foreland and the upper sheets of the orogenic belt in the northern and southern segments (Smith and Rasmussen, 2008). Caledonian granites is only present in the southern segment of the orogen (Gasser, 2014).

Svalbard (Fig. 1.1a, Fig. 2.1) can be divided into the Eastern (Western Ny Friesland and Nordaustlandet terranes), Northwestern and Southwestern basement terranes (Gee and Teben'kov, 2004). The relative origins of the different terranes are poorly constrained (e.g. Gasser, 2014). The oldest rocks on Svalbard are late Archean to late Paleoproterozoic (c. 2.71 - 1.75 Ga) granitic gneisses found in the Eastern terrane (Western Ny Friesland) (Johansson et al. 1995; Wellman et al., 2001;). Apart from the gneisses, the Eastern terrane is dominated by Mesoproterozoic metasedimentary and metavolcanic rocks, intruded by Tonian and subsequent Silurian aged granitoids (McClelland et al., 2019), overlain by Neoproterozoic to Early Palaeozoic siliciclastic and carbonate sedimentary rocks (Witt-Nilsson et al., 1998; Sandelin et al., 2001). The Northwestern Terrane is dominated by late Mesoproterozoic to Neoproterozoic metasedimentary rocks intruded by Tonian (c. 0.96 Ga) and Silurian (c. 0.42 Ga) aged granitoids (Petterson et al., 2009). The Southwestern Terrane constitutes Mesoproterozoic and Neoproterozoic ortho gneisses (c. 1.2 Ga

and 0.95Ga) (Majka et al., 2014) and metaclastic rocks subjected to metamorphism reaching upper amphibolite facies conditions during Torellian (c. 640 Ma) (e.g., Majka et al 2010) and eclogite facies during Ordovician (e.g., Kosminska et al., 2014). These units are followed by Ordovician to Silurian siliciclastic and carbonate sedimentary rocks with no evidence of Silurian magmatism or high-grade metamorphism (Gasser & Andresen, 2013).

The interior of north-western part of Baltica (Fig. 2.1) is dominated by Archean (3.5 – 2.73 Ga) granitoid–gneiss (Holttta et al, 2012), Paleoproterozoic (2.44 - 1.92 Ga) metavolcanic and metasedimentary rocks (Fig. 2.1, Lahtinen et al., 2010; Koykka et al., 2019). During the latest Neoproterozoic Timanide orogen (610-640 Ma), the northeastern margin of the Baltica experienced subduction and accretion (Pease, 2011). Late Neoproterozoic metaclastic rocks along the orogenic belt and granitoids of same age are common along the Timanide segment in Arctic Russia (Lorenz et al., 2004). The Timanide structures in the northern Norway are overprinted by the Scandinavian Caledonides in the north-western margin of Baltica (Fig. 2.1, Pease et al., 2014). The Scandinavian Caledonides are interpreted to extends into the Barents Sea region (e.g. Corfu et al., 2014). Along the length of Scandinavian Caledonides, Archaean to Neoproterozoic rocks of Baltica are followed by metaigneous and metasedimentary rocks of Neoproterozoic and younger age, interpreted to represent the thrust sheets of the orogenic belt. During the closure of the Iapetus Ocean (paleocean between Laurentia and Baltica), the Baltic plate is interpreted to be subducted below Laurentia, stacking the Baltica crust, followed by Iapetan oceanic crust and Laurentia crust (Roberts and Gee, 1985; Corfu et al., 2014). The subduction phase is constrained to Mid-Silurian and the continent–continent collision phase, referred to as the Scandian Orogeny, to mid-Silurian to late Devonian (430-380 Ma, Corfu et al., 2014).

The Uralian Orogeny closed the Uralian Ocean along the eastern margin of Baltica (Fig. 2.1, Arctic Russia) during the Carboniferous-Early Triassic times. By the end of Triassic, erosional products covered the entire eastern Barents shelf (Petrov et al., 2008). The Pai-Khoi-Novaya Zemlya fold belt is a segment of the orogenesis (Korago et al., 2004) and was deformed during thrusting of Novaya Zemlya (Fig. 1.1a, Fig. 2.1) above the Barents plate in the Triassic times (Petrov et al., 2008). The basement of the Novaya Zemlya Archipelago is dominated by Meso- to Neoproterozoic metasedimentary rocks and associated igneous intrusions. The igneous intrusions are of mafic to felsic composition and the oldest granite is dated as 1300 Ma (Korago et al., 2004).

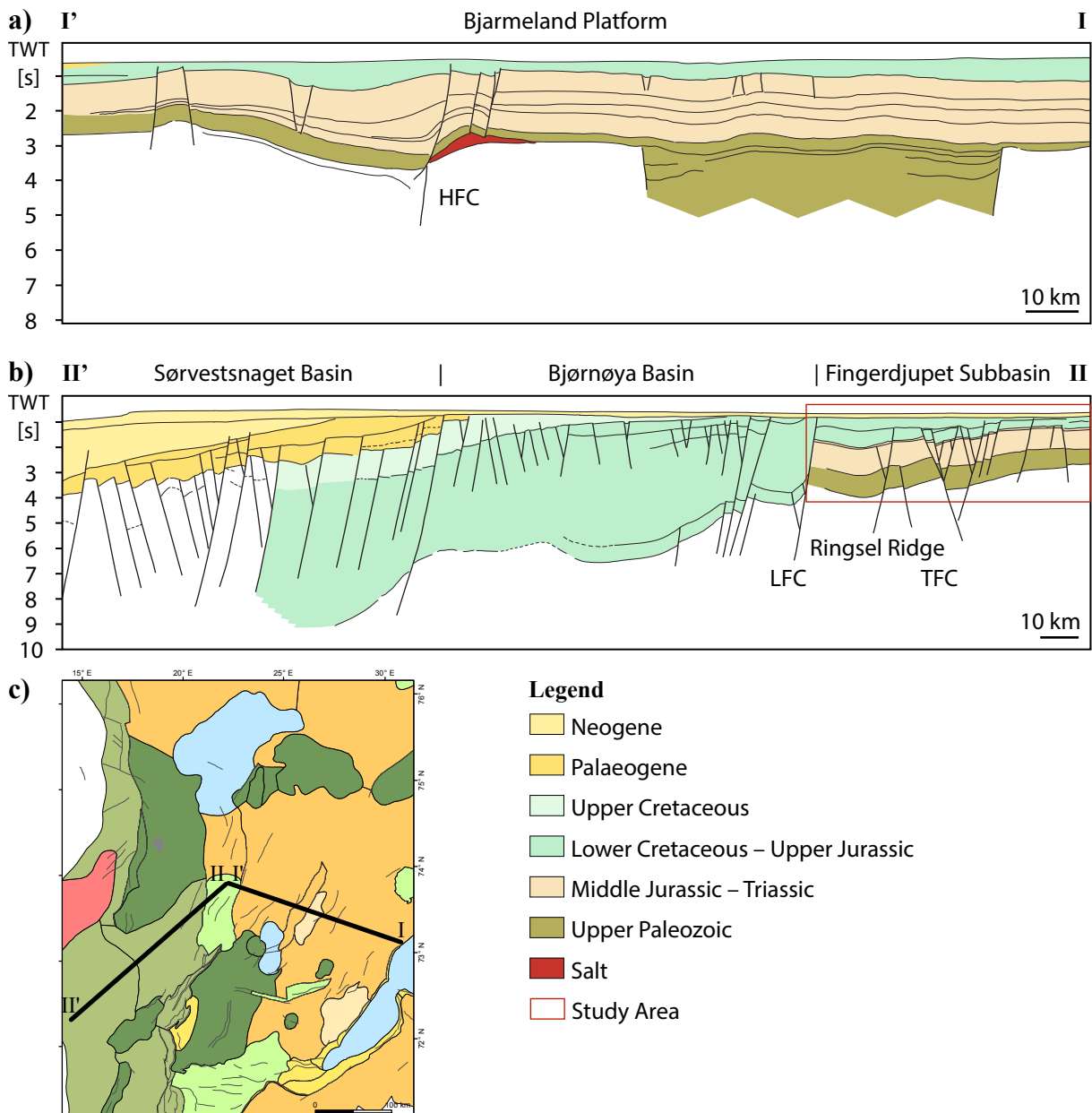
## **2.2 Tectonic Framework**

North-northwest-oriented structural trends of the Scandinavian Caledonides are inherited in the basement of the Barents shelf (e.g. Gernigon and Brønner, 2012). These lineaments are interpreted to have played a significant role in the generation of the present-day rift systems in the Barents shelf (Fig. 1.1b, Doré et al., 1991; Gudlaugsson et al., 1998; Ritzmann and Faleide., 2009; Faleide et al., 2010; Gernigon and Brønner, 2012). The post-Caledonian evolution of the shelf can be divided into three main extensional phases: late Paleozoic, Late Jurassic to Early Cretaceous and early Cenozoic. In addition to two uplift phases during Early Cretaceous and Palaeogene to recent (Faleide et al., 1993; Dore, 1995; Anell et al., 2014; Faleide et al., 2015; Riis & Fjeldskaar 1992; Riis 1996; Henriksen et al., 2011).

The tectonic setting during the Early Cretaceous was influenced by the opening of the Amerasia Basin (Grantz et al., 2011) and the North Atlantic rift system (Faleide et al., 1993). The breakup of the Amerasia Basin and the associated High Arctic Large Igneous Province (HALIP, Petrov et al., 2016) is interpreted to cause uplift of the northern margin of the Barents shelf (Gjelberg & Steel, 1995; 2012; Worsley, 2008; Midtkandal & Nystuen, 2009). This is supported by southeast directed fluvio-deltaic and clinofold systems on Svalbard and in the south-western Barents shelf, respectively (e.g. Midtkandal & Nystuen, 2009; Midtkandal, 2019). Intrusive and extrusive igneous rocks on Svalbard and Franz Josef Land suggest that the uplift peaked around latest Barremian (Maher, 2001; Corfu et al., 2013; Polteau, 2015).

The North Atlantic rift system, between Greenland and Baltica (Fig. 2.1), continued into the present day south-western Barents shelf during the Late Jurassic to Early Cretaceous. The earliest Cretaceous (Berresian–Hauterivian) structuring is characterized by north-northeast trending normal faults. Large throws and deep depocenters defined by the Harstad, Tromsø and Bjørnøya basins at the present day western margin (Fig. 2.2b, Faleide et al., 1993). Strike-slip movements accompanying the normal displacement are also recognized, especially in the eastern margin of the Bjørnøya Basin (Bjørnøyrenna Fault Complex). Further east, the north to northeast trending faults developed significant subsidence in the eastern flank of the Bjørnøya Basin and established the Fingerdjupet Subbasin (Fig. 1.1b, Fig. 2.2, Serck et al., 2017). Minor fault activity is also recognized in the Hoop Fault Complex on the Bjarmeland Platform (Fig. 2.2a, Faleide et al., 2019). The fault activity is suggested to have terminated in the northernmost part of the Barents shelf (Faleide et al., 1993), where compressional tectonic activity occurred (Kairanov et al., 2018). North oriented fault activity on Svalbard is suggested to be expressed as collapse structures within delta fronts on the eastern Spitsbergen (Onderdonk & Midtkandal, 2010).

## Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development



**Fig. 2.2 Regional Seismic Line** Regional seismic line showing the structural elements of the Fingerdjupet Subbasin in a regional context (figure modified from Faleide et al., 2015).

Tectonic quiescence is inferred during the Barremian in the Bjørnøya Basin (Blaich et al., 2017), Fingerdjupet Subbasin (Fig. 1.1b, Fig. 2.2b, Serck et al., 2017) and in the Hoop area on the Bjarmeland Platform (Fig. 2.2a, Faleide et al., 2019). Crustal stretching and thinning in Tromsø and Bjørnøya basins from the earliest Cretaceous fault activity is inferred to have caused uplift and sub-aerial exposure of the Loppa High during early Barremian (Indrevær et al., 2016). The uplift has been suggested to have occurred through different stages from Late Jurassic/earliest Cretaceous times (Gabrielsen et al., 1990; Glørstad-clark, 2011) and to have reached its peak during the Barremian (Indrevær et al., 2016). The differential uplift was accompanied by inversion tectonics in the faults surrounding the Loppa High and in the adjacent Hammerfest Basin (Fig. 1.1b, Indrevær et al., 2016). Indrevær et al. (2017) suggests that the uplift model

for the Loppa High (phase change driven vertical movements) also could explain the uplift of the Stappen High, but the timing of the uplift phase developing the present-day Stappen High is poorly constrained (Anell et al., 2016). The present structural configuration of the Stappen High is suggested to have developed during Cretaceous–Cenozoic (Anell et al., 2016) and early Cenozoic (Faleide, 1993; Worsley et al., 2001, Blaich et al., 2017).

A second extensional event occurred during Aptian, caused renewed fault activity along north-northeast trending faults. Previous established depocenters such as the Tromsø Basin, Bjørnøya Basin and Fingerdjupet Subbasin experienced subsidence (Fig. 1.1b, Fig. 2.2b, Faleide et al., 1993; Blaich et al., 2017; Serck et al., 2017), as well as the Hoop area on Bjarmeland Platform (Fig. 2.2a, Faleide et al., 2019). The Albian succession is characterized by post-rift subsidence and rapid infill of previous established paleotopography with increasing magnitude towards the western basins (Faleide et al., 1993).

## 2.3 Stratigraphy

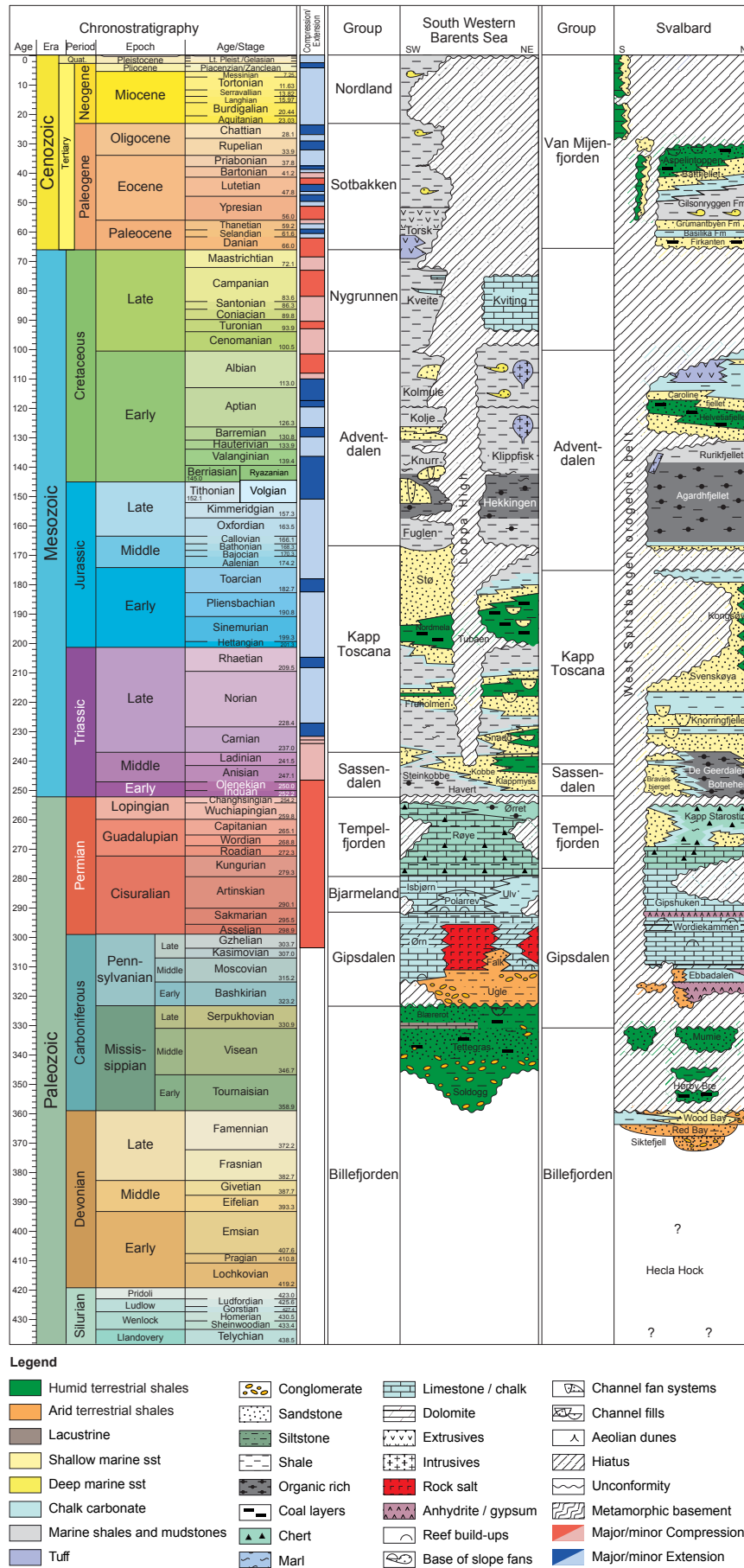
### Svalbard

The Lower Cretaceous succession in Svalbard consists of Rurikfjellet (Valanginian–Hauterivian/early Barremian), Helvetiafjellet (Barremian–early Aptian) and Carolinefjellet formations (early Aptian–Albian) (Mørk et al., 1999, Fig. 2.3). The base of the Lower Cretaceous succession is a regional marker (Dypvik et al., 2017) interpreted to be age equivalent to the Base Cretaceous Unconformity (BCU) offshore in the Barents shelf (Fig. 2.3, Grundvåg et al., 2017). The Rurikfjellet Formation (Fig. 2.3) is dominated by open marine shales that pass vertically into shallow marine, delta-front sandstones in the southeast Svalbard. The coarsening upwards trend are interpreted as a regression induced by the uplift and southern tilt of Svalbard (Gjelberg and Steel, 1995). The source of coarse-grained sediments is linked to the west-northwest hinterlands based on the southeast directed paleocurrents and thinning of the unit (Grundvåg et al., 2017). The lower boundary of Helvetiafjellet Formation (Fig. 2.3) is a regional unconformity of Barremian age (Fig. 2.3, ranging from 129–117Ma, Midtkandal et al., 2016; Vicker et al., 2016). The Barremian unconformity has been inferred to represent a subaerial unconformity associated with deep incisions in the south-west Svalbard (Nathorst Land) as a result of increased slope gradient caused by the emerging northern margin of Svalbard (Midtkandal and Nystuen, 2009). The Helvetiafjellet Formation (Fig. 2.3) is subdivided into an extensive sand sheet of the Festningen member followed by the heterolithic Glitrefjellet Member (Midtkandal et al., 2008). The Festningen Member (Fig. 2.3) consists of coarse grain fluvial braided-plain deposits interpreted as a forced regressive unit extending across the entire southern



Spitsbergen (Gjelberg & Steel, 1995). At the southern edge of Spitsbergen, these sandstone packages are interpreted to either pass laterally into marine mudstone (Gjelberg & Steel 1995, 2012) or continue into the Barents shelf (Midtkandal et al. 2009). Southeast paleocurrents indicate that a north-western source region also existed for the Festingen Member (Grundvåg et al., 2017; Midtkanal et al., 2019a). Glitrefjellet Member consists of fluvio-deltaic deposits with alternating mudstones, sandstones and thin coal beds reflecting a tide dominated coast (Gjelberg and Steel, 1995; Midtkandal and Nystuen, 2009). The base of the Carolinefjellet Formation (Fig. 2.3) forms an regional flooding event marking a return from a coast to an open marine shelf environment (Grundvåg et al, 2017). This unit consists of storm-influenced and storm-dominated sand sheets alternating with outer shelf mudstone(e.g. Hurum et al., 2016).

# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development



**Fig. 2.3 Lithostratigraphic chart** Lithostratigraphy of the southwestern Barents shelf and Svalbard (figure modified from Gradstein et al., 2010)

### Barents shelf

The Lower Cretaceous succession in the south-western Barents shelf is divided into the time-equivalent Knurr and Klippfisk formations (Berriasian to early Barremian age), overlain by Kolje Formation (early Barremian to late Barremian/early Aptian age), followed by Kolmule Formation (Aptian to mid-Cenomanian, Fig. 2.3, Dalland et al., 1988; Mørk et al., 1999). The base of the Lower Cretaceous succession is referred to as the Base Cretaceous Unconformity (BCU), representing a regional unconformity and a correlative conformity (e.g. Mørk et al., 1999). The BCU separates the deep marine deposits of the Upper Jurassic Hekkingen formation and shallow to open marine deposits of the Klippfisk and Knurr formations (Fig. 2.3, Arhus et al., 1990; Faleide et al., 1993; Mørk et al., 1999). The shallow marine Klippfisk formation consists of limestone and marl interpreted as a transgressive condensed deposits on the western platforms and highs, such as on the Ringsel Ridge and on the Bjarmeland Platform (Fig. 1.1b, Fig. 2.2a, Smelror et al., 1998; Århus and Kelly, 1990). The Klippfisk Formation passes laterally into the Knurr Formation (Fig. 2.3) which is composed of open-marine claystone interbedded with thin limestone and dolomite beds (Mørk et al., 1999). Sandstones and conglomerates are also present within the Knurr Formation in shallow marine wedges (well 7120/1-2 and 7122/2-1), deep marine wedges and fans (well 7120/10-1 and 7120/10-2) in the Hammerfest Basin (Fig. 1.1b, e.g. Seldal, 2005; Marin et al., 2018). The coarse grained deposits are interpreted to be erosional products of older rocks on the adjacent Loppa High and the Troms–Finnmark Platform (Fig. 1.1b, Seldal, 2005; Marin et al., 2018). The Klippfisk and Knurr formations are age correlated to the Rurikfjellet Formation in Svalbard (Fig. 2.3, Grundvåg et al., 2017; Midtkandal et al., 2019). The base of the Kolje Formation (Fig. 2.3) has been interpreted as a regional unconformity and correlative conformity (Smelror et al., 1998) referred to as Lower Cretaceous Unconformity (LCU) (Midtkandal et al., 2019). The LCU is inferred to represent a regressive surface of marine erosion (Midtkandal et al., 2019). The Kolje Formation is dominated by shales and claystones with minor interbeds of limestone and dolomite interpreted as open marine deposits. Sandstones and siltstones also occur in the upper part of the formation (Dalland et al., 1998). The Kolje Formation is age correlated to the Helvetiafjellet Formation on Svalbard (Fig. 2.3, Grundvåg et al., 2017; Midtkandal et al., 2019a). Kolmule Formation (Fig. 2.3) is also considered to be open marine deposits of claystone and shale, with minor interbeds of thin siltstones in addition to limestone and dolomite stringers (Dalland et al., 1998). Sandstones (well 7120/6-3S, 7120/2-3S and 7220/10-1) and

conglomerates (well 7120/6-3S) are also encountered within this unit (NPD, 2020). The Kolmule Formation is considered to be time equivalent to the Carolinefjellet Formation on Svalbard (Fig. 2.3, Grundvåg et al., 2017; Midtkanal et al., 2019).

### Lower Cretaceous Clinofolds

Clinofolds within the Lower Cretaceous succession in the south-western Barents shelf are characterized as two large scale progradation systems; a system with an dominant southeast direction and a system with south-west direction.

The southeast-directed system in the south-western Barents shelf extends from the Bjørnøya Basin (Midtkandal et al., 2019), across Fingerdjupet Subbasin and onto the Bjarmeland platform (Grundvåg et al., 2017; Hinna, 2017; Marin et al., 2017; Bryn et al., 2019; Faleide et al., 2019; Midtkandal et al., 2019). The clinofolds on the eastern Bjarmeland Platform are interpreted to be older than on the western part of the platform and in the Fingerdjupet Subbasin based on onlap relations (Faleide, 2017; Faleide et al., 2019). The southeast-directed system has been interpreted to be of Hauterivian–early Barremian (Grundvåg et al., 2017), Barremian–Aptian (Marin et al., 2017) and Barremian age (Bryn et al., 2019; Faleide et al., 2019; Midtkandal et al., 2019). Several sources located to the west-northwest of the Barents shelf have been suggested for the southeast-directed system, including northeastern Greenland, the Lomonosov high, Chukchi Borderland and Crockerland (Grundvåg et al., 2017; Midtkandal et al., 2019).

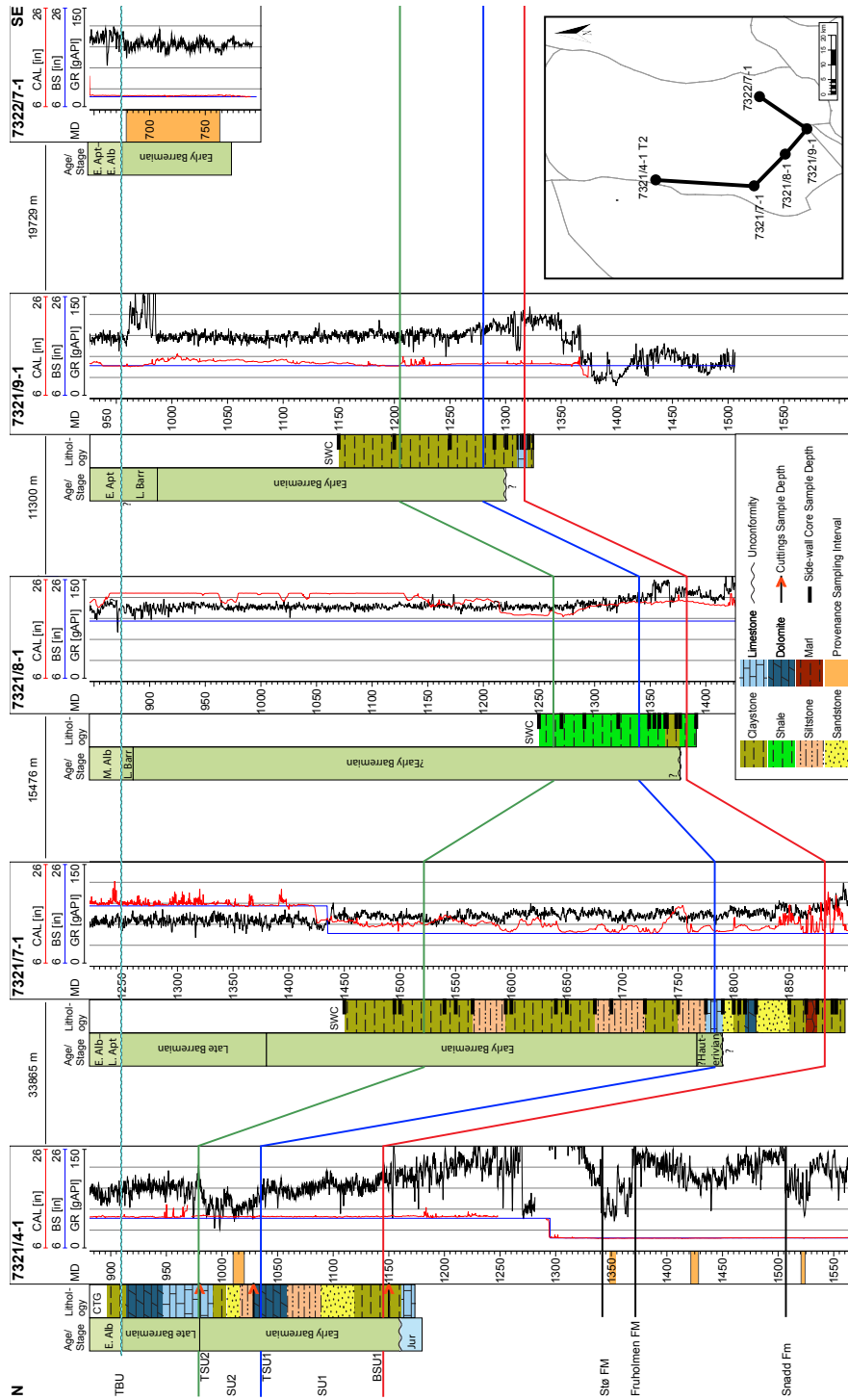
The southwest-directed system advanced across almost the entire Bjarmeland Platform and Nordkapp Basin in the eastern part of the Barents shelf during Hauterivian–Albian (Grundvåg et al., 2017), Valanginian to middle Cenomanian (Marin et al., 2017), or Aptian–Cenomanian time (Midtkandal et al., 2019) age. The source regions are suggested to be in the east-northeast (Marin et al., 2017; Midtkandal et al., 2019), such as the Taimyr and North Kara region in north Russia (Midtkandal et al., 2019).

## 3 DATA AND METHODOLOGY

### 3.1 Data

#### Well and seismic data

The wells used in this study include five exploration wells (7321/4-1, 7321/7-1, 7321/8-1, 7321/9-1 and 7322/7-1, Fig. 3.1 and ) located in the southern and in the north-western part of the study area (Fig. 3.2). All of the wells have check shot survey, petrophysical well logs and cuttings descriptions (Table 3.1 ). Well 7321/4-1 and 7322/7-1 were used for the provenance analysis (detrital zircon U/Pb geochronology) and all of the wells were included for biostratigraphic age dating (Table 3.1 ). The caliper log (hole diameter) relative to the bit size shows large size difference for wells 7321/7-1, 7321/8-1 and intervals in well 7321/9-1 (Fig. 3.1). Irregular boreholes and large spacing between the tool and borehole wall may cause inaccurate well log readings. Therefore, the readings from 7321/7-1, 7321/8-1, and parts of 7321/9-1, are considered to be less reliable than those from wells 7321/4-1 and 7322/7-1(Fig. 3.1).

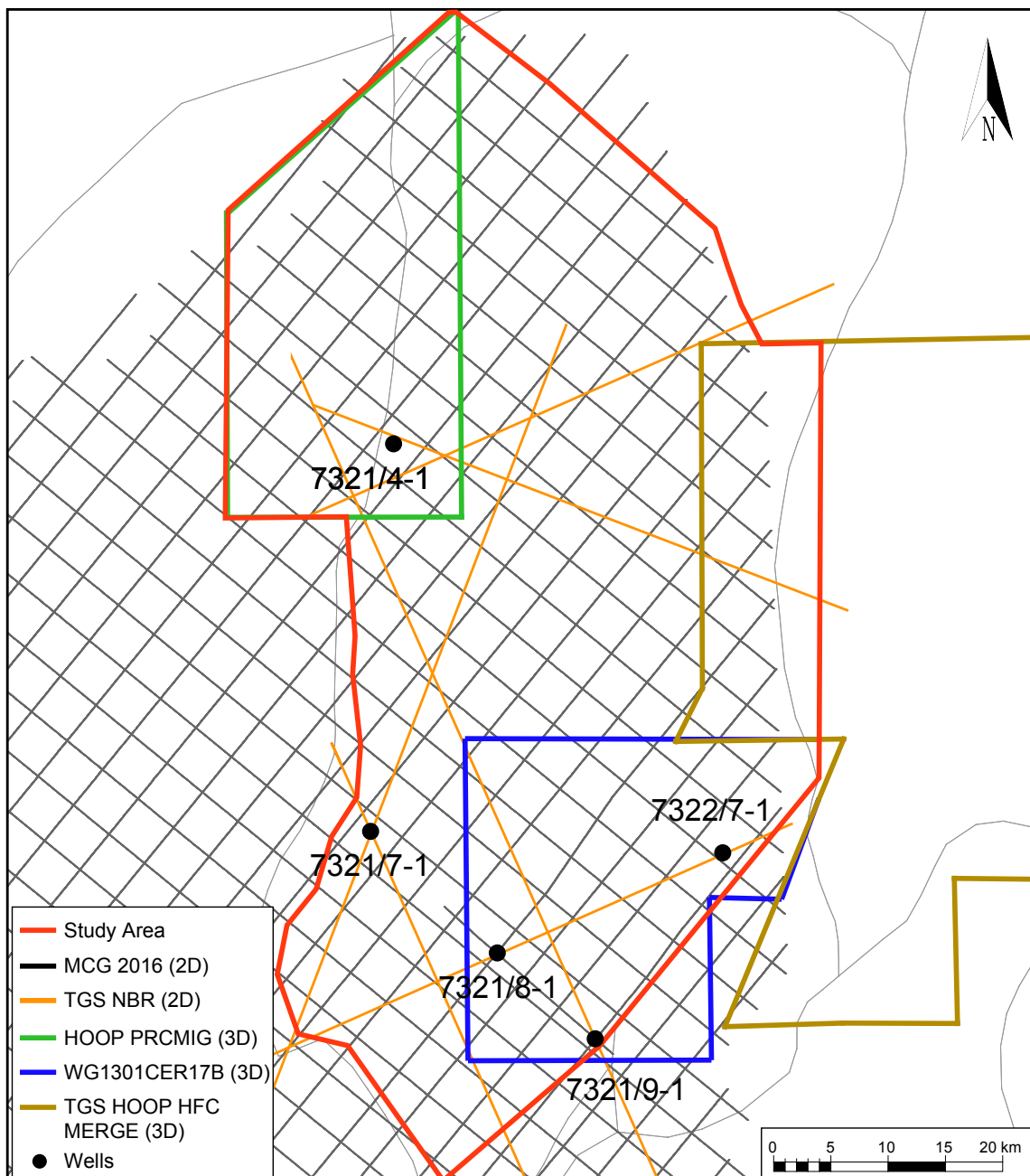


**Fig. 3.1 Well Correlation** North-Southeast Well Correlation. MD = Measured Depth, CAL = Caliper, BS = Bit Size, GR = Gamma Ray, TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1 = Base Seismic Unit 1, FM = Formation, E. Alb = Early Albian, Jur = Jurassic, L. Apt = Late Aptian, M. Alb = Middle Albian, L. Barr = Late Barremian, E. Apt = Early Aptian, CTG = Cuttings, SWC = Side-wall Core. The well logs are flattened on TBU and include Age/Stage columns showing the result of the palynological analyses. Lithology columns are included for well 7321/4-1 (based on cutting descriptions), 7321/7-1, 7321/8-1, 7321/8-1 (based on side-wall core descriptions). The MD column for well 7321/4-1 and 7322/7-1 display intervals used for provenance sampling. Inset map in the lower right shows the location of the wells. The seismic units (SU1 and SU2) are penetrated by four of the wells and not well 7322/7-1.

## Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

**Table 3.1 Well Data** PL= Petrophysical Logs (caliper, gamma ray, neutron porosity, sonic velocity and resistivity), L = Lithostratigraphic Information, CTG = Cuttings, SWC = Side-wall Core, C = Conventional Core, DML = Dip-meter Log, IL = Image Log, B = Biostratigraphic Information, P = Provenance Data, CS = Check-shot

Well	Year of completion	Oldest penetrated age	Type									
			PL	L	CTG	SWC	C	DML	IL	B	P	CS
7321/8-1	1987	Late Permian	✓	✓	✓	✓		✓		✓		✓
7321/7-1	1988	Middle Triassic	✓	✓	✓	✓	✓	✓		✓		✓
7321/9-1	1988	Late Triassic	✓	✓	✓	✓		✓		✓		✓
7322/7-1	2018	Early Cretaceous	✓	✓	✓	✓			✓	✓	✓	✓
7321/4-1	2018	Late Triassic	✓	✓	✓					✓	✓	✓



**Fig. 3.2 Wells and seismic data** Map of well data and seismic data within the study area.

The seismic data covering the study area consist of 2D and 3D reflection seismic data (, , Fig. 3.2). The 2D dataset includes four TGS NBR 2D lines and the MCG2016 survey which consists of a dense grid of 2D lines, with around 4 km spacing, covering the study area. The 3D dataset covers the north-western corner (HOOP PRCMIG) and south-eastern part of the study area (WG1301CER17B and TGS Hoop HFC Merge). Average seismic velocities in the Lower Cretaceous, between horizon TBU and BSU1, Fig. 3.1, interval ranges from 2761 m/s to 3362 m/s (). The measured dominant frequency in the same interval and calculated wavelength indicates that the vertical resolution range from 23-35m ().

**Table 3.2 Seismic Dataset** *The five different seismic surveys used in this study.*

Name	Type	Year	Courtesy
MCG2016	2D	2016	TGS/Searcher/Spectrum/PGS/ VBPR
TGS NBR	2D	2008, 2009 and 2011	TGS
HOOP PRCMIG	3D	2017	TGS
WG1301CER17B	3D	2017	Spirit Energy
TGS Hoop HFC Merge	3D	2016	TGS/Searcher/Spectrum/PGS/VBPR

**Table 3.3 Average Seismic Velocity** *Table with calculated average velocities at four of the wells. MD = Measured Depth, TWT = Two-way-time, D = Difference (e.g. MD top - MD base).*

Well	MD top [m]	MD base [m]	TWT top [s]	TWT base [s]	D MD [m]	D TWT [s]	Average Velocity (v=2xDMD/DTWT) [m/s]
7321/4-1	910	1145	0,962	1,105	235	0,143	3287
7321/7-1	1250	1882	1,232	1,608	632	0,376	3362
7321/8-1	875	1383	0,985	1,353	508	0,368	2761
7321/9-1	955	1317	1,039	1,299	362	0,260	2785

**Table 3.4 Vertical Resolution** *Table with approximate dominant frequency measured in the seismic, calculated wavelength and vertical resolution.*

Seismic Survey	Well	Approximate Dominant Frequency (fd) [Hz]	Wavelength (wl) (v/fd) [m]	Vertical Resolution, (wl/4) [m]
HOOP PRCMIG	7321/4-1	25-35	131-94	33-23
TGS NBR	7321/7-1	25-35	134-96	33-24
MCG1401	7321/8-1	20-30	138-92	34-23
WG1301CER17B	7321/9-1	20-30	139-93	35-23

## 3.2 Methodology

### 3.2.1 Biostratigraphy and detrital zircon U/Pb geochronology

Palynological analyses and detrital zircon U/Pb geochronology was carried out by the Geological Survey of Denmark and Greenland (GEUS) in Denmark. Palynological dating (for details, see Nøhr-Hansen et al., 2019) was carried out in order to establish age control and correlate key events (e.g. unconformity)



## Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

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through the five wells in the Fingerdjupet Subbasin (Fig. 3.1). For the palynological analysis, 30 ditch cutting samples were studied from well 7321/4-1, two in the Upper Jurassic and 28 in the Lower Cretaceous. The Lower Cretaceous interval was sampled from 29 ditch cutting samples in well 7321/7-1, 23 ditch cutting samples in 7321/8-1 and 38 ditch cutting samples in well 7321/9-1. Three sidewall cores and 34 ditch cutting samples were studied from well 7322/7-1.

The samples for the detrital zircon U-Pb geochronology were made from average samples in well 7321/4-1 and in well 7322/7-1 ( ). Four average samples from cuttings were made from the following four intervals in well 7321/4-1; (1) 1521 – 1524 m MD (Snadd Formation), (2) 1422 – 1428 m MD (Fruholmen Formation), (3) 1348 – 1354 m MD (Stø Formation) and (4) 1010 – 1020 m MD (SU2, Fig. 3.1). For well 7322/7-1, one average sample were made from cuttings and sidewall cores in the interval between 679 to 763 MD (Kolje FM, Fig. 3.1).

**Table 3.5 Zircon U/Pb Samples**

Well	Sample Interval [MD]	Unit
7322/7-1	679-763	Kolje FM (Lower Cretaceous)
7321/4-1	1010-1020	SU2 (Lower Cretaceous)
7321/4-1	1348-1354	Stø FM (Lower – Middle Jurassic)
7321/4-1	1422-1428	Fruholmen FM (Upper Triassic)
7321/4-1	1521-1524	Snadd FM (Middle – Upper Triassic)

The analysis of zircon for U/Pb isotopic dating were performed using laser ablation inductively coupled plasma mass spectrometry (LA-ICPMS). The equipment used included a NWR 213 laser ablation instrument from Elemental Scientific lasers (ESL) that was coupled to an Element2 magnetic sector-field ICPMS from Thermo-Fisher Scientific. The analyses was carried out on zircon grains mounted in epoxy pucks, polished, and imaged by SEM using either cathodeluminescence (CL) or back-scattered elecetrons (BSE) prior to the LA-ICPMS analyses.

For quality control of the geochronology analyses, the secondary standards, Plešovice and Harvard 91500 was measured during the zircon analyses, yielding an average accuracy and precision ( $2\sigma$ ) on the dates within 3% deviation. Zircon U-Pb ages ( $^{206}\text{Pb}/^{238}\text{U}$ ,  $^{206}\text{Pb}/^{238}\text{U}$  and  $^{207}\text{Pb}/^{206}\text{Pb}$ ) presented in this study have  $2\sigma$  error and are constrained by a concordance-discordance criterion of 10% (see appendix).

The ISOPLOT program (Ludwig, 2008) was used to present the zircon U-Pb ages in histograms and probability density plots. These plots includes concordant  $^{206}\text{Pb}/^{238}\text{U} < 1000 \text{ Ma}$  and  $^{207}\text{Pb}/^{206}\text{Pb} > 1000 \text{ Ma}$ .

### **3.2.2 Well logs and Seismic interpretation**

The seismic stratigraphic framework was interpreted using petrophysical logs, palynological analyses and seismic interpretation. Two seismic units (SU1 and SU2) bound by three key seismic surfaces (BSU1, TSU1 and TSU2, Fig. 3.1) were defined mainly based on seismic downlap terminations and seismically guided well correlations. This is due to poor identification of maximum flooding surfaces on the gamma ray logs and poor correlation of the NPD lithostratigraphic units. The interval of interest (between BSU1 and TSU2) is locally eroded by an unconformity in the top of the Barremian aged interval (TBU, Fig. 3.1) and by an unconformity in the top of Lower Cretaceous succession which is interpreted to represent Base Quaternary Unconformity (BQU). These reflectors were mapped but will not be described in any further detail.

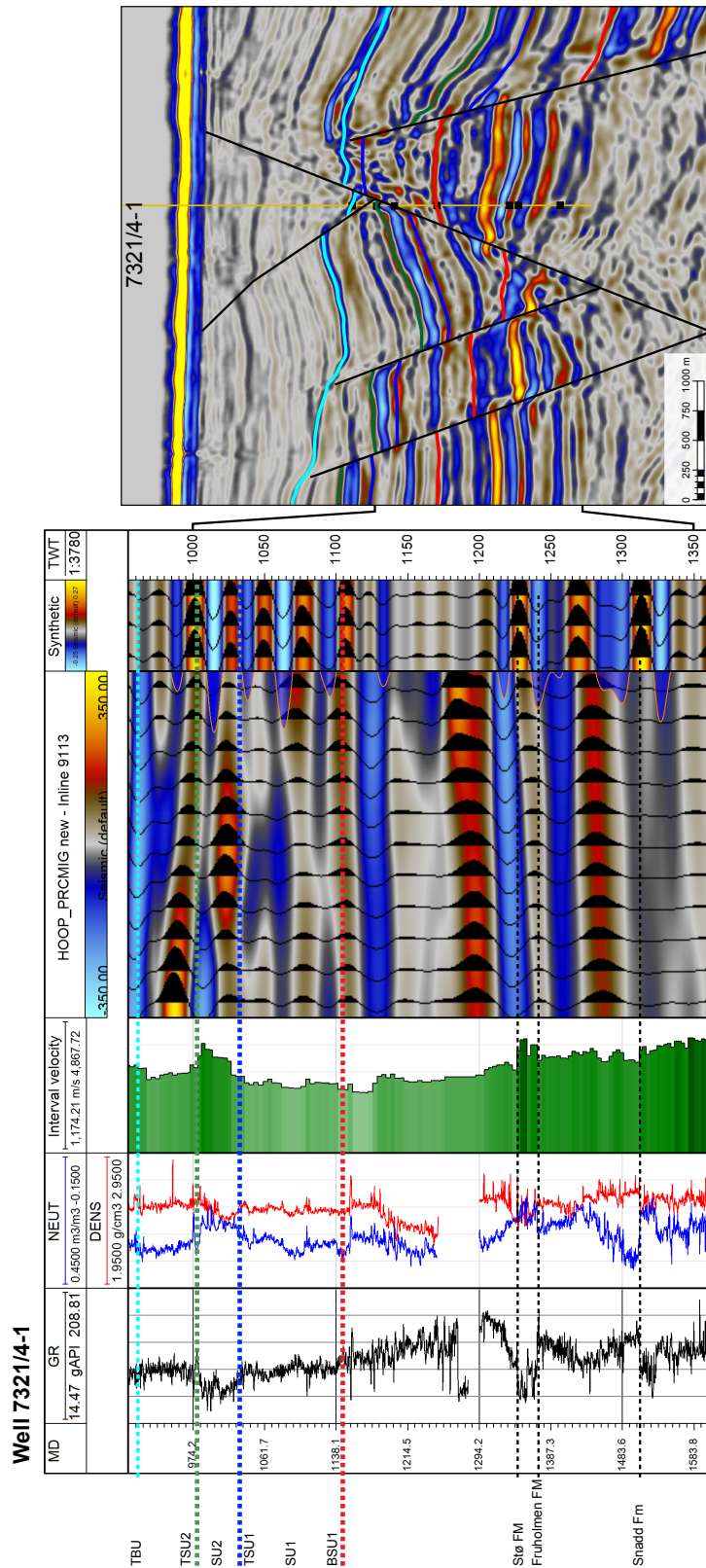
The seismic units were correlated in the wells and the faults affecting the units were mapped throughout the seismic dataset. Fault interpretations and time thickness maps generated from the seismic interpretation were used to describe the main structures and to interpret the infill of the basin. The seismic character of the seismic units is described based on main reflector characteristics including continuity, amplitude and geometry. Seismic sections parallel to the clinoform progradation direction are flattened along the clinoform downlap surface (BSU1), considered to be closest to paleo-horizontal datum in the studied interval, to describe the clinoform geometry and the trajectory of the rollover point (shelf-edge).

Identification of lithological trends and interpretation of the sandstone distribution were based on gamma ray logs, cuttings from well 7321/4-1 and sidewall cores from well 7321/7-1, 7321/8-1 and 7321/9-1.

### **3.2.3 Seismic-Well Tie**

The integrated seismic well-tie suite in the Petrel Schlumberger software were used to perform the seismic-well tie. Checkshot data from the wells were used to establish a time-depth relationship, by calibrating (correct the sonic log to seismic times) the sonic log. In order to produce a synthetic seismogram, the density log and the sonic log were used to compute acoustic impedance and the reflection coefficient for normal incidence. Furthermore, the reflection coefficient was convolved with a zero phase wavelet to generate synthetic traces to match with the seismic traces. The Ricker wavelet were used to compare synthetic traces with the real seismic traces. The frequency of the Ricker wavelet was adjusted to obtain

the best possible match with the frequency spectrum of the seismic traces. The BSU1 and TBU were used to tie the seismic data and the well data in the study area. The BSU1 is modelled as a positive reflector with high amplitude in the synthetic seismogram and correlates good with the seismic data (Fig. 3.3). A negative amplitude reflector were picked for the TBU which shows medium to high amplitude in the synthetic and high amplitude in the seismic, giving a moderate good correlation (Fig. 3.3).

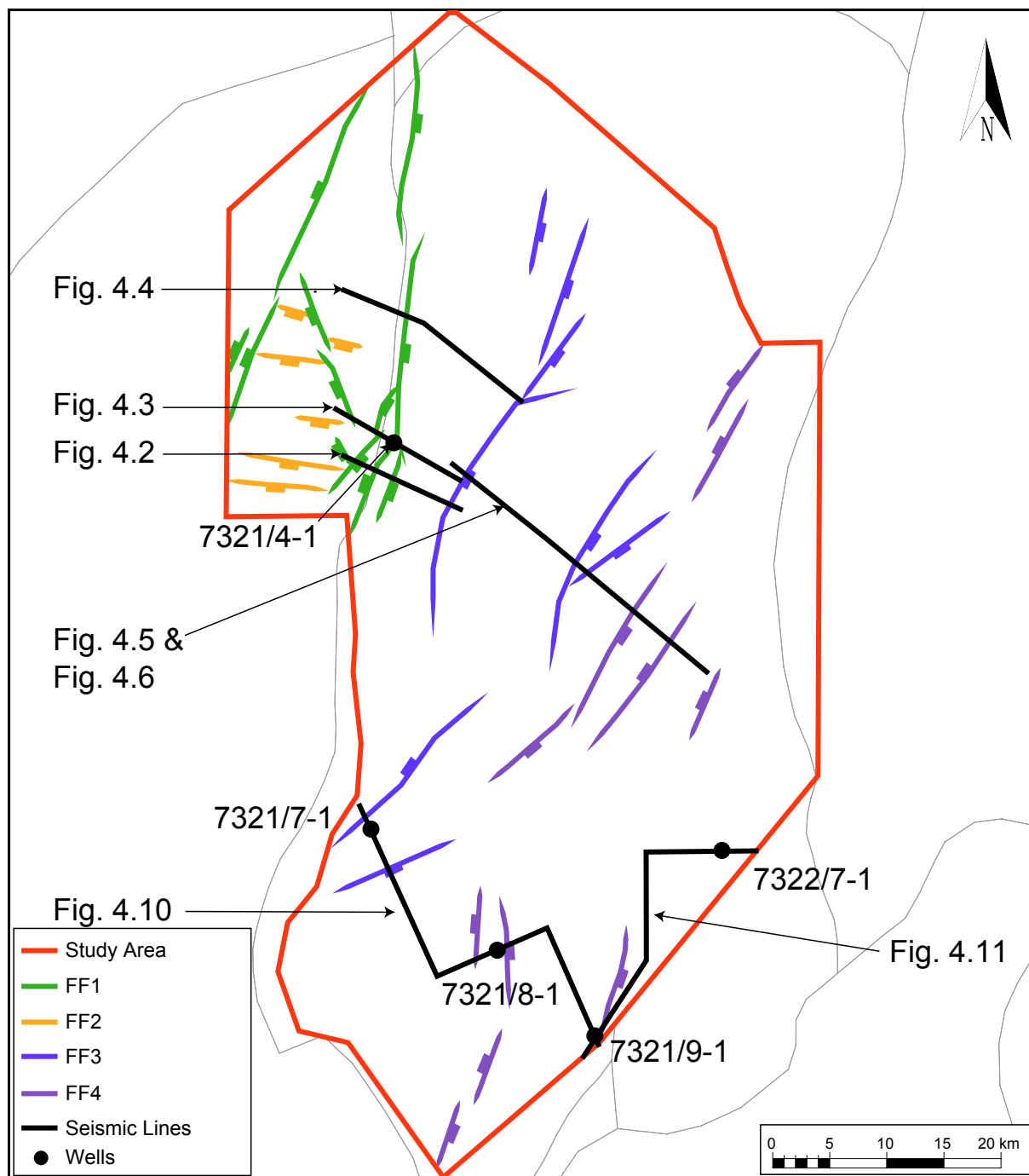


**Fig. 3.3 Synthetic Seismogram** Well logs, seismic data and synthetic seismicogram are shown to the left. An NW-SE arbitrary seismic line across well 7321/4-1 is shown to the right. MD = Measured Depth, GR = Gamma-ray log, NEUT = Neutron log, DENS = Density log, TWT = two-way-time, Unc. = Unconformity. The seismic data display an N-S inline across the well. A 25-30-35 Hz zero phase Ricker wavelet was used to compute the synthetic seismicogram. A increase in acoustic impedance yields positive amplitude displayed as a red reflector. The picked horizons for the interval of interest is defined two positive reflectors named Base Seismic Unit 1 and Top Seismic Unit 2.

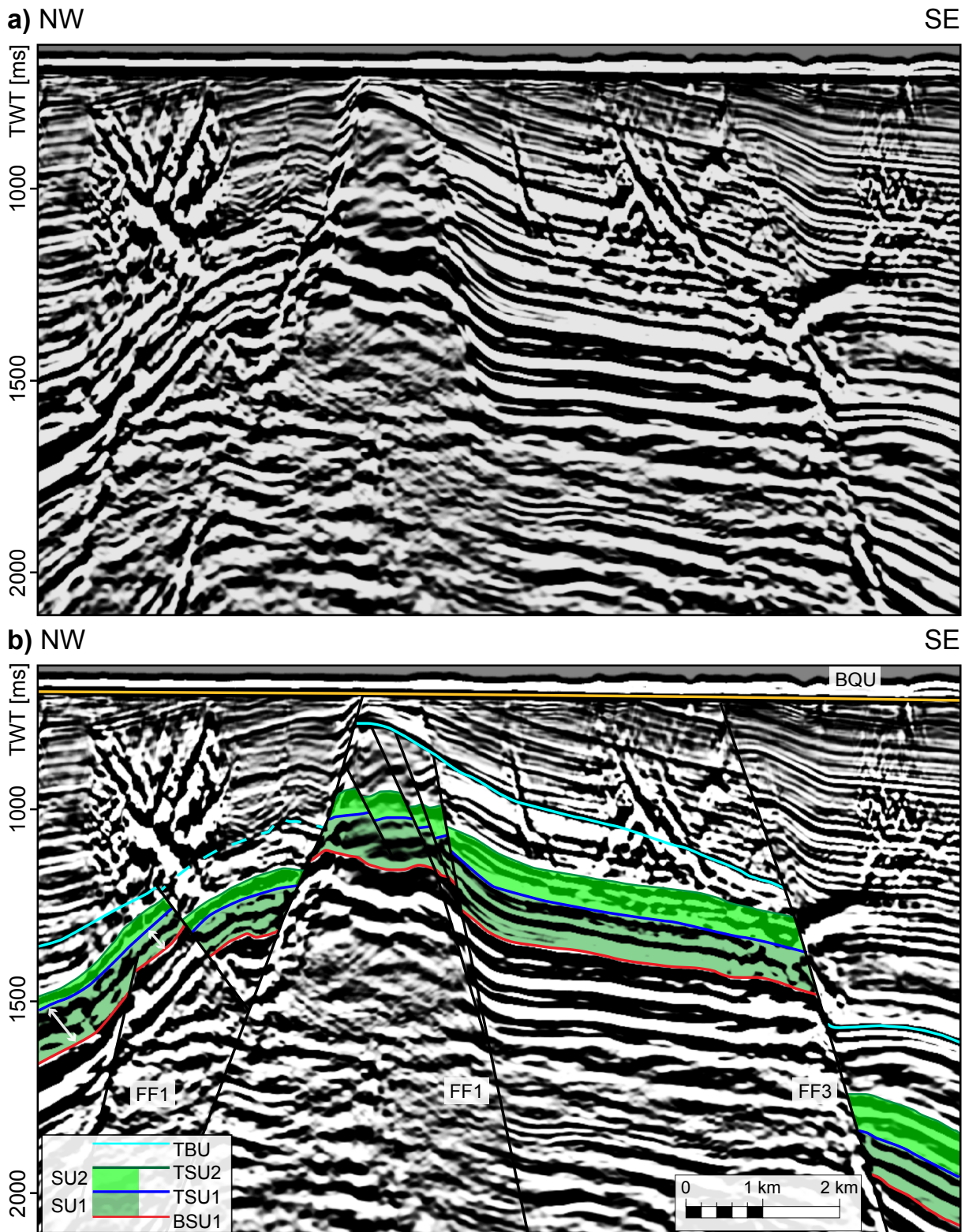
## 4 RESULTS

### 4.1 Fault Families

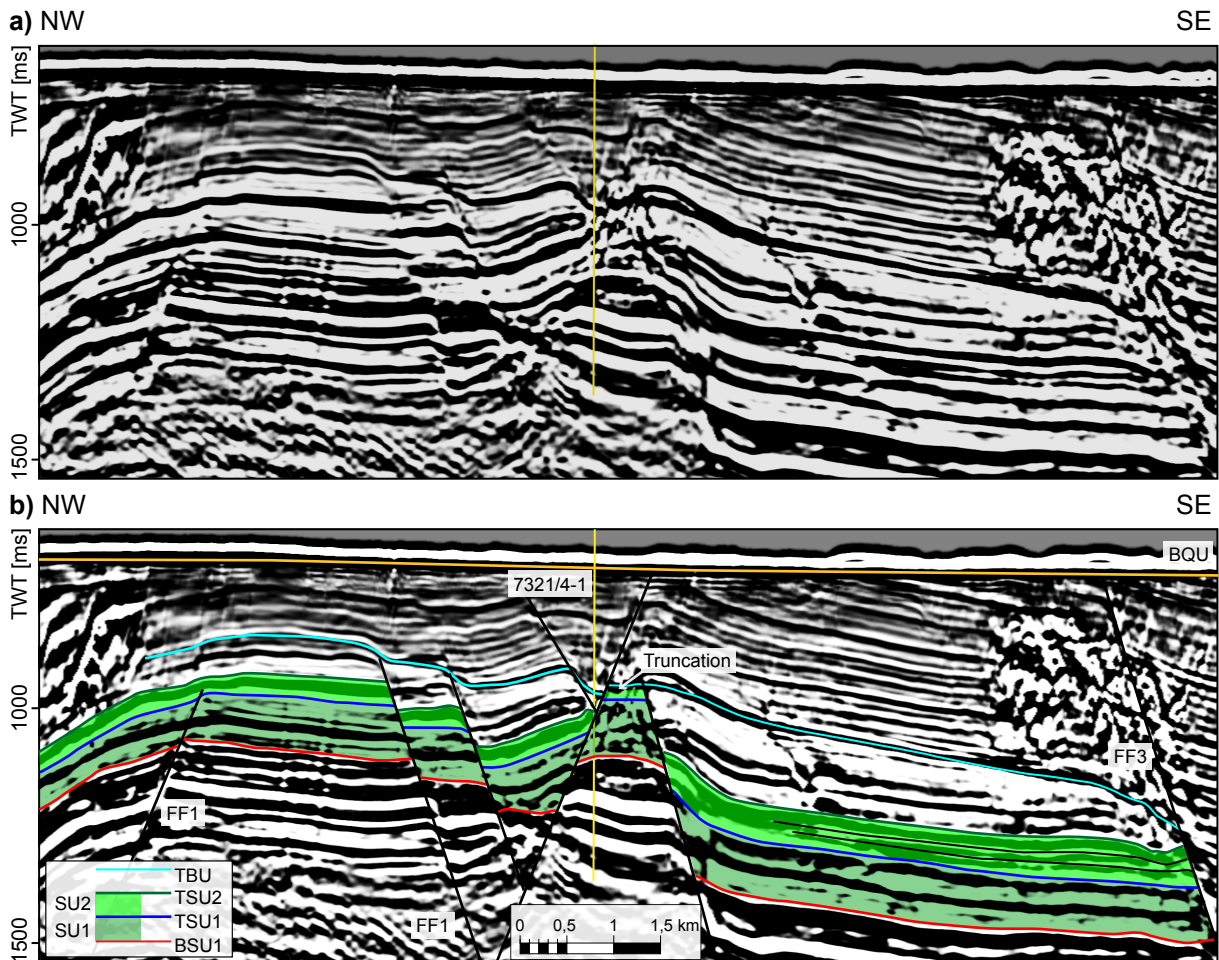
The main faults affecting the seismic units in the study area are divided into four fault families (FF1-FF4) of similar trend and age (Fig. 4.1). Fault Family 1 (FF1) offset both of the seismic units and constitute normal fault striking NNW-SSE, N-S and NNE-SSW located in the north-western part of the study area (Fig. 4.1, Fig. 4.2, Fig. 4.3, Fig. 4.4). Well 7321/4-1 penetrates the fault plane of a NNE-SSW striking fault segment of FF1 at 1035 m MD, in this location the TSU1 is displaced out of the wellbore and the well penetrated the middle part SU1 (Fig. 4.3). Fault Family 2 (FF2) consist of south-dipping normal faults striking E-W between the individual segments of FF1 (Fig. 4.1). Most these fault offset both of the seismic unit in the north-western part of the study area. Fault family 3 (FF3) offset both of the seismic units and constitute normal faults striking N-S to NE-SW along the central part of the study area (Fig. 4.5 and Fig. 4.6). Fault Family 4 (FF4) is located in the south and northeast and consists of normal fault striking N-S to NE-SW and offset both of the seismic units (Fig. 4.1, Fig. 4.6).



**Fig. 4.1 Fault Families and Presented Seismic Lines** Map of the study area (outlined in red) including four fault families defined in this study, exploration wells and location of presented seismic lines.

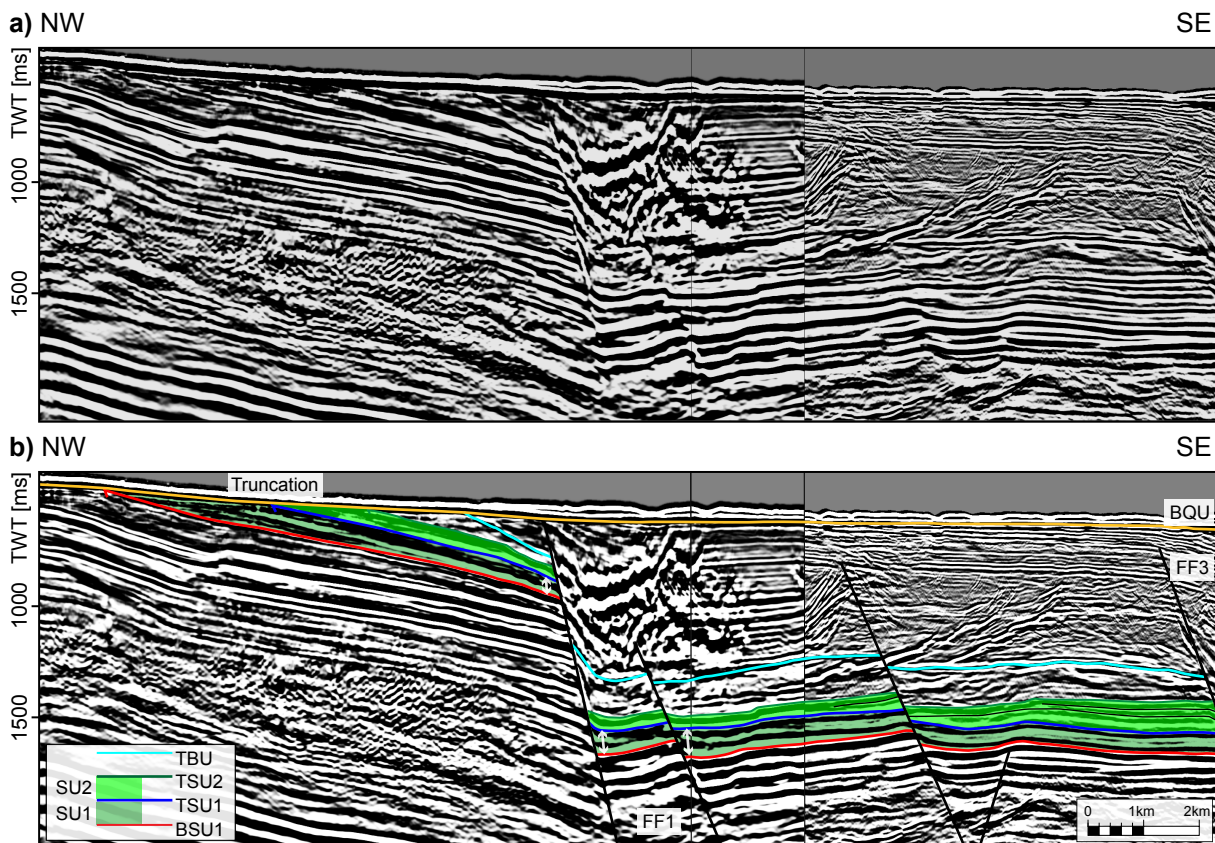


**Fig. 4.2** Crossline 1 Uninterpreted (a) and interpreted (b) northwest-southeast crossline (see Fig. 4.1 for orientation of the line). BQU = Base Quaternary Unconformity, TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF1 = Fault Family 1, FF3 = Fault Family 3. The SU1 show a thickness growth in the hanging wall of FF1. The SU2 display a thickness increase from the NW towards the SE.

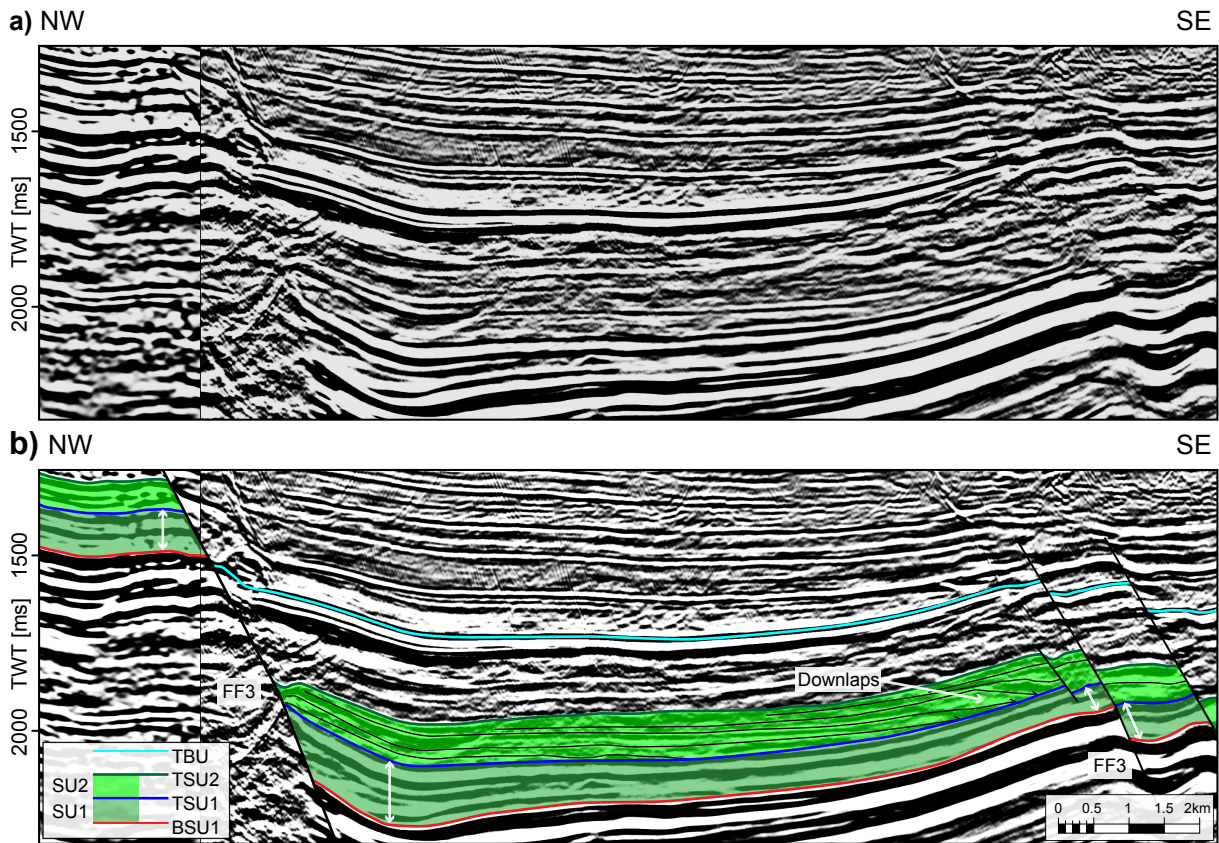


**Fig. 4.3** Crossline 2 Uninterpreted (a) and interpreted (b) northwest-southeast crossline (see Fig. 4.1 for orientation of the line). BQU = Base Quaternary Unconformity, TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF1 = Fault Family 1, FF3 = Fault Family 3. The TSU1 is not penetrated by well 7321/4-1 and the TSU2 is truncated by TBU.

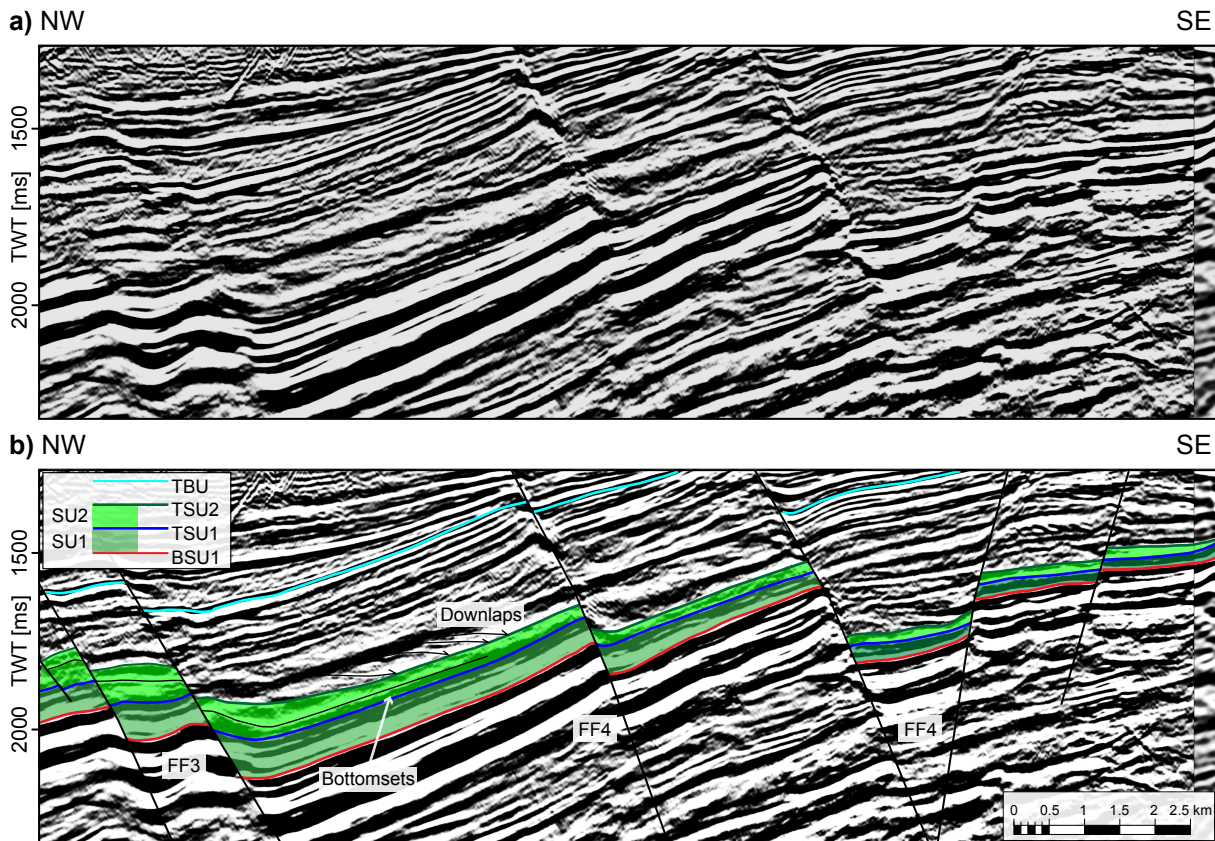




**Fig. 4.4 Composite Line 1** Uninterpreted (a) and interpreted (b) northwest-southeast composite line (see Fig. 4.1 for orientation of the line). BQU = Base Quaternary Unconformity, TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF1 = Fault Family 1, FF3 = Fault Family 3. Minor thickness differences in SU1 is observed between the footwall and hanging wall of FF1. The SU2 display a thickness increase from NW towards SE.



**Fig. 4.5 Composite Line 2a** Uninterpreted (a) and interpreted (b) northwest-southeast composite line (see Fig. 4.1 for orientation of the line). TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF3 = Fault Family 3. The SU1 show thickness growth towards the fault plane of FF3 and internal reflectors in SU2 show a sigmoidal clinoform geometry.



**Fig. 4.6 Composite Line 2b** Uninterpreted (a) and interpreted (b) northwest-southeast composite line (see Fig. 4.1 for orientation of the line). TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF3 = Fault Family 3, FF4 = Fault Family 4. The SU1 thins towards the SE and reflectors above SU2 downlaps onto the the TSU2.

## 4.2 Age Model and Seismic Unit Definition

The age of the Lower Cretaceous succession in Fingerdjupet Subbasin spans from ?Hauterivian to middle Albian (Fig. 4.7, Fig. 4.8, Fig. 4.9 and Fig. 3.1). The Jurassic-Cretaceous transition is only indicated in well 7321/4-1 at 1145 m (Fig. 4.7a and Fig. 3.1).) and has not been identified in the other wells (Fig. 4.7- Fig. 4.9 and Fig. 3.1). A unconformity is present in the base of the ?Hauterivian in well 7321/7-1 (Fig. 4.7a and Fig. 3.1) and the early Barremian in well 7321/4-1, 7321/8-1 and 7321/9-1 (Fig. 4.7a-Fig. 4.9 and Fig. 3.1). A second unconformity is present in the top of an interval dated ?Hauterivian–late Barremian which is penetrated by all of the wells (Fig. 4.7a-Fig. 4.9 and Fig. 3.1). The upper part, above the second unconformity, is dated early Aptian to middle Albian (Fig. 4.7a-Fig. 4.9 and Fig. 3.1).

# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

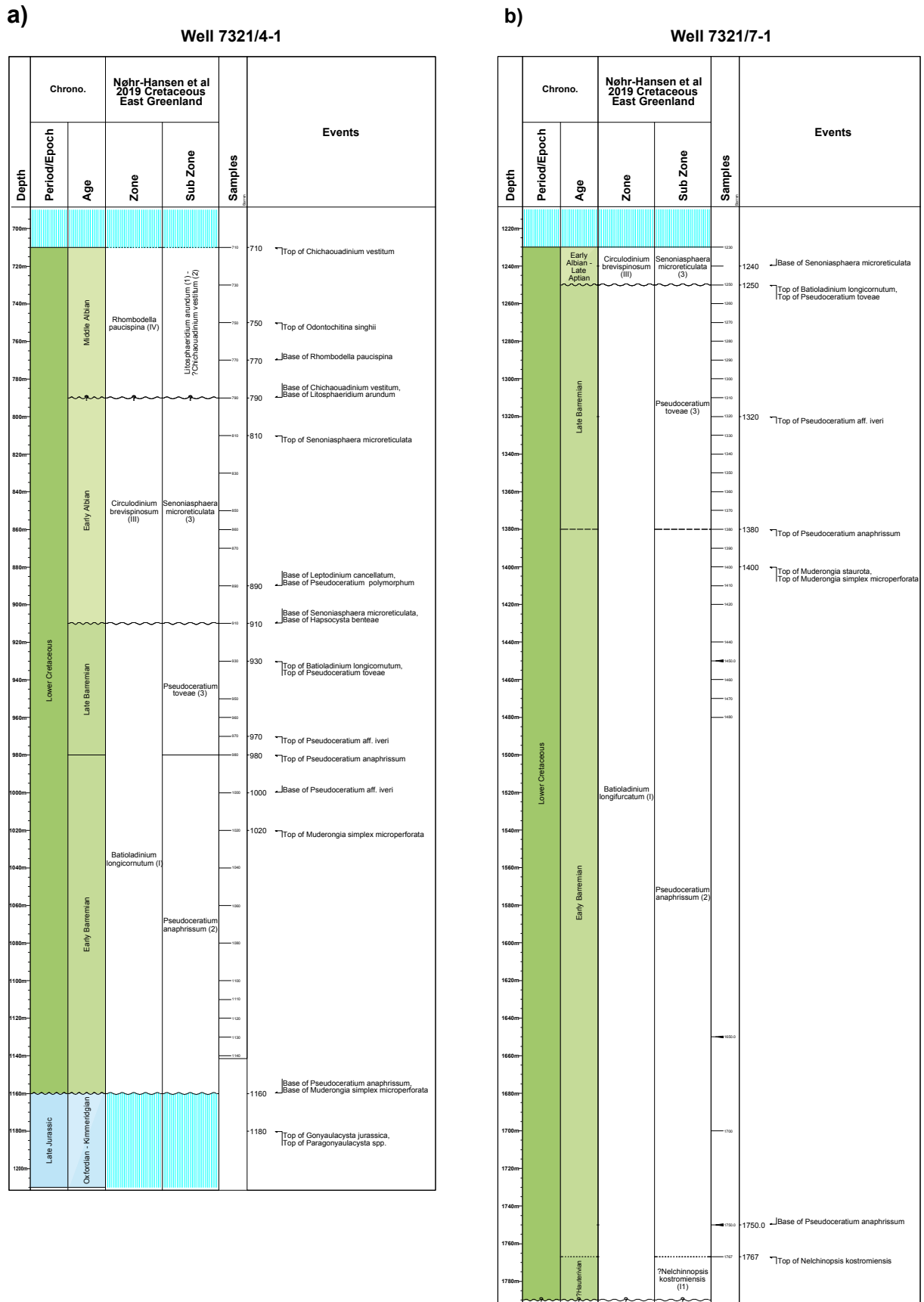


Fig. 4.7 Biostratigraphy for well 7321/4-1 and 7321/7-1

# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

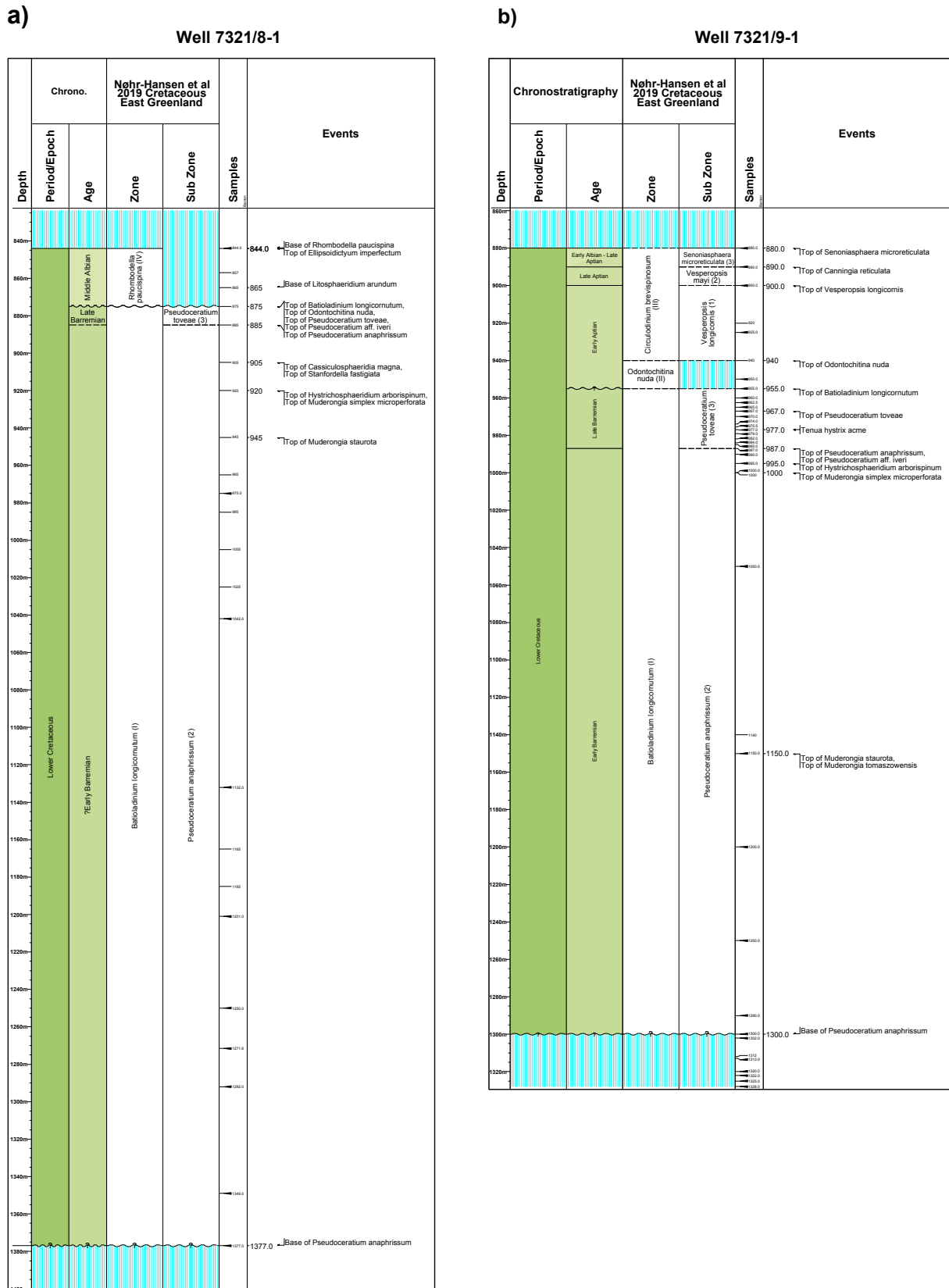


Fig. 4.8 Biostratigraphy for well 7321/8-1 and 7321/9-1

### Well 7322/7-1

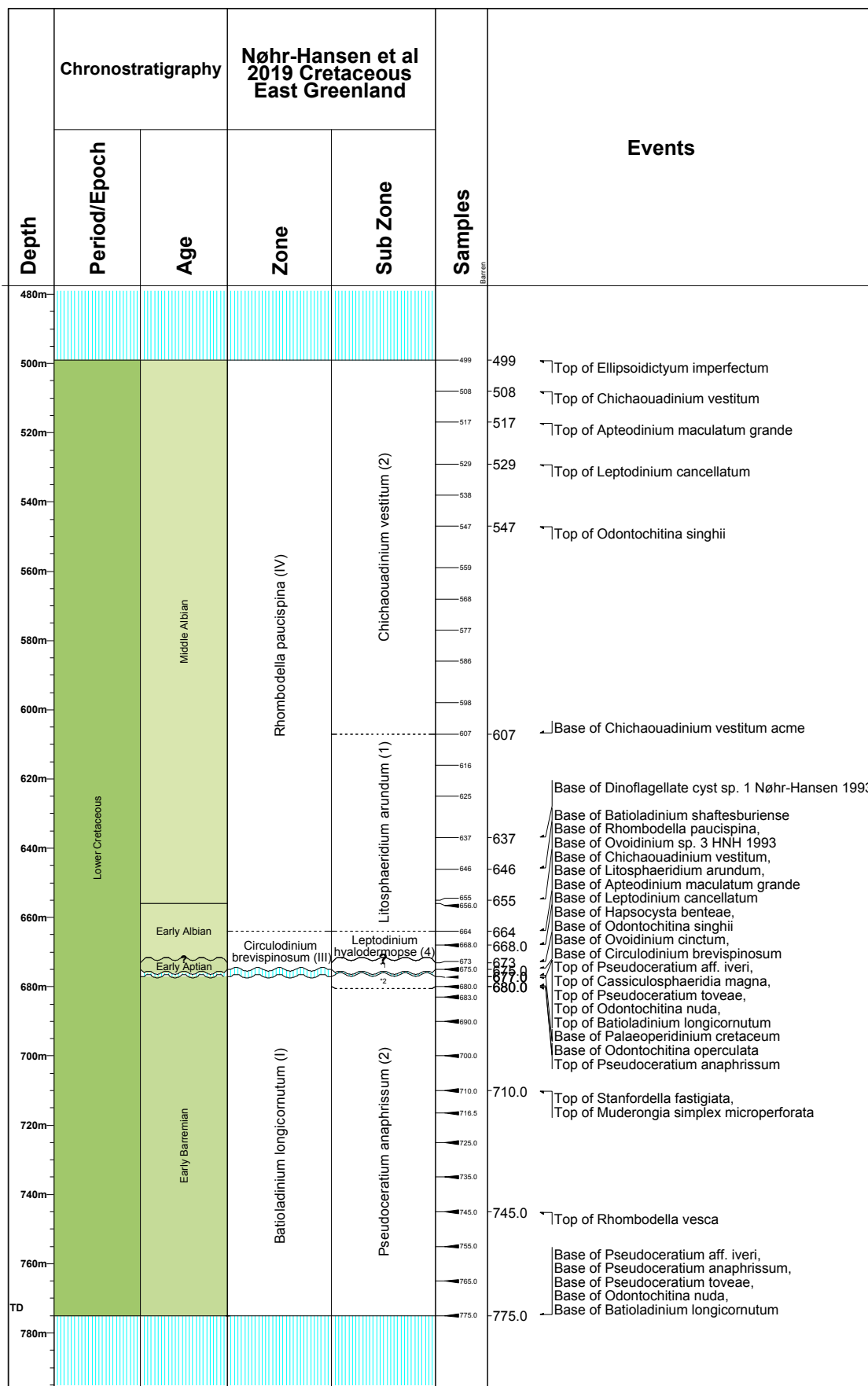
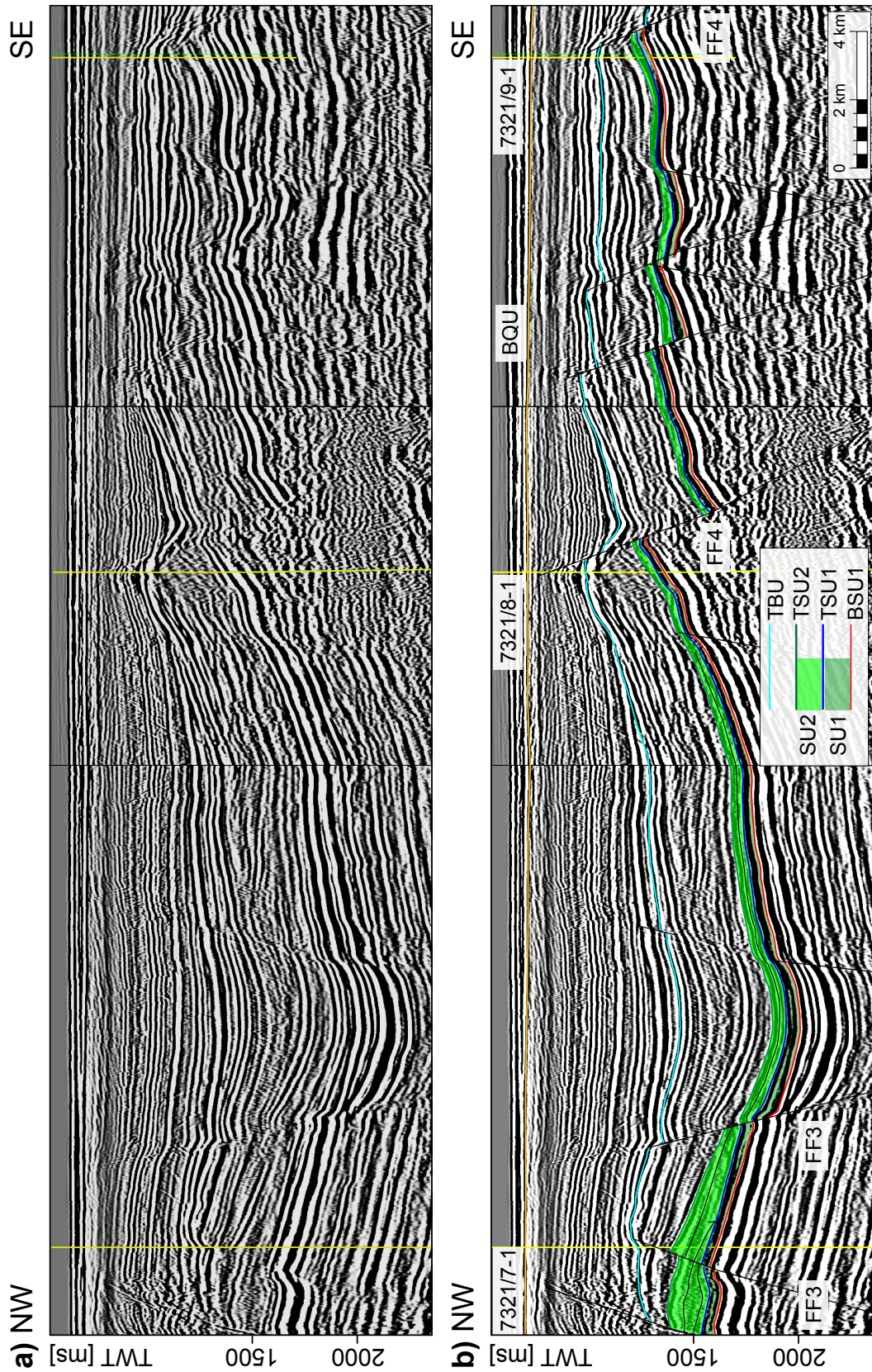


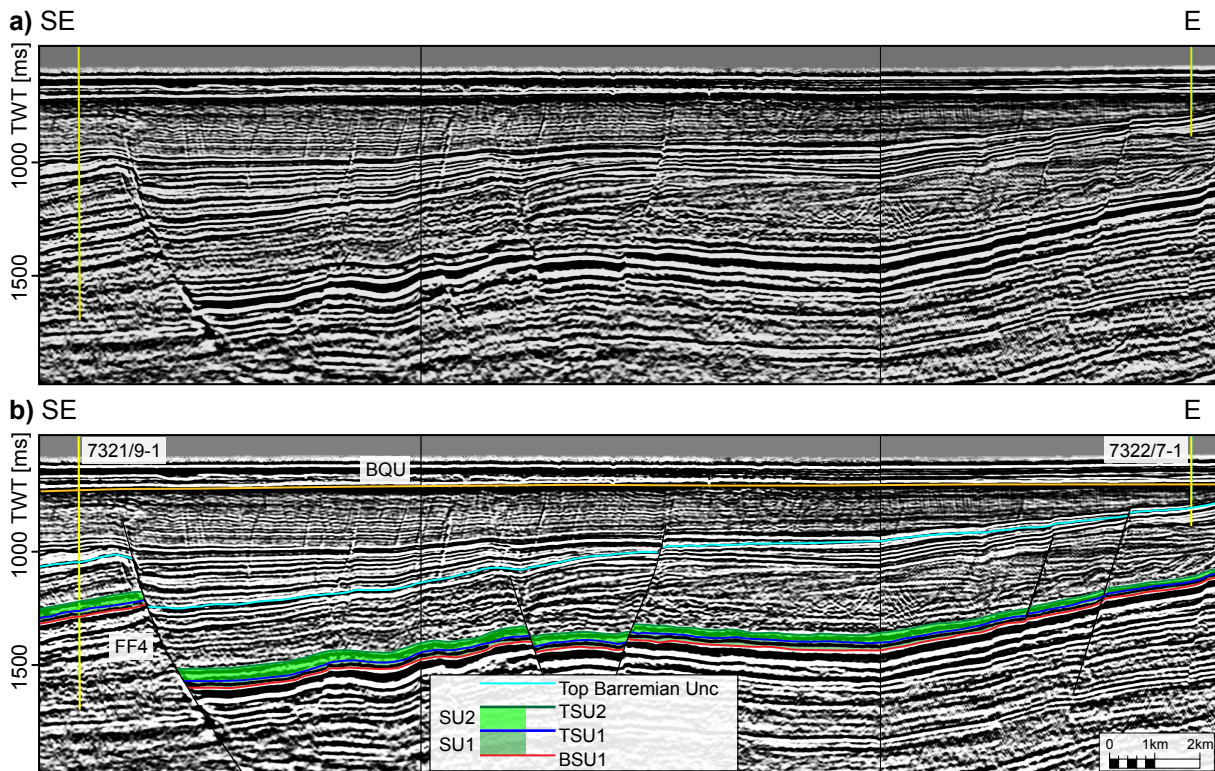
Fig. 4.9 Biostratigraphy for well 7322/7-1

The seismic units are penetrated by all of the wells except for well 7322/7-1 (Fig. 3.1, Fig. 4.3, Fig. 4.10 and Fig. 4.11). The seismic horizon defining the base of unit 1 (BSU1) is a seismic surface near the regional Base Cretaceous Unconformity characterized as a reflector with medium to high amplitude. For BSU1, the gamma ray log show high values with increasing and decreasing trends at the pick of this horizon (Fig. 3.1). The age of BSU1 is early Barremian in well 7322/4-1 and age control is missing for well 7322/7-1, 7321/8-1 and 7321/9-1 (Fig. 3.1). By comparing the BSU1 with official well tops from NPD, the BSU1 correlates to the Kolje Formation (Barremian to early Aptian) in well 7321/7-1 and Hekkingen Formation (late Oxfordian/early Kimmeridgian to Ryazanian) in well 7321/8-1 and 7321/9-1 (NPD, 2020), suggesting a Ryazanian to early Barremian age for BSU1 (Fig. 3.1). The top of Seismic Unit 1 (TSU1) is characterized as a reflector with medium to high amplitude and seismic reflectors downlaps onto the top of this surface (Fig. 4.5, Fig. 4.10). High gamma ray values with decreasing and increasing trends are observed for the TSU1 (Fig. 3.1). The age of TSU1 spans from ?Hauterivian to early Barremian, suggesting a Ryazanian to early Barremian for SU1 (Fig. 3.1).



**Fig. 4.10 Composite Line 3** Uninterpreted (a) and interpreted (b) northwest-southeast composite line (see Fig. 4.1 for orientation of the line). TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1 = Base Seismic Unit 1, FF3 = Fault Family 3, FF4 = Fault Family 4. The thins in the footwall of FF3. Well 7321/7-1 penetrates the lower part of the foresets in SU2. Well 7321/8-1 and 7321/9-1 penetrates the bottomsets.





**Fig. 4.11 Composite Line 4** Uninterpreted (a) and interpreted (b) southeast-East composite line (see Fig. 4.1 for orientation of the line). TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF3 = Fault Family 3, FF4 = Fault Family 4. Well 7322/7-1 do not penetrate the units.

The top of Seismic Unit 2 (TSU2) is characterized by a reflector with medium to high amplitude and seismic reflectors downlaps onto the top of this surface in the central (Fig. 4.6) and eastern part of the study area. The gamma ray character of the TSU2 varies from a peak of high gamma ray values in well 7321/4-1 to increasing and no change in the log response in the other wells (Fig. 3.1). The TSU2 is aged early Barremian to late Barremian, suggesting a ?Hauterivian to late Barremian age for SU2 (Fig. 3.1).

## 4.3 Seismic Unit 1

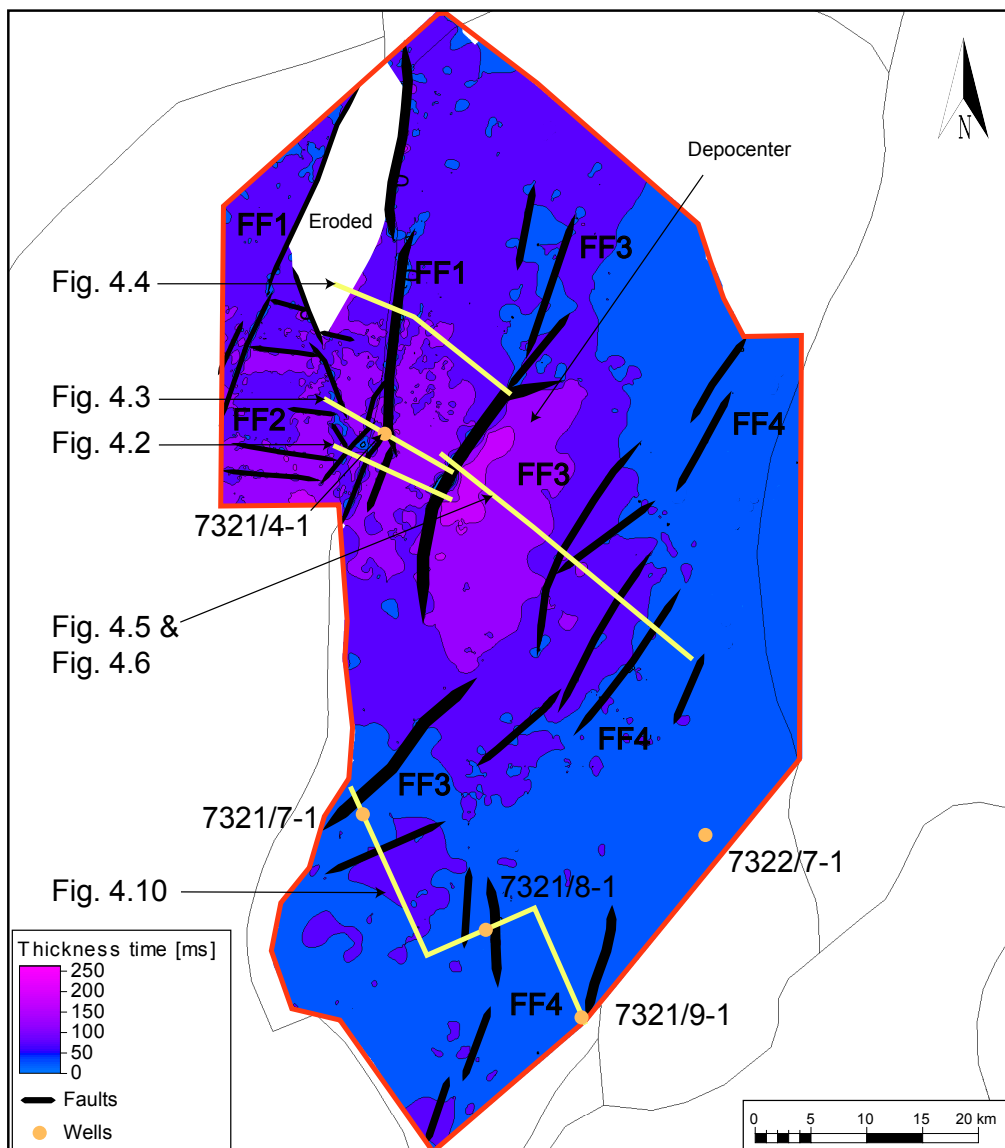
### Observation

#### Well Character

The gamma ray pattern in well 7321/4-1 and 7321/7-1 show two intervals of coarsening to fining upward trends followed by a coarsening trend in the top (Fig. 3.1). Drill cuttings and SWC from these wells show sand, silt and caly in addition to dolomite, limestone and marl. Well 7321/8-1 and 7321/9-1 display overall higher reading with a signature characterized by low gamma ray values in the lower part followed by a fining to coarsening trend in the upper part. The SWC in these wells are dominated by fine-grained sediments.

**Thickness Variations**

Seismic unit 1 show a thickness decrease from the from 110 meter in well 7321/4-1 to 90 meter in well 7321/9-1 in the east (Fig. 3.1). The time thickness map indicates that the unit is thickest in the north-western part of the study area and thins towards the southeast (Fig. 4.12). A thickness increase is observed towards the fault plane of FF3 in the southwest (Fig. 4.10) and in the center of the study area, where the thickness reaches around 200 ms (Fig. 4.5 and Fig. 4.12). The thickness of the unit increase in the hanging wall of the southern segments of FF2 (Fig. 4.12) and locally in the hanging wall of FF1 (Fig. 4.2, Fig. 4.4 and Fig. 4.12). No thickness variations is observed in hangingwall of FF4 (Fig. 4.6, Fig. 4.10, Fig. 4.11 and Fig. 4.12).



**Fig. 4.12 Time Thickness Map of SU1** Time thickness map of Seismic Unit 1 (SU1), fault families and location of seismic lines.

### Seismic Character

The SU1 is truncated by the BQU in footwall of FF1 in the northwest and is not present in the northeast trending horst structure defined by FF1 in north-western part of the study area (Fig. 4.12 and Fig. 4.4). The internal reflectors of SU1 vary from discontinuous to continuous reflectors with low to high amplitude (e.g. Fig. 4.2) and characterized by sub-parallel reflector patterns. Growth wedges related to FF3 occur in the central part (Fig. 4.5) and in the southwest ( Fig. 4.10). Reflectors downlap onto the top of SU1 in the northeastern, central (Fig. 4.5) and southwestern part of the study area (Fig. 4.10).

### Interpretation

The SU1 unit was deposited during active faulting controlled by FF2 and locally by FF1 in the north-western part of the study area. A depocenter developed in the central part of the basin and a syn-rift wedge developed along the fault plane of FF3. The constant thickness of the unit in the southeastern part indicates that FF4 was not active (Fig. 4.12). The discontinuous to continuous reflectors with sub-parallel geometry indicate stable and unstable low energy during deposition of SU2. The SU1 comprises dominantly fine-grained rocks with the occurrence of carbonates, which could indicate a shallow to deep marine environment. Sandstone is encountered by the wells in the west (7321/4-1 and 7321/7-1) which indicate a source of coarse-grained sediments existed for this unit. The source of these sediments was most likely in the west or north based on the southeastward thinning of the unit.

## 4.4 Seismic Unit 2

### Observation

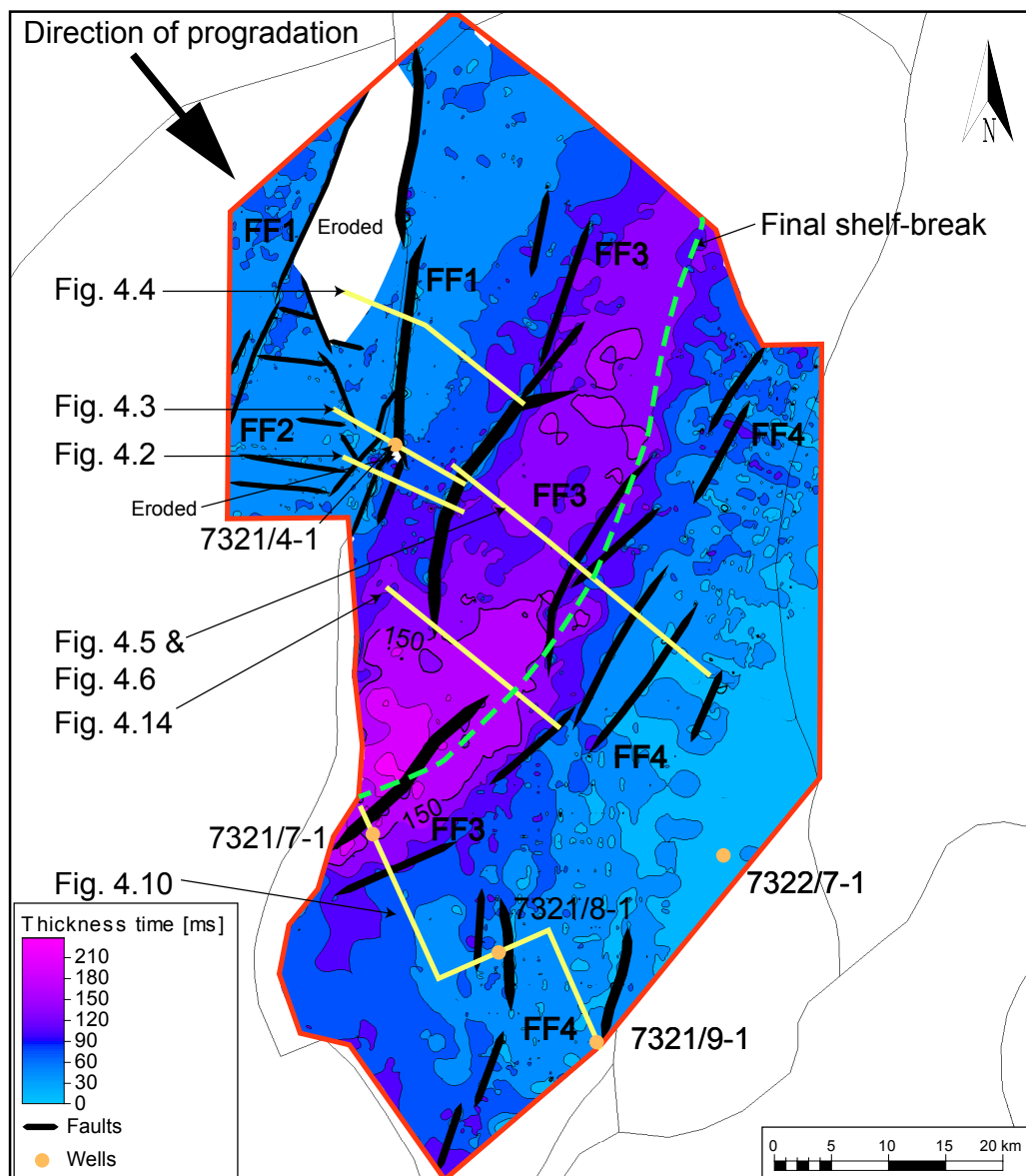
#### Well Log Character

Well 7321/4-1 display a coarsening to fining upward trend in the lower part (Fig. 3.1). The upper part is characterized by a coarsening upward trend followed by a serrated signature with intervals of low gamma ray and a fining upward trend reaching a peak of high values in the top of the SU2 (TSU2). Drill cuttings from the same interval includes clay, silt, sand and limestone. The lower part of well 7321/7-1 shows two intervals of coarsening to fining upward trends, followed by an interval with a serrated signature and two intervals of coarsening to fining upward trends in the upper part. The sidewall cores from well 7321/7-1

shows limestone in the lowermost part followed by alternating siltstone and claystone. The gamma ray pattern for well 7321/8-1 and 7321/9-1 is characterised by a coarsening upward trend followed by a serrated signature. The sidewall cores from these wells is dominated by shale and claystone.

**Thickness Variations**

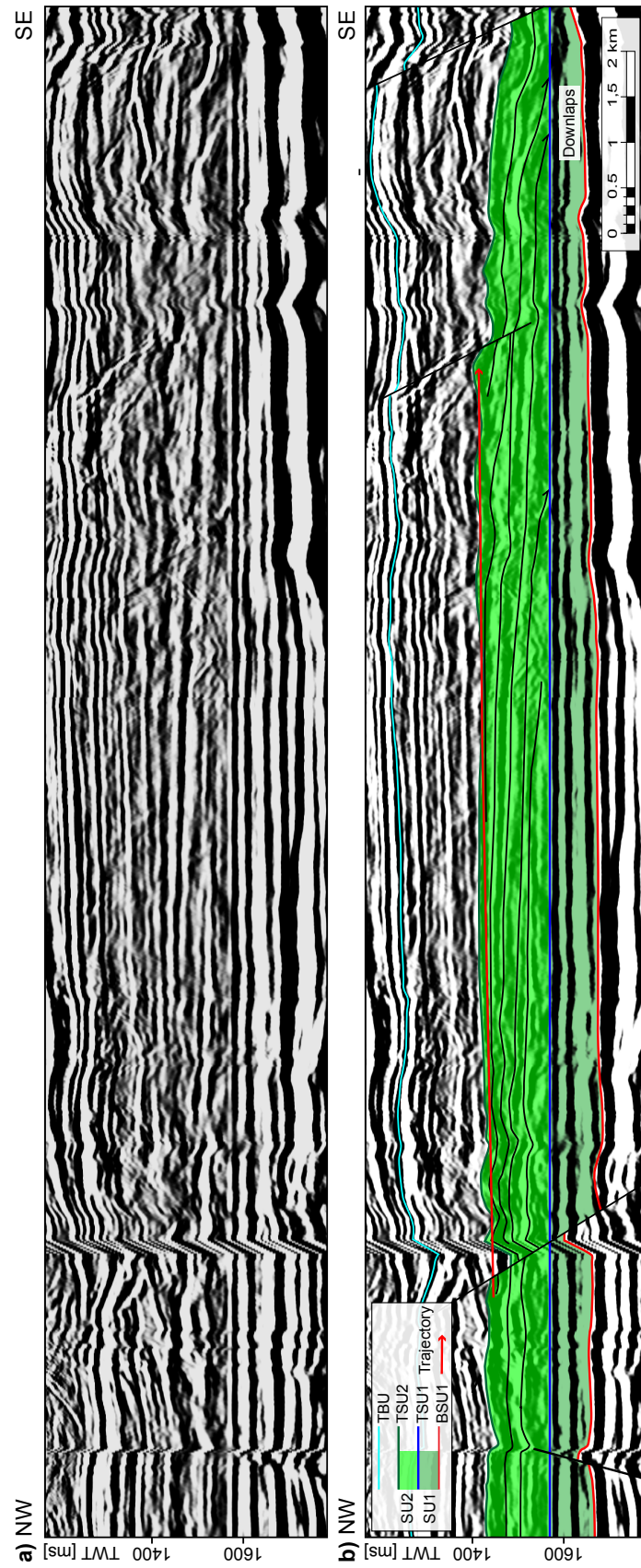
The thickness increase from the northwest to a depocenter located in central axis of the basin and thins towards the southeast of the study area (Fig. 4.13). The same trend is observed in the wells, thickens from 56 meter thick in well 7321/4-1 to 261 meter in well 7321/7-1 and thins to 75 meter in well 7321/9-1 (Fig. 3.1). The SU2 show a gradual increasing to decreasing thickness trend.



**Fig. 4.13 Time Thickness Map of SU2** Time thickness map of Seismic Unit 1 (SU2), fault families and location of seismic lines.

### Seismic Character

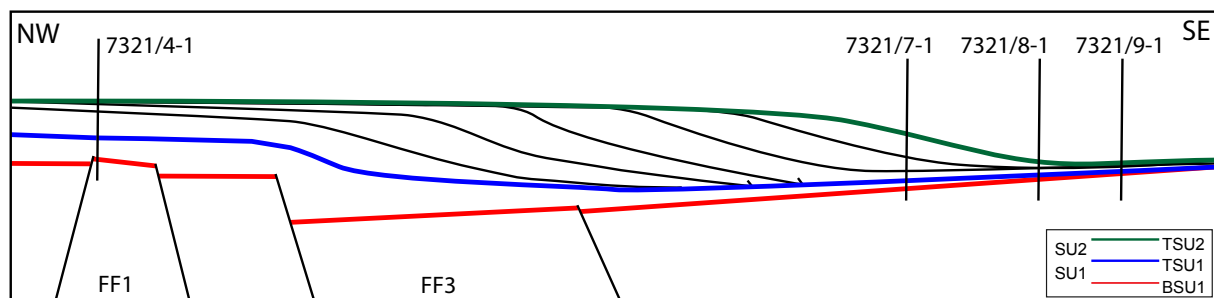
The TSU2 is truncated by the TBU in BQU in the footwall of FF1 in the northwest and is not present in the structural high defined by FF1 in the northwest (Fig. 4.4 and Fig. 4.13). The internal reflectors of SU2 consists of continuous reflectors with medium to high amplitude in the north-western part of the study area (Fig. 4.2-Fig. 4.4). Further southeast, subparallel to gently dipping continuous reflectors occur with low to medium amplitude (Fig. 4.5, Fig. 4.6, Fig. 4.10, Fig. 4.11). These reflectors displaying a sigmoidal clinoform geometry prograding towards the southeast. The upper boundary of the clinoforms is characterized by thin topsets with medium to high amplitude and low angle foresets which display a rising trajectory (Fig. 4.14). The clinoforms downlap onto the TSU1 in the central part of the study area and bottomsets are observed further southeast-east in the the study area (Fig. 4.5, Fig. 4.6, Fig. 4.10, Fig. 4.11).



**Fig. 4.14 Crossline 3** Uninterpreted (a) and interpreted (b) northwest-southeast crossline line (see Fig. 4.13 for orientation of the line). TBU = Top Barremian Unconformity, TSU2 = Top Seismic Unit 2, SU2 = Seismic Unit 2, TSU1 = Top Seismic Unit 1, SU1 = Seismic Unit 1, BSU1= Base Seismic Unit 1, FF3 = Fault Family 3, FF4 = Fault Family 4. The crossline is flattened on TSU1. Internal reflectors in SU2 downlaps onto the TSU1.

## Interpretation

The sigmoidal clinoform geometries are interpreted to represent clinoforms prograding towards the southeast. Well 7321/4-1 is interpreted to penetrate clinoform topsets (Fig. 4.15) consisting of shales coarsening upwards into silty sandstones overlaid by shales and possibly carbonates. This stacking pattern reflects a prograding trend in the lower part followed by a middle aggrading trend and an upper retrograding trend. Well 7321/7-1 is interpreted to penetrate the middle slope which consists of alternating siltstone and claystone. The final shelf-break occurs a few km north of well 7321/7-1 (Fig. 4.13). The bottomsets are interpreted to be penetrated by well 7321/8-1 and 7321/9-1, fine-grained basin floor deposits. The southeast prograding direction of the clinoforms indicates that the source of sediment was located in the northwest. Based on the gradual increasing to decreasing thickness trend of the unit, the faults in the study area were not active during deposition of SU2.



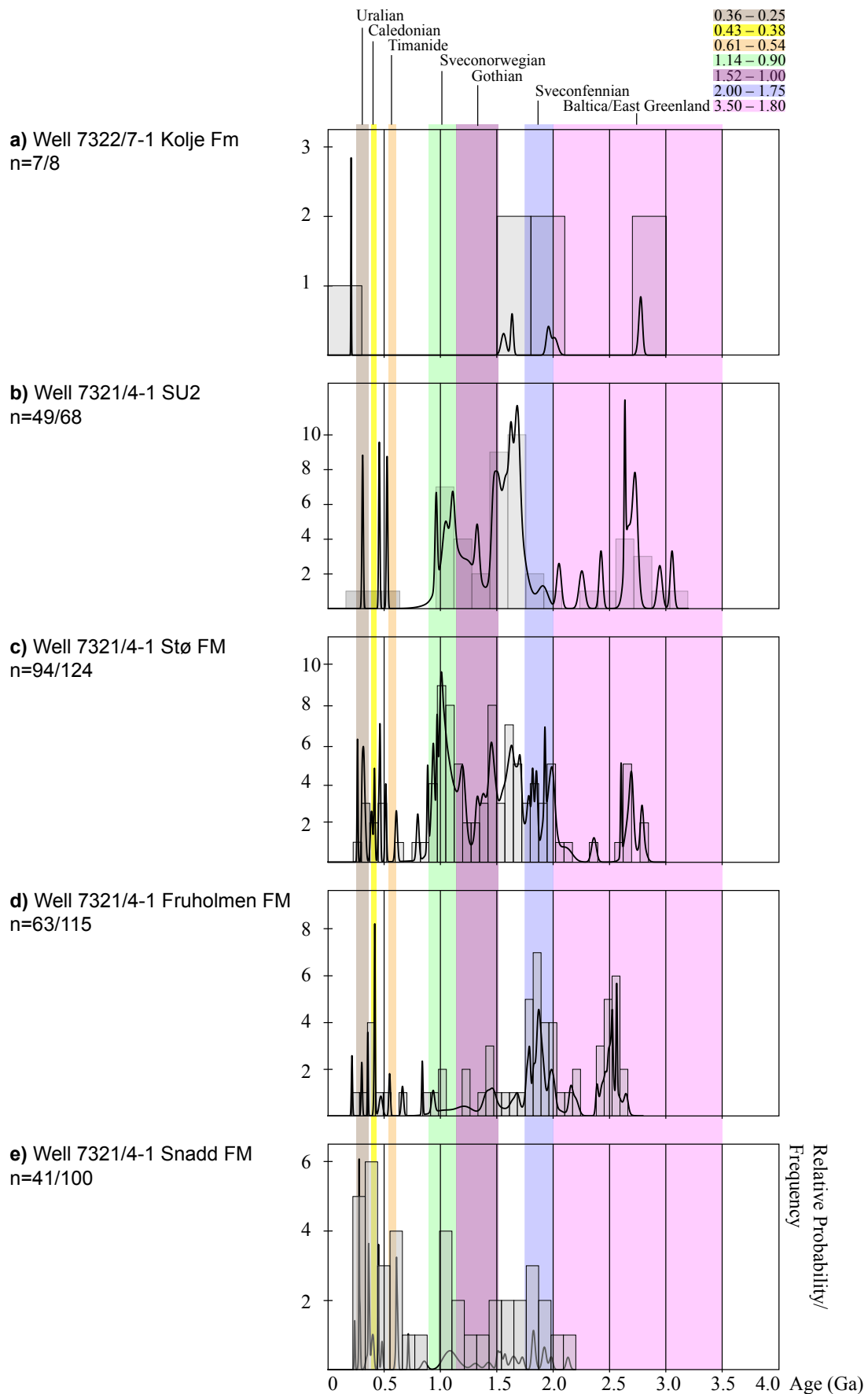
**Fig. 4.15 Interpretation of SU2** Conceptual sketch of a cross-section showing the SU1 and the SU2 just after deposition of S2. The wells have been projected to show where they are inferred to penetrate the clinoforms. Well 7321/4-1 is inferred to penetrate the topsets, Well 7321/7-1 penetrates the middle part of the foresets. Well 7321/8-1 and Well 7321/9-1 penetrate the bottomsets.

## 4.5 Detrital Zircon Ages

### Snadd FM (Middle – Upper Triassic)

The detrital zircon age spectra within the Snadd FM samples in well 7321/4-1 lack grains with Archean ages and display two main younger age populations (Fig. 4.16). A population from 2 to 1 Ga with similar frequency, except for the populations from 1.1 to 1 Ga (Late Mesoproterozoic). The second population, between 0.6 – 0.24 Ga (Late proterozoic – Early Mesozoic), are all of similar frequency.

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**Fig. 4.16 Detrital Zircon Ages** Zircon data from well 7321/4-1 and 7322/7-1. Probability density plots and histograms. *n* = number of <10% discordant samples/total number of samples.



### **Fruholmen FM (Upper Triassic)**

The Fruholmen FM samples (well 7321/4-1) contain abundant grains of Archean ages from 2.6 to 2.4 Ga (Fig. 4.16). The Proterozoic age distribution show a significant population between 1.75 – 2.0 Ga (late Paleoproterozoic) and smaller amounts of 1 to 1.5 Ga (late Mesoproterozoic). Grains of 0.47 Ga to 0.4 Ga (mid-Ordovician to early Devonian) are abundant among the younger dated samples.

### **Stø FM (Lower – Middle Jurassic)**

The Stø Formation (well 7321/4-1) consist of Archean dated grains in the interval between 2.8 – 2.6 Ga (Fig. 4.16). The Proterozoic age distribution is dominated by grains dated between 2.0 – 1.0 Ga characterized by two main age populations. The oldest population, between 2 – 1.3 Ga, show the highest frequency at c. 1.4 Ga (early Mesoproterozoic) and the younger population, between 1.2 – 0.9 Ga, display the highest frequency from 1.1 Ga to Ga (late Mesoproterozoic). The most significant population among the youngest dated grains are the intervals 0.46 – 0.4 Ga (mid-Ordovician – Early Devonian) and 0.4 – 3.0 Ga (Early Devonian – Late Carboniferous).

### **Seismic Unit 2 and Kolje FM (Lower Cretaceous)**

The SU2 (well 7321/4-1) have highest frequency of Archean aged grains in the 2.6 – 2.75 Ga interval (Fig. 4.16). The dominant age populations in the Proterozoic ages range from 1.7 to 1.5 Ga (Paleo-Mesoproterozoic) and between 1.2 – 1 Ga (late Mesoproterozoic). Three younger dated grains are of 0.52 Ga (late Cambrian), 0.45 Ga (late Ordovician) and 0.30 Ga (late Carboniferous) age.

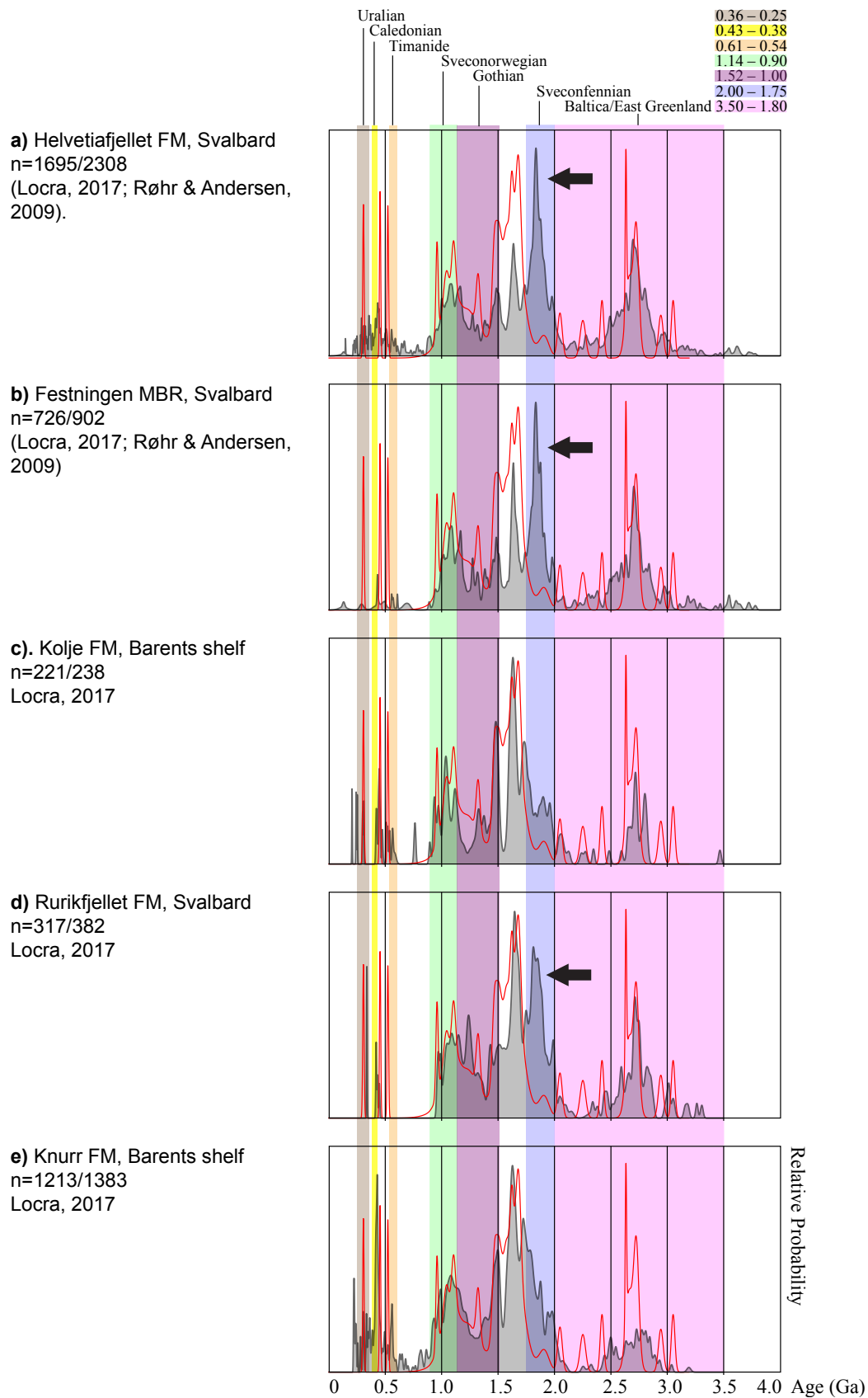
Kolje FM in well 7322/7-1 contains seven dated grains. The two oldest are dated Archean (c. 2.8 Ga), followed four Proterozoic aged grains (2, 1.9, 1.6, and 1.5 Ga) and the youngest sample is of Late Triassic age (0.21 Ga).

## 5 DISCUSSION

### 5.1 Provenance Evaluation and Implications for Reservoir Development

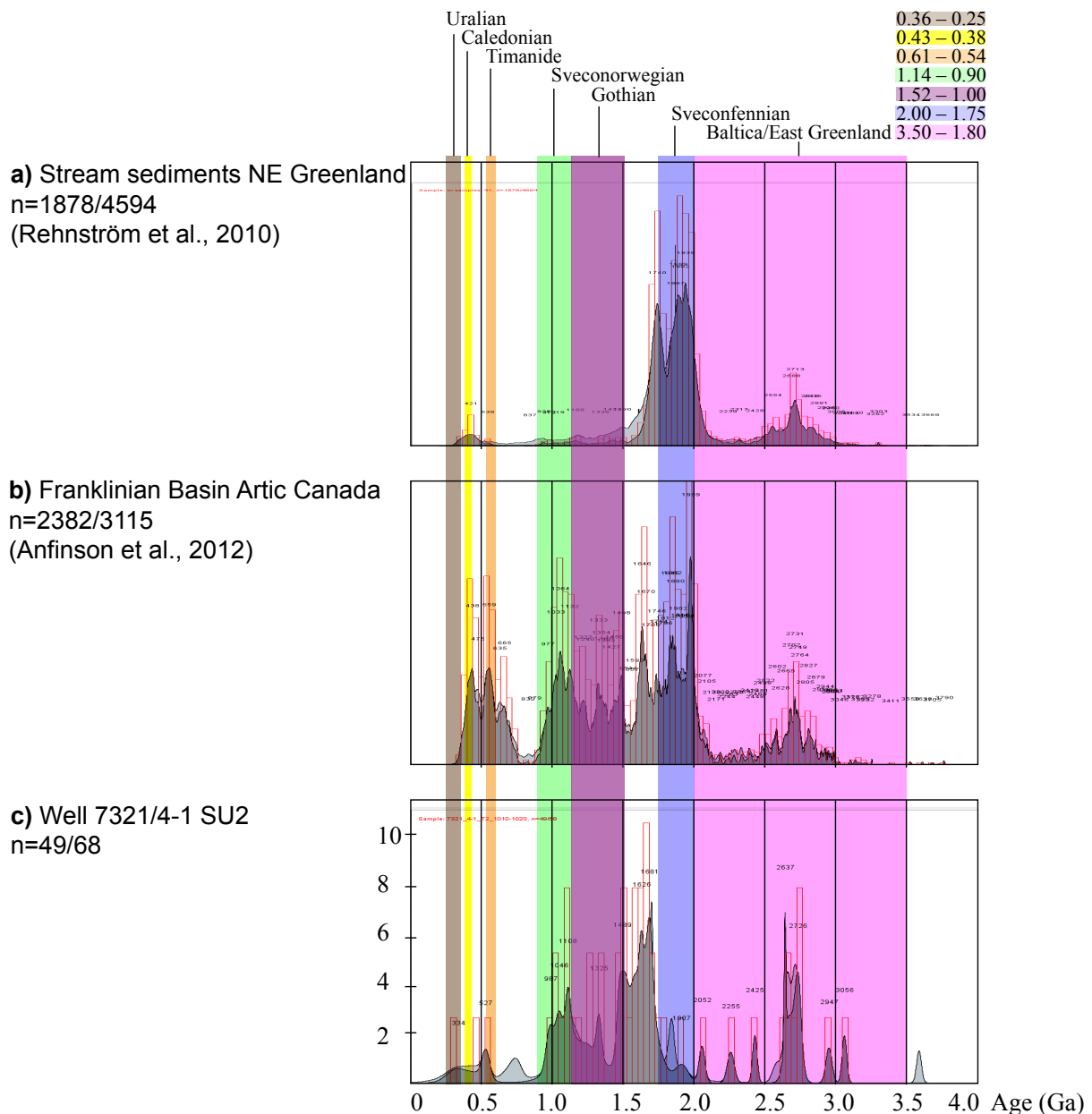
The northern part of Greenland and Canada (Laurentia) are potential source regions situated in the north-western corner of the Barents shelf during the Early Cretaceous. Because the clinofolds observed in the ?Hauterivian to late Barremian aged unit (SU2) in this study evolved from the northwest to the southeast in the Fingerdjupet Subbasin. The detrital zircon abundances of the studied SU2 unit (Well 7321/4-1) is similar to the zircon age signatures in Lower Cretaceous strata on Svalbard including Rurikfjellet Formation, Festningen Member and Helvetiafjellet Formation (Fig. 5.1 a,b and d). The main difference is the 1.8–1.9 Ga age population in the Svalbard samples appears to be different (Fig. 5.1). The Helvetiafjellet Formation on Svalbard, Lower Cretaceous sediments in Wandel Sea Basin (northern Greenland) and Sverdrup Basin (Arctic Canada) also shares similar detrital zircon abundances (Røhr et al., 2008, 2010; Røhr & Andersen 2009). Detrital zircon ages from Lower–Middle Triassic sediments on Svalbard also show similar age distribution as the Helvetiafjellet FM (Bue and andersen, 2013). In addition to older deposits, Cambrian aged sediments in Northwest Territories (northern Canada, Hadlari et al., 2012), Neoproterozoic to late Devonian aged sediments in Franklinian Basin (Arctic Canada, Anfinson et al. 2012) and Mesoproterozoic to early Cambrian sediments in Peary Land (North Greenland, Kirkland et al. 2009), display similar age distribution as the Helvetiafjellet FM and Lower–Middle Triassic sediments on Svalbard (Bue & andersen, 2013). The provenance of the Lower–Middle Triassic (Svalbard) and Lower Cretaceous (Svalbard, Wandel Sea Basin and Sverdrup Basin) detritus are interpreted to be older deposits in the Franklinian Basin (Arctic Canada) and the North Greenland (Fig. 5.2, Røhr et al., 2008, 2010; Røhr & Andersen 2009; Bue% Andersen, 2013). The similar detrital zircon abundances of the Barremian strata in the Fingerdjupet Subbasin (this study), the Lower Cretaceous succession in Svalbard, Wandel Sea Basin and Sverdrup Basin indicate that they originate from the same source area located in the northern Greenland and Franklinian Basin (Fig. 5.2) The reason why the dominant 1.8–1.9 Ga peak is not present in studied unit could be due to the distribution of the detritus characterized 1.8–1.9 Ga protosource could be limited to the north and did not reach the Barents shelf.

# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development



**Fig. 5.1 Lower Cretaceous Comparison of Zircon Ages** The red curve represents the SU2 in well 7321/4-1.

Probability density plots and histograms.  $n$  = number of <10% discordant samples/total number of samples.



**Fig. 5.2 NE Greenland and Franklinian Basin** a) Stream sediments in NE Greenland. b) Franklinian Basin. c) SU2

The uplift and present day structuring of the Stappen high is suggested to occur during early Cenozoic (Faleide, 1993; Worsley et al., 2001, Blaich et al., 2017). However, the northern part of Stappen High (not open for petroleum exploration) is poorly constrained (e.g. Anell et al., 2016; Blaich et al., 2017) and could represent a potential source region. Permian to Middle Triassic rocks outcrops in the highest point of Stappen High (Bjørnøya, Vigran et al. 2014) and the Middle Jurassic Stø Formation is interpreted to be present in the southern margin of the Stappen High (Blaich et al., 2017). The detrital zircon abundances of the studied SU2 unit is similar to age signatures of the Stø Formation, in well 7321/4-1 ( Fig. 4.16) and in the Barents shelf (Klausen et al., 2017), Knurr Formation and Kolje Formation (Fig. 5.1 c and e). Exposure and recycling of underlying Stø Formation on the northern Stappen High could explain the

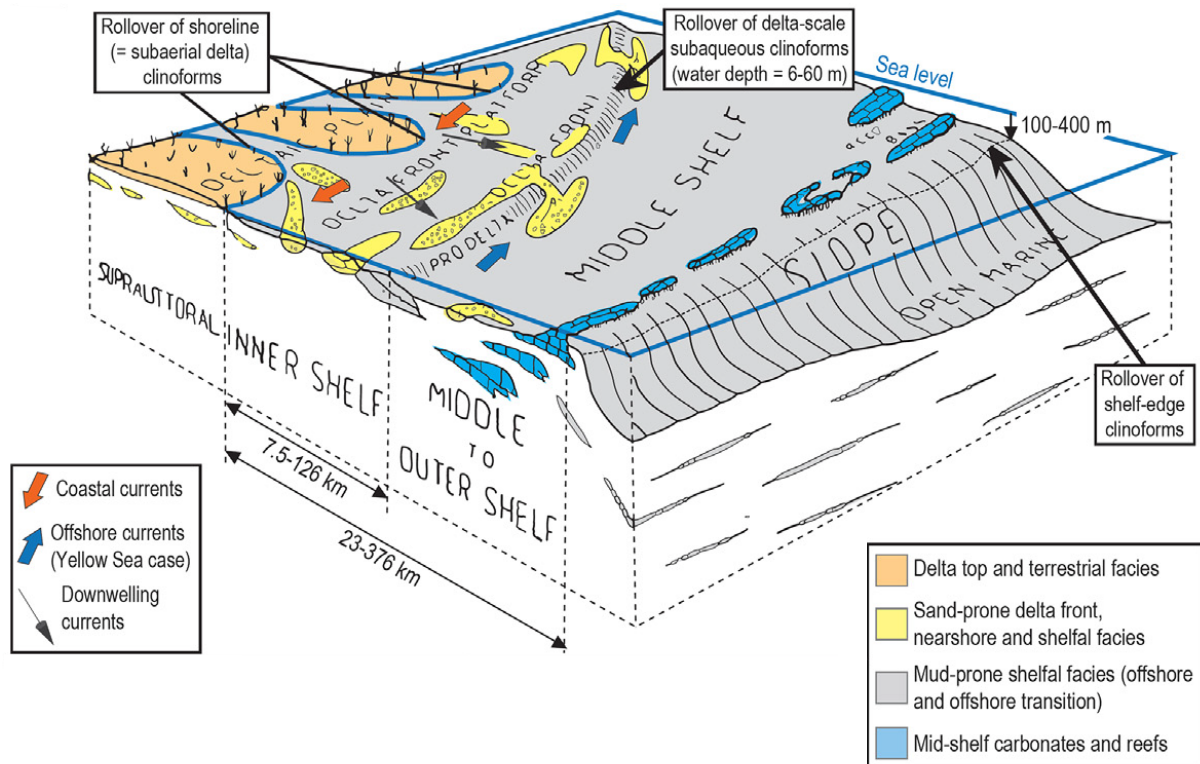
similar detrital zircon abundances observed in the studied SU2 unit (Well 7321/4-1) and in Knurr and Kolje formations. This may also explain why the dominant 1.8–1.9 Ga age population observed in the Lower Cretaceous succession on Svalbard is not present in the SU2 unit, Knurr and Kolje formations (Fig. 5.1). The inferred northern position of the northern Stappen High also supports the southeast prograding direction of the clinoforms in SU2. Although, petrographic and geochemical analysis performed on the Knurr and Kolje formations in the Hammerfest and Tromsø basins do not indicate a complete recycling of younger or older strata (Locra, 2017). Locra (2017) suggest Baltica as the main provenance for the Lower Cretaceous succession in addition to a exotic source area located to the north of the Barents shelf. The exotic source is characterized by 0.4 – 0.225 Ga (Early Devonian – Late Triassic) aged sediments, intermediate to mafic detritus and strongly reworked material of Gothian and Neoproterozoic age (Locra, 2017).

The Loppa high is another potential source area situated in the southeastern margin of Fingerdjupet Subbasin (Fig. 1.1) The uplift the Loppa High is interpreted to reach a peak during Barremian (Indrevær et al., 2016) and may provide sediments influx from the southeastern margin of the Fingerdjupet Subbasin. However, based on the southeast prograding direction of the clinoforms in SU2, the Loppa High was unlikely the the source of SU2.

### Implications for Reservoir Development

The clinoforms penetrated by well 7231/7-1 suggesting they are of shelf-edge delta scale (relief between 100 - 1000 m, e.g. Patruno et al., 2015) based on the thickness in well 7321/7-1 (261m) penetrating the middle part of the slope . The topset segment of shelf-edge delta clinoforms includes paralic to shallow marine facies which can be prolific reservoirs (Porebski and Steel 2003). Transport of sediment across the shelf is largely controlled by the successive migration of repeated regressive-transgressive delta scale clinoforms and draping clinoforms (compound clinoform systems, Patruno Helland-Hansen, Fig. 5.3) The interaction of mud-prone draping with mud-prone and sand-prone shoreline progradation gives a mixed facies relationship within shelf prism clinoforms (, Patruno & Helland-Hansen, 2018).

## Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development



**Fig. 5.3 Depositional Environment** 3D cartoon illustrating facies related to a shoreline, subaqueous delta-scale and time equivalent shelf-edge clinoform system (figure modified from Patrunoa & Helland-Hansen, 2018).

Well 7321/4-1 penetrates the topset segment of the clinoforms at a position which could represent inner to middle shelf environment. The overall coarsening to fining upward trend in SU2 can indicate a regressive lower part followed by a transgressive interval in the top. This trend may represent a shoreline system followed by transgressive shales in the top. Thus, be a system of a potentially laterally extensive shorelines systems. Example, the Battfjellet Formation (Eocene) on Svalbard consists of 20 sand-prone shelf-margin clinoforms. The topsets of these clinoforms consists of a regressive lower part of fluvial-wave dominated delta-front deposits overlain by tide-influenced estuarine deposits (Steel and Olsen, 2002). Thick sand wedges can develop at the shelf-edge consisting of upward coarsening delta-front, mouth bar and channel system. Sand sheets of turbidite deposits commonly form in the upper to middle slope and deep water fans sand may accumulate in both during rising and falling sea levels (porębski & Steel, 2003). Hence shelf-edge delta systems can form good reservoirs in the shelf, slope and in the basin floor.

The only indicator of sand in the SU2 clinoforms is in well 7321/4-1 which corresponds to the topset segment of the clinoforms. The middle part of the slope is penetrated by well 7321/7-1 which show dominant fine-grain rocks. Well 7321/8-1 and well 7321/9-1 penetrated the bottomsets which is also

## **Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development**

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dominated by fine-grain rocks. There are no seismic evidence of delta scale clinoforms with steep foresets and base on dominant fine grain deposits in the slope and bottomsets penetrated by the other wells, sand is probably restricted to the topsets.

## **6 CONCLUSION**

Well 7321/4-1 penetrates a unit with relative low gamma ray values compared to the other exploration wells in the Fingerdjupet Subbasin. Palynological analyses (carried out by GEUS) indicated a ?Hauterivian to late Barremian age for this interval and seismic interpretation shows that this interval corresponds to southeast prograding clinoforms.

Detrital zircon U/Pb geochronology (carried out by GEUS) from samples in well 7321/4-1 and from in well 7322/7-1 was performed to interpret the source of these sediments. The samples from well 7321/4-1 are dominated by three age populations: (1) 2.6 – 2.75 Ga, (2) 1.7 to 1.5 Ga and (3) 1.2 – 1 Ga. While the samples from well 7322/7-1 contain only seven dated grains, thus making it difficult to interpret the origin of these siltstones. The dominant age populations observed in SU2 turns out to be quite similar to the age distribution that dominates several chronostratigraphic units in the northern Greenland, the Arctic Canada and on Svalbard. Based on the similar age signatures of SU2 with these and southeast prograding direction of the clinoforms in SU2, the source of sediments to the Fingerdjupet is suggested to be North Greenland and/or Arctic Canada.

The sandstones in well 7321/4-1 is interpreted to be the topsets of shelf-edge delta scale clinoforms that prograded from the Ringsel Ridge in the northwest and into the Fingerdjupet Subbasin. The gamma ray signature of the topsets indicates a coarsening to fining upwards stacking pattern which can potentially form good shallow marine sandstones. The lack of coarse clastics in wells penetrating the slope and basin floor segment of the system constrains the lateral distribution of sandstones in the study area.



# 7 APPENDIX

Fig. 7.1

Well	Sample Interval [MD]	Unit
7322/7-1	679-763	Kolje FM (Lower Cretaceous)
7321/4-1	1010-1020	SU2 (Lower Cretaceous)
7321/4-1	1348-1354	Stø FM (Lower – Middle Jurassic)
7321/4-1	1422-1428	Fruholmen FM (Upper Triassic)
7321/4-1	1521-1524	Snadd FM (Middle – Upper Triassic)

Kolje FM			CONCENTRATIONS <sup>a</sup>					RATIOS					AGES					Concordance				
Sample	Analysis	U [ppm]	2 σ	Th [ppm]	2 σ	Pb [ppm]	2 σ	U/Th <sup>b</sup>	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>206</sup> Pb/ <sup>238</sup> U <sup>b</sup>	2 σ <sup>d</sup>	rho <sup>c</sup>	<sup>207</sup> Pb/ <sup>206</sup> Pb <sup>b</sup>	2 σ <sup>d</sup>	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>206</sup> Pb/ <sup>238</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>207</sup> Pb/ <sup>206</sup> Pb <sup>b</sup>	2 σ <sup>d</sup>	
7322/7-1_679-763	Zircon_sample-007.FIN2	429	45	151	25	141	22	2.91	0.267	0.018	0.037	0.001	0.459	0.0526	0.0038	240	15	233	7.5	300	160	77.8
7322/7-1_679-763	Zircon_sample-008.FIN2	646	56	168	26	145	35	3.88	0.231	0.030	0.033	0.001	0.889	0.0501	0.0055	211	25	209	4.8	200	240	104.4
7322/7-1_679-763	Zircon_sample-009.FIN2	337	99	49	14	365	91	7.60	5.863	0.091	0.332	0.018	0.848	0.1239	0.0038	1956	13	1850	87	2013	55	91.9
7322/7-1_679-763	Zircon_sample-010.FIN2	886	61	299	39	1920	170	3.04	3.840	0.140	0.271	0.008	0.929	0.1006	0.0013	1600	29	1546	40	1636	23	94.5
7322/7-1_679-763	Zircon_sample-012.FIN2	215	27	64	3	504	33	3.38	5.780	0.340	0.337	0.026	0.895	0.1201	0.0024	1942	50	1870	130	1957	36	95.6
7322/7-1_679-763	Zircon_sample-013.FIN2	400	100	61	16	670	180	6.48	13.690	0.430	0.500	0.012	0.660	0.1936	0.0044	2727	30	2636	63	2772	37	95.1
7322/7-1_679-763	Zircon_sample-014.FIN2	310	120	52	12	488	99	5.90	14.100	1.200	0.521	0.039	0.990	0.1944	0.0033	2754	77	2700	160	2779	28	97.2
Common-Pb corrected <sup>d</sup>																						
7322/7-1_679-763	Zircon_sample-011.FIN2	546	71	132	32	870	140	4.94	3.860	0.180	0.284	0.015	0.906	0.0968	0.0023	1604	35	1613	75	1561	44	103.3
SU2			CONCENTRATIONS <sup>a</sup>					RATIOS					AGES					Concordance				
Sample	Analysis	U [ppm]	2 σ	Th [ppm]	2 σ	Pb [ppm]	2 σ	U/Th <sup>b</sup>	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>206</sup> Pb/ <sup>238</sup> U <sup>b</sup>	2 σ <sup>d</sup>	rho <sup>c</sup>	<sup>207</sup> Pb/ <sup>206</sup> Pb <sup>b</sup>	2 σ <sup>d</sup>	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>206</sup> Pb/ <sup>238</sup> U <sup>b</sup>	2 σ <sup>d</sup>	<sup>207</sup> Pb/ <sup>206</sup> Pb <sup>b</sup>	2 σ <sup>d</sup>	
7321-41T2_1010-1020	Zircon_sample-007.FIN2	175	24	99	10	187	18	1.82	0.569	0.043	0.073	0.002	0.000	0.0563	0.0041	456	28	455	11	450	160	101.1
7321-41T2_1010-1020	Zircon_sample-008.FIN2	182	6	72	2	673	48	2.62	8.550	0.420	0.438	0.014	0.776	0.1422	0.0040	2290	45	2341	63	2254	49	103.9
7321-41T2_1010-1020	Zircon_sample-009.FIN2	472	70	30	3	116	11	16.60	1.661	0.669	0.168	0.008	0.856	0.0717	0.0017	993	26	1003	43	977	48	102.7
7321-41T2_1010-1020	Zircon_sample-013.FIN2	380	110	148	48	199	55	2.73	0.411	0.035	0.057	0.002	0.486	0.0522	0.0057	349	25	360	12	280	240	128.6
7321-41T2_1010-1020	Zircon_sample-025.FIN2	266	56	22	6	239	48	13.30	15.180	0.990	0.595	0.038	0.940	0.1852	0.0041	2824	62	3010	150	2699	37	111.5
7321-41T2_1010-1020	Zircon_sample-028.FIN2	1260	330	7	3	100	28	200.00	0.625	0.019	0.071	0.003	0.926	0.0641	0.0022	493	12	443	19	745	71	59.5
7321-41T2_1010-1020	Zircon_sample-036.FIN2	52	3	47	4	548	60	1.13	18.900	0.940	0.596	0.042	0.981	0.2305	0.0046	3036	48	3010	170	3055	32	98.5
7321-41T2_1010-1020	Zircon_sample-041.FIN2	195	30	41	9	267	68	4.82	4.179	0.087	0.294	0.009	0.498	0.1035	0.0030	1670	17	1663	42	1687	53	98.6
7321-41T2_1010-1020	Zircon_sample-050.FIN2	249	16	56	3	408	61	4.45	4.258	0.095	0.307	0.008	0.551	0.1000	0.0016	1685	18	1725	38	1623	29	106.3
7321-41T2_1010-1020	Zircon_sample-052.FIN2	301	16	128	9	203	21	2.34	0.633	0.096	0.067	0.003	0.096	0.0690	0.0100	496	61	416	16	860	340	48.4
7321-41T2_1010-1020	Zircon_sample-053.FIN2	132	9	96	4	1040	100	1.37	13.240	0.740	0.533	0.018	0.976	0.1801	0.0052	2696	54	2754	76	2653	48	103.8
7321-41T2_1010-1020	Zircon_sample-056.FIN2	480	100	42	7	262	49	11.60	3.830	0.130	0.268	0.010	0.867	0.1034	0.0021	1599	27	1532	50	1686	37	90.9
7321-41T2_1010-1020	Zircon_sample-064.FIN2	864	53	302	28	1760	130	2.84	3.370	0.140	0.266	0.011	0.907	0.0920	0.0017	1497	34	1518	57	1467	35	103.5
7321-41T2_1010-1020	Zircon_sample-066.FIN2	346	46	60	8	680	120	5.69	10.770	0.690	0.498	0.026	0.986	0.1570	0.0030	2502	60	2600	110	2424	32	107.3
7321-41T2_1010-1020	Zircon_sample-070.FIN2	530	190	56	10	239	51	8.80	1.733	0.071	0.173	0.007	0.815	0.0741	0.0017	1020	26	1029	40	1043	47	98.7
7321-41T2_1010-1020	Zircon_sample-083.FIN2	207	18	70	12	430	35	2.95	3.890	0.340	0.279	0.011	0.016	0.1009	0.0087	1609	69	1587	55	1630	160	97.4
7321-41T2_1010-1020	Zircon_sample-093.FIN2	111	24	38	5	397	30	2.96	13.580	0.730	0.546	0.009	0.874	0.1806	0.0079	2720	51	2808	36	2657	72	105.7
7321-41T2_1010-1020	Zircon_sample-094.FIN2	151	32	89	26	600	150	1.84	3.720	0.120	0.278	0.012	0.558	0.0979	0.0053	1576	26	1579	61	1580	100	99.9
7321-41T2_1010-1020	Zircon_sample-099.FIN2	114	7	110	9	464	38	1.03	2.010	0.150	0.190	0.005	0.690	0.0773	0.0076	1118	52	1120	27	1120	200	100
7321-41T2_1010-1020	Zircon_sample-098.FIN2	388	99	75	19	306	53	5.15	1.889	0.093	0.172	0.004	0.166	0.0780	0.0031	1076	33	1025	24	1145	79	89.5
7321-41T2_1010-1020	Zircon_sample-108.FIN2	123	26	44	5	625	81	2.74	26.140	0.690	0.587	0.027	0.854	0.3224	0.0071	3351	26	2980	110	3581	34	83.2
7321-41T2_1010-1020	Zircon_sample-109.FIN2	508	39	73	21	530	140	8.00	4.120	0.220	0.292	0.011	0.612	0.1027	0.0044	1658	43	1652	53	1671	78	98.9
7321-41T2_1010-1020	Zircon_sample-112.FIN2	320	110	137	44	1010	200	2.29	4.060	0.160	0.299	0.013	0.546	0.0976	0.0027	1646	31	1684	65	1576	52	106.9
7321-41T2_1010-1020	Zircon_sample-120.FIN2	103	13	29	5	147	27	3.52	2.150	0.250	0.190	0.002	0.131	0.0820	0.0100	1162	81	1121	11	1240	250	90.4
7321-41T2_1010-1020	Zircon_sample-123.FIN2	127	6	60	3	541	33	2.01	6.050	0.260	0.376	0.010	0.623	0.1172	0.0056	1952	37	2056	59	1911	87	107.6
7321-41T2_1010-1020	Zircon_sample-125.FIN2	91	7	31	4	237	26	2.91	4.450	0.240	0.305	0.015	0.674	0.1063	0.0047	1719	45	1713	72	1733	80	98.8
7321-41T2_1010-1020	Zircon_sample-135.FIN2	820	120	180	45	1110	210	4.66	3.540	0.140	0.274	0.010	0.827	0.0940	0.0023	1536	33	1559	50	1507	47	103.5
7321-41T2_1010-1020	Zircon_sample-136.FIN2	560	160	85	19	414	83	6.28	2.228	0.091	0.208	0.006	0.647	0.0765	0.0013	1189	29	1216	31	1107	35	109.8
7321-41T2_1010-1020	Zircon_sample-138.FIN2	272	58	141	50	312	75	2.07	0.790	0.048	0.094	0.002	0.222	0.0626	0.0031	590	27	576	13	690	100	83.5
7321-41T2_1010-1020	Zircon_sample-139.FIN2	233	22	43	5	554	53	5.38	15.860	0.610	0.534	0.012	0.804	0.2154	0.0057	2868	37	2759	52	2946	43	93.7
7321-41T2_1010-1020	Zircon_sample-149.FIN2	205	32	65	20	399	89	3.37	3.620	0.350	0.282	0.010	0.019	0.0928	0.0089	1551	78	1600	48	1470	190	108.8
7321-41T2_1010-1020	Zircon_sample-153.FIN2	91	16	20	3	78	6	4.63	2.230	0.120	0.197	0.005	0.406	0.0824	0.0039	1188	38	1157	26	1252	92	92.4
7321-41T2_1010-1020	Zircon_sample-154.FIN2	341	20	10	5	58	24	42.00	2.020	0.120	0.186	0.003	0.961	0.0792	0.0042	1123	43	1102	17	1180	100	93.4
7321-41T2_1010-1020	Zircon_sample-163.FIN2	127	10	45	3	526	25	5.11	14.510	0.110	0.569	0.015	0.810	0.1881	0.0047	2784	74	2903	61	2725	41	106.5
7321-41T2_1010-1020	Zircon_sample-165.FIN2	221	35	24	16	108	74	6.80	1.630	0.180	0.161	0.004	0.175	0.0736	0.0091	978	68	960	20	1020	250	94.1
7321-41T2_1010-1020	Zircon_sample-176.FIN2	233	37	98	10	1121	62	2.37	13													



# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

Fig. 7.3

Frøholmen FM		CONCENTRATIONS <sup>a</sup>					RATIOS							AGES					Concordance							
Sample	Analysis	U [ppm]	Th [ppm]	2 $\sigma$	Pb [ppm]	2 $\sigma$	U/Th <sup>b</sup>	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 $\sigma$	<sup>206</sup> Pb/ <sup>238</sup> Pb <sup>b</sup>	2 $\sigma$	rho <sup>c</sup>	<sup>207</sup> Pb/ <sup>206</sup> Pb <sup>b</sup>	2 $\sigma$	<sup>207</sup> Pb/ <sup>235</sup> U <sup>b</sup>	2 $\sigma$	<sup>206</sup> Pb/ <sup>238</sup> Pb <sup>b</sup>	2 $\sigma$	AGES	2 $\sigma$	AGES	2 $\sigma$	AGES	2 $\sigma$	AGES	Concordance
7321-4-172_1422-1428	Zircon_sample-023.FIN2	1287	81	237	97		330	150	5.90	0.517	0.021	0.067	0.002	0.304	0.0545	0.0027	423	14	417	14	390	110	106.9			
7321-4-172_1422-1428	Zircon_sample-026.FIN2	1530	850	119	62		480	270	11.70	4.860	0.930	0.306	0.040	0.991	0.1143	0.0016	1790	110	1720	200	1868	25	92.1			
7321-4-172_1422-1428	Zircon_sample-028.FIN2	334	63	146	22		138	29	2.15	0.290	0.030	0.039	0.002	0.031	0.0546	0.0069	258	24	244	12	370	280	65.9			
7321-4-172_1422-1428	Zircon_sample-036.FIN2	345	61	110	14		714	83	3.13	4.720	0.220	0.307	0.007	0.906	0.1095	0.0036	1770	40	1727	47	1790	60	96.5			
7321-4-172_1422-1428	Zircon_sample-040.FIN2	405	59	138	35		670	120	3.00	2.922	0.091	0.218	0.003	0.598	0.0977	0.0019	1388	24	1268	17	1581	37	80.2			
7321-4-172_1422-1428	Zircon_sample-041.FIN2	350	120	106	33		126	38	3.37	0.452	0.070	0.046	0.002	0.049	0.0711	0.0091	377	50	290	13	930	270	31.2			
7321-4-172_1422-1428	Zircon_sample-042.FIN2	72	3	39	3		230	19	1.87	4.873	0.096	0.302	0.006	0.504	0.1150	0.0033	1797	17	1699	28	1879	52	90.4			
7321-4-172_1422-1428	Zircon_sample-050.FIN2	730	480	126	67		1100	1100	5.50	11.000	0.410	0.458	0.016	-0.607	0.1740	0.0110	2531	34	2428	69	2590	100	93.7			
7321-4-172_1422-1428	Zircon_sample-051.FIN2	472	74	93	25		118	45	5.24	0.566	0.088	0.068	0.003	0.842	0.0601	0.0077	455	57	426	18	600	270	71.0			
7321-4-172_1422-1428	Zircon_sample-053.FIN2	361	37	96	17		176	23	3.86	0.587	0.021	0.072	0.001	0.163	0.0596	0.0022	469	13	447	7.3	586	83	76.3			
7321-4-172_1422-1428	Zircon_sample-055.FIN2	610	140	217	45		1073	91	2.84	5.030	0.100	0.299	0.010	0.006	0.1201	0.0042	1824	17	1684	49	1958	61	86.0			
7321-4-172_1422-1428	Zircon_sample-056.FIN2	240	120	88	21		410	130	2.71	3.480	0.650	0.240	0.010	0.743	0.1040	0.0180	1520	140	1386	53	1670	310	83.0			
7321-4-172_1422-1428	Zircon_sample-064.FIN2	630	140	259	28		262	34	2.45	0.309	0.019	0.042	0.001	0.355	0.0529	0.0029	274	15	266	6.6	330	120	83.0			
7321-4-172_1422-1428	Zircon_sample-066.FIN2	292	32	85	6		765	20	3.48	10.140	0.290	0.452	0.017	0.975	0.1634	0.0016	2447	36	2405	76	2491	17	95.5			
7321-4-172_1422-1428	Zircon_sample-068.FIN2	800	210	127	17		740	300	6.40	4.950	0.300	0.294	0.017	0.968	0.1233	0.0009	1811	51	1660	85	2004	13	82.8			
7321-4-172_1422-1428	Zircon_sample-069.FIN2	387	34	77	4		102	15	5.16	0.427	0.033	0.057	0.001	0.246	0.0549	0.0044	361	23	357	7	370	170	91.4			
7321-4-172_1422-1428	Zircon_sample-070.FIN2	447	20	166	7		137	35	2.75	0.322	0.009	0.037	0.002	0.123	0.0700	0.0140	284	7	233	13	870	350	26.8			
7321-4-172_1422-1428	Zircon_sample-078.FIN2	460	110	157	77		153	72	3.32	0.341	0.020	0.042	0.002	-0.642	0.0589	0.0051	297	15	266	10	550	190	48.4			
7321-4-172_1422-1428	Zircon_sample-079.FIN2	918	80	287	25		334	38	3.25	0.350	0.023	0.048	0.002	0.273	0.0523	0.0031	305	17	303	11	290	130	104.5			
7321-4-172_1422-1428	Zircon_sample-081.FIN2	510	240	250	130		190	100	2.07	0.236	0.019	0.034	0.002	-0.592	0.0503	0.0056	215	15	216	9.8	200	230	108.0			
7321-4-172_1422-1428	Zircon_sample-082.FIN2	610	100	67	23		480	140	9.00	5.080	0.210	0.325	0.015	0.909	0.1142	0.0017	1833	34	1811	71	1866	27	97.1			
7321-4-172_1422-1428	Zircon_sample-092.FIN2	325	64	27	7		263	20	14.60	11.330	0.430	0.475	0.018	0.872	0.1704	0.0010	2550	36	2506	78	2561.4	9.3	97.8			
7321-4-172_1422-1428	Zircon_sample-094.FIN2	190	5	94	5		192	15	1.99	0.980	0.068	0.107	0.004	0.907	0.0668	0.0069	693	35	655	23	820	220	75.9			
7321-4-172_1422-1428	Zircon_sample-095.FIN2	1155	94	447	44		1860	130	2.55	1.824	0.999	0.146	0.000	0.473	0.0884	0.0030	1054	35	881	2.3	1390	64	63.4			
7321-4-172_1422-1428	Zircon_sample-096.FIN2	516	76	116	10		137	20	4.41	0.358	0.028	0.050	0.002	0.659	0.0520	0.0024	310	21	313	11	280	110	111.8			
7321-4-172_1422-1428	Zircon_sample-097.FIN2	147	32	42	9		189	44	3.45	2.230	0.150	0.199	0.004	0.553	0.0817	0.0062	1190	46	1167	21	1240	150	94.1			
7321-4-172_1422-1428	Zircon_sample-107.FIN2	213	14	56	4		314	32	3.74	2.806	0.080	0.229	0.007	0.381	0.0889	0.0038	1357	21	1328	35	1397	81	95.1			
7321-4-172_1422-1428	Zircon_sample-108.FIN2	852	66	37	8		321	95	23.60	5.250	0.160	0.325	0.006	0.960	0.1174	0.0019	1860	26	1812	27	1917	28	94.5			
7321-4-172_1422-1428	Zircon_sample-109.FIN2	1700	100	45	23		200	99	42.00	2.163	0.091	0.180	0.002	0.975	0.0701	0.0036	1169	29	1069	9.6	1362	80	78.5			
7321-4-172_1422-1428	Zircon_sample-110.FIN2	2050	430	138	73		1050	610	19.20	5.480	0.470	0.318	0.017	0.922	0.1236	0.0037	1895	73	1777	62	2008	54	88.5			
7321-4-172_1422-1428	Zircon_sample-111.FIN2	616	59	216	43		2370	200	2.76	11.800	0.430	0.476	0.014	0.949	0.0109	0.0009	2537	25	2510	62	2566.4	9.3	97.8			
7321-4-172_1422-1428	Zircon_sample-112.FIN2	830	130	330	100		3400	780	2.64	10.180	0.400	0.441	0.020	0.811	0.1665	0.0025	2451	36	2357	88	2523	25	93.4			
7321-4-172_1422-1428	Zircon_sample-121.FIN2	680	20	132	15		757	58	3.780	1.310	0.273	0.004	-0.264	0.0983	0.0049	1588	28	1554	18	1915	94	97.7				
7321-4-172_1422-1428	Zircon_sample-122.FIN2	1390	250	240	130		1740	800	13.00	5.690	0.390	0.341	0.015	0.763	0.1207	0.0041	1927	58	1890	70	1964	61	96.2			
7321-4-172_1422-1428	Zircon_sample-123.FIN2	1170	74	408	43		660	83	2.90	0.523	0.014	0.067	0.001	0.871	0.0564	0.0017	427	9.6	419	7.7	465	65	90.0			
7321-4-172_1422-1428	Zircon_sample-124.FIN2	618	82	293	45		2670	270	2.06	10.130	0.420	0.440	0.018	0.773	0.1656	0.0032	2446	38	2350	80	2513	32	93.5			
7321-4-172_1422-1428	Zircon_sample-125.FIN2	381	31	82	6		564	22	4.67	4.945	0.305	0.310	0.005	0.847	0.1158	0.0020	1810	59	1742	25	1892	31	92.1			
7321-4-172_1422-1428	Zircon_sample-134.FIN2	257	67	50	17		99	22	5.23	0.709	0.038	0.079	0.004	0.292	0.0661	0.0060	544	23	488	27	790	180	61.8			
7321-4-172_1422-1428	Zircon_sample-135.FIN2	746	72	312	33		324	27	2.36	0.315	0.031	0.040	0.001	0.904	0.0575	0.0056	278	24	252	7.9	490	230	51.4			
7321-4-172_1422-1428	Zircon_sample-137.FIN2	54	195	20	13		170	19	2.07	0.279	0.014	0.036	0.007	0.508	0.0509	0.0062	520	79	428	19	520	190	51.8			
7321-4-172_1422-1428	Zircon_sample-138.FIN2	574	59	183	16		1670	110	3.12	10.480	0.690	0.454	0.028	0.970	0.1650	0.0020	2476	61	2410	120	2508	20	93.4			
7321-4-172_1422-1428	Zircon_sample-139.FIN2	584	52	222	40		1600	280	2.64	5.810	0.210	0.338	0.008	0.819	0.1236	0.0024	1947	31	1878	38	2009	33	93.5			
7321-4-172_1422-1428	Zircon_sample-139.FIN2	1580	360	64	12		420	41	25.40	4.590	0.410	0.299	0.026	0.967	0.1116	0.0014	1745	73	1680	130	1826	23	92.0			
7321-4-172_1422-1428	Zircon_sample-149.FIN2	795	59	289	34		60	7	2.65	0.514	0.033	0.067	0.001	0.525	0.0552	0.0025	421	22	418	7.7	420	100	99.4			
7321-4-172_1422-1428	Zircon_sample-150.FIN2	694	44	398	31		66	7	1.70	0.442	0.026	0.056	0.002	0.291	0.0572	0.0033	371	18	353	9	490	130	72.1			
7321-4-172_1422-1428	Zircon_sample-151.FIN2	300	150	270	170		64	27	1.18	0.655	0.063	0.080	0.005	0.462	0.0600	0.0050	510	39	493	32	590	180	83.6			
7321-4-172_1422-1428	Zircon_sample-152.FIN2	1148	57	607	57		83	5	1.90	0.358	0.023	0.045	0.002	0.223	0.0578	0.0025	1392	17	1284	12	1517	96	54.9			
7321-4-172_1422-1428	Zircon_sample-153.FIN2	325	42	98	5		70	8	3.30	2.940	0.028	0.237	0.005	0.537	0.0893	0.0025	3199	99	1973	28	1409	54	97.4			
7321-4-172_1422-1428	Zircon_sample-155.FIN2	21	21	27	37		180	19	2.60	4.250																

# Provenance Evaluation of Lower Cretaceous in the Stappen High Area and Implications for Reservoir Development

Fig. 7.4

Snaad FM		CONCENTRATIONS <sup>a</sup>										RATIOS										AGES										Concordance
		U (ppm)		Th (ppm)		Pb (ppm)		2 <sup>35</sup> U/2 <sup>38</sup> U		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb		2 <sup>35</sup> Pb/2 <sup>38</sup> Pb								
7321-4 1T2_1521-1524	Zircon_sample-260.FIN2	690	290	163	66	280	130	3.50	0.391	0.024	0.056	0.005	0.814	0.0533	0.0033	349	30	351	29	340	140	103.2										
7321-4 1T2_1521-1524	Zircon_sample-261.FIN2	249	7	42	2	376	35	5.76	6.980	0.440	0.383	0.019	0.558	0.1326	0.0024	2108	56	2091	90	2133	32	98.0										
7321-4 1T2_1521-1524	Zircon_sample-265.FIN2	216	44	64	18	105	27	3.39	0.650	0.150	0.075	0.002	0.975	0.0625	0.0092	507	89	463	32	680	310	68.1										
7321-4 1T2_1521-1524	Zircon_sample-266.FIN2	409	53	34	6	307	43	15.10	5.360	0.200	0.348	0.013	0.777	0.1109	0.0021	1878	32	1923	61	1828	24	105.2										
7321-4 1T2_1521-1524	Zircon_sample-276.FIN2	336	59	199	55	221	44	1.78	0.480	0.100	0.055	0.005	0.701	0.0640	0.0120	397	68	347	32	660	410	52.6										
7321-4 1T2_1521-1524	Zircon_sample-289.FIN2	450	61	149	20	376	87	2.98	0.964	0.068	0.105	0.002	0.221	0.0665	0.0042	685	35	643	12	820	130	78.4										
7321-4 1T2_1521-1524	Zircon_sample-290.FIN2	200	18	39	6	97	23	5.15	0.777	0.067	0.091	0.001	0.360	0.0623	0.0055	583	38	561	4.6	670	200	83.7										
7321-4 1T2_1521-1524	Zircon_sample-291.FIN2	759	64	249	30	323	42	3.02	0.394	0.036	0.050	0.001	0.392	0.0573	0.0048	337	27	315	6.2	490	190	64.2										
7321-4 1T2_1521-1524	Zircon_sample-293.FIN2	640	110	154	39	358	94	0.22	0.869	0.022	0.101	0.002	0.624	0.0624	0.0013	635	12	621	12	688	45	90.3										
7321-4 1T2_1521-1524	Zircon_sample-294.FIN2	383	60	172	22	250	48	2.17	0.400	0.300	0.052	0.002	0.625	0.0558	0.0039	342	22	325	13	440	160	73.9										
7321-4 1T2_1521-1524	Zircon_sample-302.FIN2	1690	220	185	31	1440	220	9.16	5.500	0.180	0.345	0.010	0.938	0.1180	0.0022	1900	29	1909	50	1926	34	99.1										
7321-4 1T2_1521-1524	Zircon_sample-303.FIN2	1150	190	137	42	180	50	8.60	0.429	0.042	0.058	0.001	0.231	0.0540	0.0056	362	30	362	5.9	360	240	100.6										
7321-4 1T2_1521-1524	Zircon_sample-304.FIN2	230	78	52	35	81	46	5.00	0.740	0.028	0.064	0.002	0.704	0.0533	0.0022	391	19	400	13	340	91	117.6										
7321-4 1T2_1521-1524	Zircon_sample-305.FIN2	221	40	103	31	213	35	2.25	0.596	0.093	0.068	0.004	0.619	0.0637	0.0076	471	57	421	23	700	250	60.1										
7321-4 1T2_1521-1524	Zircon_sample-308.FIN2	449	45	167	20	277	35	2.64	0.492	0.028	0.065	0.003	0.889	0.0552	0.0027	406	19	405	16	420	110	96.4										
7321-4 1T2_1521-1524	Zircon_sample-316.FIN2	640	150	147	49	219	41	4.48	0.455	0.047	0.057	0.002	0.824	0.0580	0.0046	380	33	357	12	520	170	68.7										
7321-4 1T2_1521-1524	Zircon_sample-317.FIN2	197	22	70	22	400	100	3.10	2.150	0.150	0.202	0.002	0.159	0.0769	0.0060	1163	50	1187	9.4	1110	160	106.9										
7321-4 1T2_1521-1524	Zircon_sample-318.FIN2	118	29	28	11	33	9	4.72	0.454	0.057	0.058	0.002	0.281	0.0552	0.0060	379	39	365	11	400	240	91.3										
7321-4 1T2_1521-1524	Zircon_sample-319.FIN2	118	29	28	11	33	9	3.48	0.499	0.057	0.063	0.003	0.563	0.0055	0.0055	405	39	392	17	430	220	91.2										
7321-4 1T2_1521-1524	Zircon_sample-320.FIN2	392	19	170	0	202	26	2.29	0.399	0.018	0.044	0.000	0.846	0.0511	0.0029	273	14	275	15	240	130	114.8										
7321-4 1T2_1521-1524	Zircon_sample-322.FIN2	164	13	47	3	83	14	3.56	0.612	0.067	0.071	0.002	0.254	0.0622	0.0068	483	42	445	13	640	240	69.5										
7321-4 1T2_1521-1524	Zircon_sample-323.FIN2	274	72	172	59	360	120	1.62	0.740	0.210	0.081	0.002	0.725	0.0660	0.0190	560	120	504	11	770	550	65.5										
7321-4 1T2_1521-1524	Zircon_sample-333.FIN2	286	62	102	26	440	120	5.30	3.650	0.280	0.262	0.014	0.972	0.1010	0.0029	1557	62	1497	72	1641	54	91.2										
7321-4 1T2_1521-1524	Zircon_sample-334.FIN2	510	120	102	30	570	160	5.08	3.187	0.073	0.256	0.004	0.253	0.0991	0.0026	1454	18	1471	21	1427	55	103.1										
7321-4 1T2_1521-1524	Zircon_sample-335.FIN2	313	36	85	10	218	37	3.72	0.953	0.081	0.104	0.001	0.113	0.0658	0.0055	678	42	640	7.7	810	190	82.0										
7321-4 1T2_1521-1524	Zircon_sample-336.FIN2	425	42	130	14	160	15	3.31	0.445	0.043	0.049	0.001	0.769	0.0664	0.0061	374	31	306	3.7	810	190	117.6										
7321-4 1T2_1521-1524	Zircon_sample-344.FIN2	1080	210	550	130	930	170	2.00	0.455	0.018	0.062	0.001	0.132	0.0531	0.0023	381	13	390	7.2	329	99	118.6										
7321-4 1T2_1521-1524	Zircon_sample-345.FIN2	191	16	66	8	174	32	3.06	0.834	0.072	0.100	0.002	0.100	0.0607	0.0052	614	39	614	11	610	180	100.7										
7321-4 1T2_1521-1524	Zircon_sample-349.FIN2	1510	230	274	46	270	74	5.59	0.341	0.031	0.045	0.000	0.106	0.0550	0.0045	298	23	282	2	740	180	68.7										
7321-4 1T2_1521-1524	Zircon_sample-358.FIN2	390	49	241	70	550	110	1.73	0.846	0.054	0.096	0.004	0.570	0.0643	0.0053	622	30	591	21	740	180	79.9										
7321-4 1T2_1521-1524	Zircon_sample-360.FIN2	145	21	80	11	220	34	1.85	0.840	0.300	0.077	0.001	0.376	0.0800	0.0270	610	170	475	5.1	1110	760	42.8										
7321-4 1T2_1521-1524	Zircon_sample-361.FIN2	1030	140	340	110	309	72	3.24	0.268	0.015	0.037	0.000	0.190	0.0529	0.0024	241	12	233	2.7	320	100	72.7										
7321-4 1T2_1521-1524	Zircon_sample-362.FIN2	333	50	79	26	342	97	5.20	1.918	0.062	0.184	0.004	0.315	0.0757	0.0026	1087	22	1088	24	1082	71	100.6										
7321-4 1T2_1521-1524	Zircon_sample-363.FIN2	241	21	114	6	690	110	2.14	2.930	0.130	0.249	0.006	0.924	0.0848	0.0029	1387	34	1435	30	1311	116	109.5										
7321-4 1T2_1521-1524	Zircon_sample-364.FIN2	1870	160	300	9	930	110	6.31	1.185	0.046	0.121	0.003	0.803	0.0707	0.0021	793	21	738	15	948	60	77.8										
7321-4 1T2_1521-1524	Zircon_sample-372.FIN2	880	240	429	27	680	160	2.47	0.590	0.032	0.069	0.002	0.291	0.0625	0.0047	470	20	429	14	680	160	63.1										
7321-4 1T2_1521-1524	Zircon_sample-373.FIN2	1260	120	426	79	500	100	3.19	0.316	0.100	0.044	0.001	0.676	0.0521	0.0014	279	7.7	279	4.1	288	60	96.8										
7321-4 1T2_1521-1524	Zircon_sample-375.FIN2	159	18	34	4	219	33	4.71	3.600	0.110	0.280	0.005	0.886	0.0941	0.0011	1559	31	1592	27	1509	23	105.5										
7321-4 1T2_1521-1524	Zircon_sample-376.FIN2	668	41	72	29	440	140	12.20	3.732	0.097	0.282	0.009	0.923	0.0954	0.0012	1578	21	1601	45	1536	24	104.2										
7321-4 1T2_1521-1524	Zircon_sample-377.FIN2	224	44	165	50	205	47	1.41	0.398	0.031	0.049	0.002	0.387	0.0589	0.0061	340	23	311	15	550	230	56.5										
7321-4 1T2_1521-1524	Zircon_sample-378.FIN2	388	48	79	10	103	12	5.00	0.383	0.039	0.053	0.002	0.524	0.0556	0.0064	343	38	331	14	400	210	84.4										
7321-4 1T2_1521-1524	Zircon_sample-386.FIN2	169	17	61	6	150	10	2.81	0.840	0.036	0.096	0.004	0.022	0.0631	0.0032	619	20	591	21	700	140	80.6										
7321-4 1T2_1521-1524	Zircon_sample-387.FIN2	442	57	188	42	1330	267	4.49	0.665	0.309	0.005	0.290	0.1058	0.0019	1731	12	1738	23	1738	33	100.6											
7321-4 1T2_1521-1524	Zircon_sample-390.FIN2	282	28	70	7	110	8	4.12	0.544	0.065	0.065	0.004	0.778	0.0609	0.0057	440	44	404	24	620	210	65.2										
7321-4 1T2_1521-1524	Zircon_sample-391.FIN2	248	40	139	32	348	72	1.87	0.823	0.049	0.099	0.003	0.261	0.0607	0.0052	609	27	608	18	610	190	99.7										
7321-4 1T2_1521-1524	Zircon_sample-402.FIN2	236	11	55	6	75	7	4.44	0.369	0.054	0.053	0.002	0.162	0.0503	0.0076	318	40	335	13	190	320	176.3										
7321-4 1T2_1521-1524	Zircon_sample-406.FIN2	188	7	67	6	290	62	2.89	1.510	0.140	0.138	0.005	0.828	0.0792	0.0049	933	58	835	28	1170	120	71.4										
7321-4 1T2_1521-1524	Zircon_sample-414.FIN2	900	140	178	24	352	44	5.30	0.612	0.030	0.075	0.002	0.694	0.0593	0.0021	491	15	465	11	572	80	81.3										
7321-4 1T2_1521-1524	Zircon_sample-415.FIN2	335	18	88	10	439	54	3.84	2.240	0.190	0.212	0.002	0.656	0.0766	0.0064	1193	59	1242	8.1	1110	160	111.0										
7321-4 1T2_1521-1524	Zircon_sample-416.FIN2	616	64	27	6	111	28	24.70	0.974	0.044	0.117	0.002	0.624	0.0619	0.0030	700	28	715	11	660	100	108.3										
7321-4 1T2_1521-1524	Zircon_sample-417.FIN2	1713	76	103	4	939	49	16.73	5.350	0.071	0.344	0.004	0.710	0.1123	0.0021	1877	11	1904	17	1837	34	103.6										
7321-4 1T2_1521-1524	Zircon_sample-418.FIN2	694	26	425	43	465	26	1.62	0.311	0.015	0.044	0.001	0.937	0.0517	0.0021	275	12	277	4	266	96	104.1										
7321-4 1T2_1521-1524	Zircon_sample-419.FIN2	1710	120	502	54	458	53	3.																								

## 8 REFERENCES

- Anell, I., Faleide, J.-I., & Braathen, A. (2016). Regional tectono-sedimentary development of the highs and basins of the northwestern Barents Shelf. *Norwegian Journal of Geology*, 96 (1), 27-41. <https://doi.org/10.17850/njg96-1-04>
- Anfinson, O. A., Leier, A. L., Embry, A. F., & Dewing, K. (2012). Detrital zircon geochronology and provenance of the Neoproterozoic to Late Devonian Franklinian Basin, Canadian Arctic Islands. *Geological Society of America Bulletin*, 124(3–4), 415–430. <https://doi.org/10.1130/B30503.1>
- Blaich, O. A., Tsikalas, F., & Faleide, J. I. (2017). New insights into the tectono-stratigraphic evolution of the southern Stappen High and its transition to Bjørnøya Basin, SW Barents Sea. *Marine and Petroleum Geology*, 85, 89–105. <https://doi.org/10.1016/j.marpetgeo.2017.04.015>
- Bryn, B. K. L., Ahokas, J., Patruno, S., Schjelderup, S., Hinna, C., Lowrey, C., & Escalona, A. (2019). Exploring the reservoir potential of Lower Cretaceous Clinofolds in the Fingerdjupet Subbasin, Norwegian Barents Sea. *Basin Research*, 32(2), 332–347. <https://doi.org/10.1111/bre.12407>
- Bue, E.P. & Andresen, A. (2013). Constraining depositional models in the Barents Sea region using detrital zircon U–Pb data from Mesozoic sediments in Svalbard. *Geological Society, London, Special Publications*, 386(1), 261–279. <https://doi.org/10.1144/SP386.14>
- Corfu, F., Andersen, T. B., & Gasser, D. (2014). The Scandinavian Caledonides: Main features, conceptual advances and critical questions. *Geological Society, London, Special Publications*, 390(1), 9–43. <https://doi.org/10.1144/SP390.25>
- Corfu, F., Polteau, S., Planke, S., Faleide, J. I., Svensen, H., Zayoncheck, A., & Stolbov, N. (2013). U–Pb geochronology of Cretaceous magmatism on Svalbard and Franz Josef Land, Barents Sea Large Igneous Province. *Geological Magazine*, 150(6), 1127–1135. <https://doi.org/10.1017/S0016756813000162>
- Dalland, A., Worsley, D., Ofstad, K., (1988). A lithostratigraphic scheme for the mesozoic and cenozoic succession offshore Norway north of 62 N. *NPD Bulletin*. 4,67.

Dewing, K., Mayr, U., Harrison, J.C., and de Freitas, T. (2008). Upper Neoproterozoic to Lower Devonian stratigraphy of northeast Ellesmere Island. *In* *Geology of northeast Ellesmere Island adjacent to Kane Basin and Kennedy Channel, Nunavut*. Edited by U. Mayr. Geological Survey of Canada, Bulletin 592, 31–108

Doré, A. G. (1991). The structural foundation and evolution of Mesozoic seaways between Europe and the Arctic. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 87(1–4), 441–492. [https://doi.org/10.1016/0031-0182\(91\)90144-G](https://doi.org/10.1016/0031-0182(91)90144-G)

Faleide, J. I., Vågnes, E., & Gudlaugsson, S. T. (1993). Late Mesozoic–Cenozoic evolution of the southwestern Barents Sea. *Geological Society, London, Petroleum Geology Conference Series*, 4(1), 933–950. <https://doi.org/10.1144/0040933>

Faleide, J. I., Bjørlykke, K., & Gabrielsen, R. H. (2010). Geology of the Norwegian Continental Shelf. *In* K. Bjørlykke, *Petroleum Geoscience* (s. 467–499). Springer Berlin Heidelberg. [https://doi.org/10.1007/978-3-642-02332-3\\_22](https://doi.org/10.1007/978-3-642-02332-3_22)

Faleide, J. I., Bjørlykke, K., & Gabrielsen, R. H. (2015). Geology of the Norwegian Continental Shelf. *In* K. Bjørlykke (Red.), *Petroleum Geoscience* (s. 603–637). Springer Berlin Heidelberg. [https://doi.org/10.1007/978-3-642-34132-8\\_25](https://doi.org/10.1007/978-3-642-34132-8_25)

Faleide, T. S., Midtkandal, I., Planke, S., Corseri, R., Faleide, J. I., Serck, C. S., & Nystuen, J. P. (2019). Characterisation and development of Early Cretaceous shelf platform deposition and faulting in the Hoop area, southwestern Barents Sea—Constrained by high-resolution seismic data. *Norwegian Journal of Geology*, 99(3), 1–20. <https://doi.org/10.17850/njg99-3-7>

Gabrielsen, R. H. (1990). *Structural elements of the Norwegian continental shelf. Part 1: The Barents Sea Region*. Oljedirektoratet.

Gasser, D. (2014). The Caledonides of Greenland, Svalbard and other Arctic areas: Status of research and open questions. *Geological Society, London, Special Publications*, 390(1), 93–129. <https://doi.org/10.1144/SP390.17>

Gasser, D., & Andresen, A. (2013). Caledonian terrane amalgamation of Svalbard: Detrital zircon provenance of Mesoproterozoic to Carboniferous strata from Oscar II Land, western Spitsbergen. *Geological Magazine*, 150(6), 1103–1126. <https://doi.org/10.1017/S0016756813000174>

Gee, D. (2015). Caledonides of Scandinavia, Greenland, and Svalbard. In *Reference Module in Earth Systems and Environmental Sciences*. Elsevier. <https://doi.org/10.1016/B978-0-12-409548-9.09133-8>

Gee, D., Fossen, H., Henriksen, N., & Higgins, A. K. (2008). From the Early Paleozoic Platforms of Baltica and Laurentia to the Caledonide Orogen of Scandinavia and Greenland. *Episodes*, 31(1), 44–51. <https://doi.org/10.18814/epiiugs/2008/v31i1/007>

Gee, D., & Teben'kov, A. M. (2004). Svalbard: A fragment of the Laurentian margin. *Geological Society, London, Memoirs*, 30(1), 191–206. <https://doi.org/10.1144/GSL.MEM.2004.030.01.16>

Gernigon, L., & Brönnert, M. (2012). Late Palaeozoic architecture and evolution of the southwestern Barents Sea: Insights from a new generation of aeromagnetic data. *Journal of the Geological Society*, 169(4), 449–459. <https://doi.org/10.1144/0016-76492011-131>

Gjelberg, J., & Steel, R. J. (1995). Helvetiafjellet formation (Barremian-Aptian), Spitsbergen: Characteristics of a transgressive succession. *Norwegian Petroleum Society Special Publications*, pp.571–593. Elsevier. [https://doi.org/10.1016/S0928-8937\(06\)80087-1](https://doi.org/10.1016/S0928-8937(06)80087-1)

Gjelberg, J. & Steel, R. (2012). Depositional model for the lower Cretaceous Helvetiafjellet Formation on Svalbard, diachronous vs layer-cake models. *Norwegian Journal of Geology*, 92, 41-54

Glørstad-Clark, E. (2011). Basin analysis in western Barents Sea area: The interplay between accommodation space and depositional system (PhD thesis). University of Oslo, Oslo.

Gradstein, F. M., Anthonissen, E., Brunstad, H., Charnock, M., Hammer, O., Hellem, T., & Lervik, K. S. (2010). Norwegian Offshore Stratigraphic Lexicon (NORLEX). *Newsletters on Stratigraphy*, 44(1), 73–86. <https://doi.org/10.1127/0078-0421/2010/0005>

Grantz, A., Hart, P. E., & Childers, V. A. (2011). Chapter 50 Geology and tectonic development of the Amerasia and Canada Basins, Arctic Ocean. *Geological Society, London, Memoirs*, 35(1), 771–799. <https://doi.org/10.1144/M35.50>

Grundvåg, S.-A., Marin, D., Kairanov, B., Śliwińska, K. K., Nøhr-Hansen, H., Jelby, M. E., ... Olaussen, S. (2017). The Lower Cretaceous succession of the northwestern Barents Shelf: Onshore and offshore correlations. *Marine and Petroleum Geology*, 86, 834–857. <https://doi.org/10.1016/j.marpetgeo.2017.06.036>

Gudlaugsson, S. T., Faleide, J. I., Johansen, S. E., & Breivik, A. J. (1998). Late Palaeozoic structural development of the South-western Barents Sea. *Marine and Petroleum Geology*, 15(1), 73–102. [https://doi.org/10.1016/S0264-8172\(97\)00048-2](https://doi.org/10.1016/S0264-8172(97)00048-2)

Hadlari, T., Davis, W. J., Dewing, K., Heaman, L. M., Lemieux, Y., Ootes, L., ... Pyle, L. J. (2012). Two detrital zircon signatures for the Cambrian passive margin of northern Laurentia highlighted by new U-Pb results from northern Canada. *Geological Society of America Bulletin*, 124(7–8), 1155–1168. <https://doi.org/10.1130/B30530.1>

Hellman, F. J., Gee, D. G., & Witt-Nilsson, P. (2001). Late Archean basement in the Bangenhuken Complex of the Nordbreen Nappe, western Ny-Friesland, Svalbard. *Polar Research*, 20(1), 49–59. <https://doi.org/10.3402/polar.v20i1.6499>

Henriksen, E., Ryseth, A. E., Larssen, G. B., Heide, T., Rønning, K., Sollid, K., & Stoupakova, A. V. (2011). Chapter 10 Tectonostratigraphy of the greater Barents Sea: Implications for petroleum systems. *Geological Society, London, Memoirs*, 35(1), 163–195. <https://doi.org/10.1144/M35.10>

Higgins, A.K. & Gilotti, J. & Smith, M.P. (2008). The Greenland Caledonides. Evolution of the Northeast Margin of Laurentia. *Geological Society of America Memoir*, 202. <https://doi.org/10.1130/MEM202>

Hinna, C. H. (2016). Seismic characterization of lower Cretaceous Cliniform packages in the Fingerdjupet sub-basin, southwestern Barents sea (MSc Thesis). University of Stavanger, Stavanger.

Hölttä, P., Heilimo, E., Huhma, H., Juopperi, H., Kontinen, A., Konnunaho, H., ... & Sorjonen-Ward, P. (2012). Archaean complexes of the Karelia Province in Finland. Geological Survey of Finland. Special Paper. 54. 7-20.

Hurum, J., Roberts, A., Dyke, G., Grundvåg, S-A., Nakrem, H., Midtkandal, I., .Olaussen, S. (2016). Bird or maniraptoran dinosaur? A femur from the Albian strata of Spitsbergen. *Palaeontologia Polonica*, 67, 137-147.

Indrevær, K., Gabrielsen, R. H., & Faleide, J. I. (2016). Early Cretaceous synrift uplift and tectonic inversion in the Loppa High area, southwestern Barents Sea, Norwegian shelf. *Journal of the Geological Society*, 174(2), 242–254. <https://doi.org/10.1144/jgs2016-066>



Indrevær, K., Gac, S., Gabrielsen, R. H., & Faleide, J. I. (2017). Crustal-scale subsidence and uplift caused by metamorphic phase changes in the lower crust: A model for the evolution of the Loppa High area, SW Barents Sea from late Paleozoic to Present. *Journal of the Geological Society*, 175(3), 497–508. <https://doi.org/10.1144/jgs2017-063>

Jakobsson, M., Mayer, L., Coakley, B., Dowdeswell, J. A., Forbes, S., Fridman, B., ... Weatherall, P. (2012). The International Bathymetric Chart of the Arctic Ocean (IBCAO) Version 3.0: IBCAO VERSION 3.0. *Geophysical Research Letters*, 39. <https://doi.org/10.1029/2012GL052219>

Johansson, Å., Gee, D. G., Björklund, L., & Witt-Nilsson, P. (1995). Isotope studies of granitoids from the Bangenhuk Formation, Ny Friesland Caledonides, Svalbard. *Geological Magazine*, 132(3), 303–320. <https://doi.org/10.1017/S0016756800013625>

Kairanov, B., Escalona, A., Mordasova, A., Śliwińska, K., & Suslova, A. (2018). Early Cretaceous tectonostratigraphic evolution of the north central Barents Sea. *Journal of Geodynamics*, 119, 183–198. <https://doi.org/10.1016/j.jog.2018.02.009>

Kalsbeek, F., Higgins, A. K., Jepsen, H. F., Frei, R., & Nutman, A. P. (2008). Granites and granites in the East Greenland Caledonides. In *Memoir 202: The Greenland Caledonides: Evolution of the Northeast Margin of Laurentia* (P. 227–249). Geological Society of America. [https://doi.org/10.1130/2008.1202\(09\)](https://doi.org/10.1130/2008.1202(09))

Kalsbeek, F., Nutman, A. P., Escher, J. C., Friderichsen, J. D., Hull, J. M., Jones, K. A., & Schack Pedersen, S. A. (1999). Geochronology of granitic and supracrustal rocks from the northern part of the East Greenland Caledonides: Ion microprobe U–Pb zircon ages. *GEUS Bulletin*, 31–48. <https://doi.org/10.34194/ggub.v184.5228>

Kirkland, C. L., Pease, V., Whitehouse, M. J., & Ineson, J. R. (2009). Provenance record from Mesoproterozoic-Cambrian sediments of Peary Land, North Greenland: Implications for the ice-covered Greenland Shield and Laurentian palaeogeography. *Precambrian Research*, 170(1–2), 43–60. <https://doi.org/10.1016/j.precamres.2008.11.006>

Klausen, T. G., Müller, R., Sláma, J., Olausson, S., Rismyhr, B., & Helland-Hansen, W. (2017). Depositional history of a condensed shallow marine reservoir succession: Stratigraphy and detrital zircon geochronology of the Jurassic Stø Formation, Barents Sea. *Journal of the Geological Society*, 175(1), 130–145. <https://doi.org/10.1144/jgs2017-024>

Korago, E. A., Kovaleva, G. N., Lopatin, B. G., & Orgo, V. V. (2004). The Precambrian rocks of Novaya Zemlya. *Geological Society, London, Memoirs*, 30(1), 135–143. <https://doi.org/10.1144/GSL.MEM.2004.030.01.12>

Kośmińska, K., Majka, J., Mazur, S., Krumbholz, M., Klonowska, I., Manecki, M., ... Dwornik, M. (2014). Blueschist facies metamorphism in Nordenskiöld Land of west-central Svalbard. *Terra Nova*, 26(5), 377–386. <https://doi.org/10.1111/ter.12110>

Köykkä, J., Lahtinen, R., & Huhma, H. (2019). Provenance evolution of the Paleoproterozoic metasedimentary cover sequences in northern Fennoscandia: Age distribution, geochemistry, and zircon morphology. *Precambrian Research*, 331, 105364. <https://doi.org/10.1016/j.precamres.2019.105364>

Lahtinen, R., Huhma, H., Kontinen, A., Kohonen, J., & Sorjonen-Ward, P. (2010). New constraints for the source characteristics, deposition and age of the 2.1–1.9Ga metasedimentary cover at the western margin of the Karelian Province. *Precambrian Research*, 176(1–4), 77–93. <https://doi.org/10.1016/j.precamres.2009.10.001>

LOCRA Final Report. (2017). LoCRA: Lower Cretaceous clastic wedges. An under-explored play in the Arctic. A multi-university collaboration. University of Stavanger and University Centre in Svalbard, Norway, internal report, 129.

Lorenz, H., Pystin, A. M., Olovyanishnikov, V. G., & Gee, D. G. (2004). Neoproterozoic high-grade metamorphism of the Kanin Peninsula, Timanide Orogen, northern Russia. *Geological Society, London, Memoirs*, 30(1), 59–68. <https://doi.org/10.1144/GSL.MEM.2004.030.01.06>

Ludwig, K. R. (1998). On the Treatment of Concordant Uranium-Lead Ages. *Geochimica et Cosmochimica Acta*, 62(4), 665–676. [https://doi.org/10.1016/S0016-7037\(98\)00059-3](https://doi.org/10.1016/S0016-7037(98)00059-3)

Maher, H. D. (2001). Manifestations of the Cretaceous High Arctic Large Igneous Province in Svalbard. *The Journal of Geology*, 109(1), 91–104. <https://doi.org/10.1086/317960>

Majka, J., Be'Eri-Shlevin, Y., Gee, D. G., Czerny, J., Frei, D., & Ladenberger, A. (2014). Torellian ( c. 640 Ma) metamorphic overprint of Tonian ( c. 950 Ma) basement in the Caledonides of southwestern Svalbard. *Geological Magazine*, 151(4), 732–748. <https://doi.org/10.1017/S0016756813000794>

Majka, J., Czerny, J., Mazur, S., Holm, D. K., & Manecki, M. (2010). Neoproterozoic metamorphic evolution of the Isbjørnhamna Group rocks from south-western Svalbard: Neoproterozoic metamorphism in southern Svalbard. *Polar Research*, 29(3), 250–264. <https://doi.org/10.1111/j.1751-8369.2010.00186.x>

Malone, S. J., McClelland, W. C., von Gosen, W., & Piepjohn, K. (2017). The earliest Neoproterozoic magmatic record of the Pearya terrane, Canadian high Arctic: Implications for Caledonian terrane reconstructions. *Precambrian Research*, 292, 323–349. <https://doi.org/10.1016/j.precamres.2017.01.006>

Marín, D., Escalona, A., Grundvåg, S.-A., Olaussen, S., Sandvik, S., & Śliwińska, K. K. (2018). Unravelling key controls on the rift climax to post-rift fill of marine rift basins: Insights from 3D seismic analysis of the Lower Cretaceous of the Hammerfest Basin, SW Barents Sea. *Basin Research*, 30(4), 587–612. <https://doi.org/10.1111/bre.12266>

Marin, D., Escalona, A., Śliwińska, K. K., Nøhr-Hansen, H., & Mordasova, A. (2017). Sequence stratigraphy and lateral variability of Lower Cretaceous clinoforms in the southwestern Barents Sea. *AAPG Bulletin*, 101(09), 1487–1517. <https://doi.org/10.1306/10241616010>

McClelland, W. C., von Gosen, W., & Piepjohn, K. (2019). Tonian and Silurian magmatism in Nordaustlandet: Svalbard's place in the Caledonian orogen. In Piepjohn K., Strauss, J.V., Reinhardt, L. & McClelland, W.C. *Circum-Arctic Structural Events: Tectonic Evolution of the Arctic Margins and Trans-Arctic Links with Adjacent Orogens*. Geological Society of America. [https://doi.org/10.1130/2018.2541\(04\)](https://doi.org/10.1130/2018.2541(04))

Midtkandal, I., & Nystuen, J. P. (2009). Depositional architecture of a low-gradient ramp shelf in an epicontinental sea: The lower Cretaceous of Svalbard. *Basin Research*, 21(5), 655–675. <https://doi.org/10.1111/j.1365-2117.2009.00399.x>

Midtkandal, I., Nystuen, J.P., Nagy, J. & Mørk, A., 2008. Lower Cretaceous lithostratigraphy across a regional subaerial unconformity in Spitsbergen: the Rurikfjellet and Helvetiafjellet formations. *Norwegian Journal of Geology*, 88, 287-304.

Midtkandal, I., Faleide, T. S., Faleide, J. I., Planke, S., Anell, I., & Nystuen, J. P. (2019). Nested intrashelf platform clinoforms—Evidence of shelf platform growth exemplified by Lower Cretaceous strata in the Barents Sea. *Basin Research*, 32(2), 216–223. <https://doi.org/10.1111/bre.12377>

Midtkandal, I., Svensen, H. H., Planke, S., Corfu, F., Polteau, S., Torsvik, T. H., ... Olausen, S. (2016). The Aptian (Early Cretaceous) oceanic anoxic event (OAE1a) in Svalbard, Barents Sea, and the absolute age of the Barremian-Aptian boundary. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 463, 126–135. <https://doi.org/10.1016/j.palaeo.2016.09.023>

Midtkandal, I., Nystuen, J.P., Nagy, J., Mørk, A., (2008). Lower Cretaceous lithostratigraphy across a regional subaerial unconformity in Spitsbergen: the Rurikfjellet and Helvetiafjellet formations. *Norwegian Journal of Geology*. 88, 287-304.

Mørk, A., Dallmann, A., Dypvik, H., Johannessen, E., Larssen, G., Nøttvedt, N., ... Worsley, D. (1999), Mesozoic lithostratigraphy, in W. K. Dallmann, ed., Lithostratigraphic lexicon of Svalbard. Upper Palaeozoic to Quaternary bedrock. Review and recommendations for nomenclature use: Tromsø, Norway, Norsk Polarinstitut, 127–214.

Norwegian Petroleum Directorate (NPD). (2020). Factpages (online). Available 4.05.2020.

Nøhr-Hansen, H., Piasecki, S., & Alsen, P. (2019). A Cretaceous dinoflagellate cyst zonation for NE Greenland. *Geological Magazine*, 1–35. <https://doi.org/10.1017/S0016756819001043>

Onderdonk, N., & Midtkandal, I. (2010). Mechanisms of collapse of the cretaceous helvetiafjellet formation at Kvalvågen, eastern Spitsbergen. *Marine and Petroleum Geology*, 27(10), 2118–2140. <https://doi.org/10.1016/j.marpetgeo.2010.09.004>

Patruno, S., Hampson, G. J., & Jackson, C. A.-L. (2015). Quantitative characterisation of deltaic and subaqueous clinoforms. *Earth-Science Reviews*, 142, 79–119. <https://doi.org/10.1016/j.earscirev.2015.01.004>

Patruno, S., & Helland-Hansen, W. (2018). Clinoforms and clinoform systems: Review and dynamic classification scheme for shorelines, subaqueous deltas, shelf edges and continental margins. *Earth-Science Reviews*, 185, 202–233. <https://doi.org/10.1016/j.earscirev.2018.05.016>

Pease, V. (2011). Chapter 20 Eurasian orogens and Arctic tectonics: An overview. *Geological Society, London, Memoirs*, 35(1), 311–324. <https://doi.org/10.1144/M35.20>

Pease, V., Drachev, S., Stephenson, R., & Zhang, X. (2014). Arctic lithosphere—A review. *Tectonophysics*, 628, 1–25. <https://doi.org/10.1016/j.tecto.2014.05.033>

Pettersson, C. H., Tebenkov, A. M., Larionov, A. N., Andresen, A., & Pease, V. (2009). Timing of migmatization and granite genesis in the Northwestern Terrane of Svalbard, Norway: Implications for regional correlations in the Arctic Caledonides. *Journal of the Geological Society*, 166(1), 147–158. <https://doi.org/10.1144/0016-76492008-023>

Petrov, O.V., Sobolev, N.N., Koren, T.N., Vasiliev, V.E., Petrov, E.O., Larssen, G.B. & Smelror, M. (2008). Palaeozoic and Early Mesozoic evolution of the East Barents and Kara Seas sedimentary basins. *Norwegian Journal of Geology*, 88, 227-234.

Petrov, O., Morozov, A., Shokalsky, S., Kashubin, S., Artemieva, I., Sobolev, N., ... Smelror, M. (2016). Crustal structure and tectonic model of the Arctic region. *Earth-Science Reviews*, 154, 29-71.

Piepjohn, K., & von Gosen, W. (2017). Structural transect through Ellesmere Island (Canadian Arctic): Superimposed Palaeozoic Ellesmerian and Cenozoic Eureka deformation. *Geological Society, London, Special Publications*, 460(1), 33–56. <https://doi.org/10.1144/SP460.5>

Porębski, S. J., & Steel, R. J. (2003). Shelf-margin deltas: Their stratigraphic significance and relation to deepwater sands. *Earth-Science Reviews*, 62(3–4), 283–326. [https://doi.org/10.1016/S0012-8252\(02\)00161-7](https://doi.org/10.1016/S0012-8252(02)00161-7)

Ritzmann, O., & Faleide, J. I. (2009). The crust and mantle lithosphere in the Barents Sea/Kara Sea region. *Tectonophysics*, 470(1–2), 89–104. <https://doi.org/10.1016/j.tecto.2008.06.018>

Roberts, D., & Gee, D.G. (1985). An introduction to the structure of the Scandinavian Caledonides, in Gee, D.G., and Sturt, B.A., editors., *The Caledonide Orogen—Scandinavia and related areas* (pp.55–68). Chichester: John Wiley & Sons.

Røhr, T. S. & Andersen, T. (2009). Detrital zircons from the high Arctic; evidence for extensive recycling of sediment from Devonian through Mesozoic times. In: Røhr, T. S. (ed.) *Sedimentary Provenance Analysis of Lower Cretaceous Sedimentary Successions in The Arctic; Constraints From Detrital Zircon data*. PhD thesis, Faculty of Mathematics and Natural Sciences, University of Oslo, Oslo, 55–105.

Røhr, T. S., Andersen, T., & Dypvik, H. (2008). Provenance of Lower Cretaceous sediments in the Wandel Sea Basin, North Greenland. *Journal of the Geological Society*, 165(3), 755–767. <https://doi.org/10.1144/0016-76492007-102>

Røhr, T. S., Andersen, T., Dypvik, H., & Embry, A. F. (2010). Detrital zircon characteristics of the Lower Cretaceous Isachsen Formation, Sverdrup Basin: Source constraints from age and Hf isotope data. *Canadian Journal of Earth Sciences*, 47(3), 255–271. <https://doi.org/10.1139/E10-006>

Sandelin, S., Tebenkov, A. M., & Gee, D. G. (2001). The stratigraphy of the lower part of the Neoproterozoic Murchisonfjorden Supergroup in Nordaustlandet, Svalbard. *GFF*, 123(2), 113–127. <https://doi.org/10.1080/11035890101232113>

Seldal, J. (2005). Lower Cretaceous: The next target for oil exploration in the Barents Sea? *Geological Society, Petroleum Geology Conference Series*, 6(1), 231–240. <https://doi.org/10.1144/0060231>

Serck, C. S., Faleide, J. I., Braathen, A., Kjølhamar, B., & Escalona, A. (2017). Jurassic to Early Cretaceous basin configuration(s) in the Fingerdjupet Subbasin, SW Barents Sea. *Marine and Petroleum Geology*, 86, 874–891. <https://doi.org/10.1016/j.marpetgeo.2017.06.044>

Smelror, M., Mørk, A., Monteil, E., Rutledge, D., & Leereveld, H. (1998). The Klippfisk Formation - a new lithostratigraphic unit of Lower Cretaceous platform carbonates on the Western Barents Shelf. *Polar Research*, 17(2), 181–202. <https://doi.org/10.1111/j.1751-8369.1998.tb00271.x>

Smelror, M., Petrov, O., Larssen, G-B & Werner, S.C.. (2009). ATLAS: Geological History of the Barents Sea. Geological Survey of Norway.

Smith, M. P., & Rasmussen, J. A. (2008). Cambrian–Silurian development of the Laurentian margin of the Iapetus Ocean in Greenland and related areas. I *Memoir 202: The Greenland Caledonides: Evolution of the Northeast Margin of Laurentia* (pp.137–167). Geological Society of America. [https://doi.org/10.1130/2008.1202\(06\)](https://doi.org/10.1130/2008.1202(06))

Steel, R., Mellere, D., Plink-Bjorklund, P., Crabaugh, Deibert, J., Loeseth & Shellpeper. (2000). Deltas vs. Rivers on the Shelf Edge: Their Relative Contributions to the Growth of Shelf-Margins and Basin-Floor Fans (Barremian and Eocene, Spitsbergen). GCSSEPM Special Publication. 20. 981-1009. [10.5724/gcs.00.15.0981](https://doi.org/10.5724/gcs.00.15.0981).

Torsvik, T.H., Carlos, D., Mosar, J., Cocks, L.R.M. & Malme, T. (2002). Global reconstructions and North Atlantic palaeogeography 400 Ma to Recent. In: Eide, E.A. (coord.). BATLAS – Mid Norway plate reconstructions atlas with global and Atlantic perspectives. Geological Survey of Norway, 18-39.

Torsvik, T. H., Van der Voo, R., Preeden, U., Mac Niocaill, C., Steinberger, B., Doubrovine, P. V., ... Cocks, L. R. M. (2012). Phanerozoic polar wander, palaeogeography and dynamics. *Earth-Science Reviews*, 114(3–4), 325–368. <https://doi.org/10.1016/j.earscirev.2012.06.007>

Trettin, H. P. (Red.). (1991). Geology of the Innuitian Orogen and Arctic Platform of Canada and Greenland. *Geological Society of America*, 3, 569. <https://doi.org/10.1130/DNAG-GNA-E>

Vickers, M. L., Price, G. D., Jerrett, R. M., & Watkinson, M. (2016). Stratigraphic and geochemical expression of Barremian–Aptian global climate change in Arctic Svalbard. *Geosphere*, 12(5), 1594–1605. <https://doi.org/10.1130/GES01344>.

Vigran, J. O. (2014). *Palynology and geology of the Triassic succession of Svalbard and the Barents Sea*. Norges geologiske undersøkelse.

Witt-Nilsson, P., Gee, D. G. & Hellman, F. J. (1998). Tectonostratigraphy of the Caledonian. Atomfjella Antiform of northern Ny Friesland, Svalbard. *Norwegian Journal of Geology*, 78, 67-80. Oslo.

Worsley, D. (2008). The post-Caledonian development of Svalbard and the western Barents Sea. *Polar Research*, 27(3), 298–317. <https://doi.org/10.1111/j.1751-8369.2008.00085.x>

Worsley, D., Agdestein, T., Gjelberg, J., Kirkemo, K., Mørk, A., Nilsson, I., ... Nilsson, A. (2001). The geological evolution of Bjørnøya, Arctic Norway: Implications for the Barents Shelf. *Norwegian Journal of Geology*, 81, 195-234.

Århus, N., Kelly, S. R. A., Collins, J. S. H., & Sandy, M. R. (1990). Systematic palaeontology and biostratigraphy of two Early Cretaceous condensed sections from the Barents Sea. *Polar Research*, 8 (2), 165–194. <https://doi.org/10.1111/j.1751-8369.1990.tb00383.x>